



Biogeochemical cycles in the Baltic Sea

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Climate of the Baltic Sea Region
International Baltic Earth Summer School
20-27 August 2018



The National Centre
for Research and Development



BONUS
SCIENCE FOR A BETTER FUTURE OF THE BALTIC SEA REGION



NATIONAL SCIENCE CENTRE
POLAND

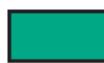
Outline:

- I. Carbon – the chemical element conditioning climate
- II. Role of marine environment in the carbon cycle
- III. Marine CO₂ system and ocean acidification
- IV. Peculiarities of the Baltic Sea

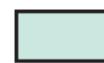
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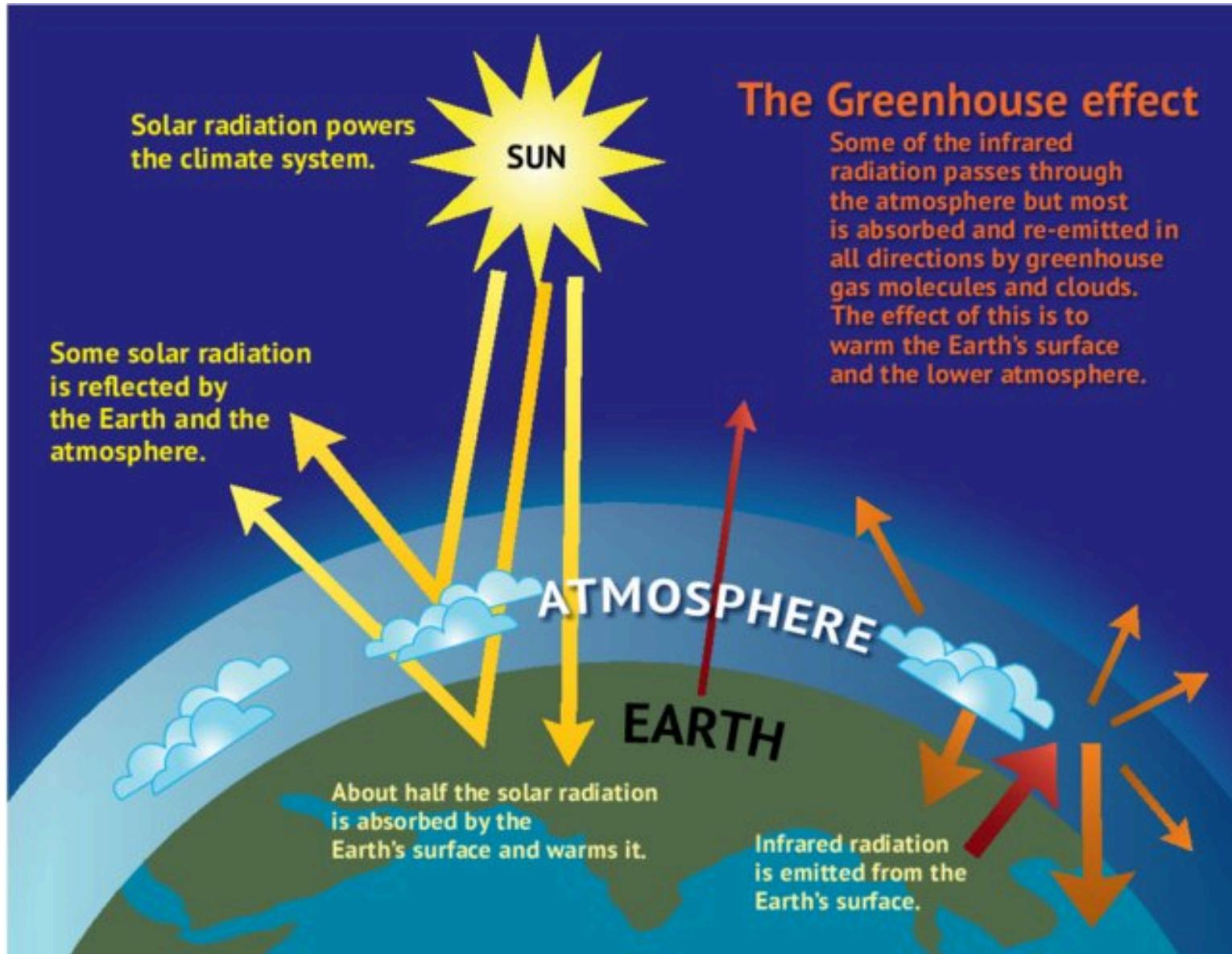
H																					He
Li	Be																				Ne
Na	Mg																				Ar
K	Ca	Sc	Ti	V	Cr	Mn	Fe	Co	Ni	Cu	Zn	Ga	Ge	As	Se	Br					Kr
R	Sr	Y	Zr	Nb	Mo	Tc	Ru	Rh	Pd	Ag	Cd	In	Sn	S	Te	I					Xe
Cs	Ba	La	Hf	Ta	W	Re	Os	Ir	Pt	Au	Hg	Tl	Pb	Bi	Po	At					Rn

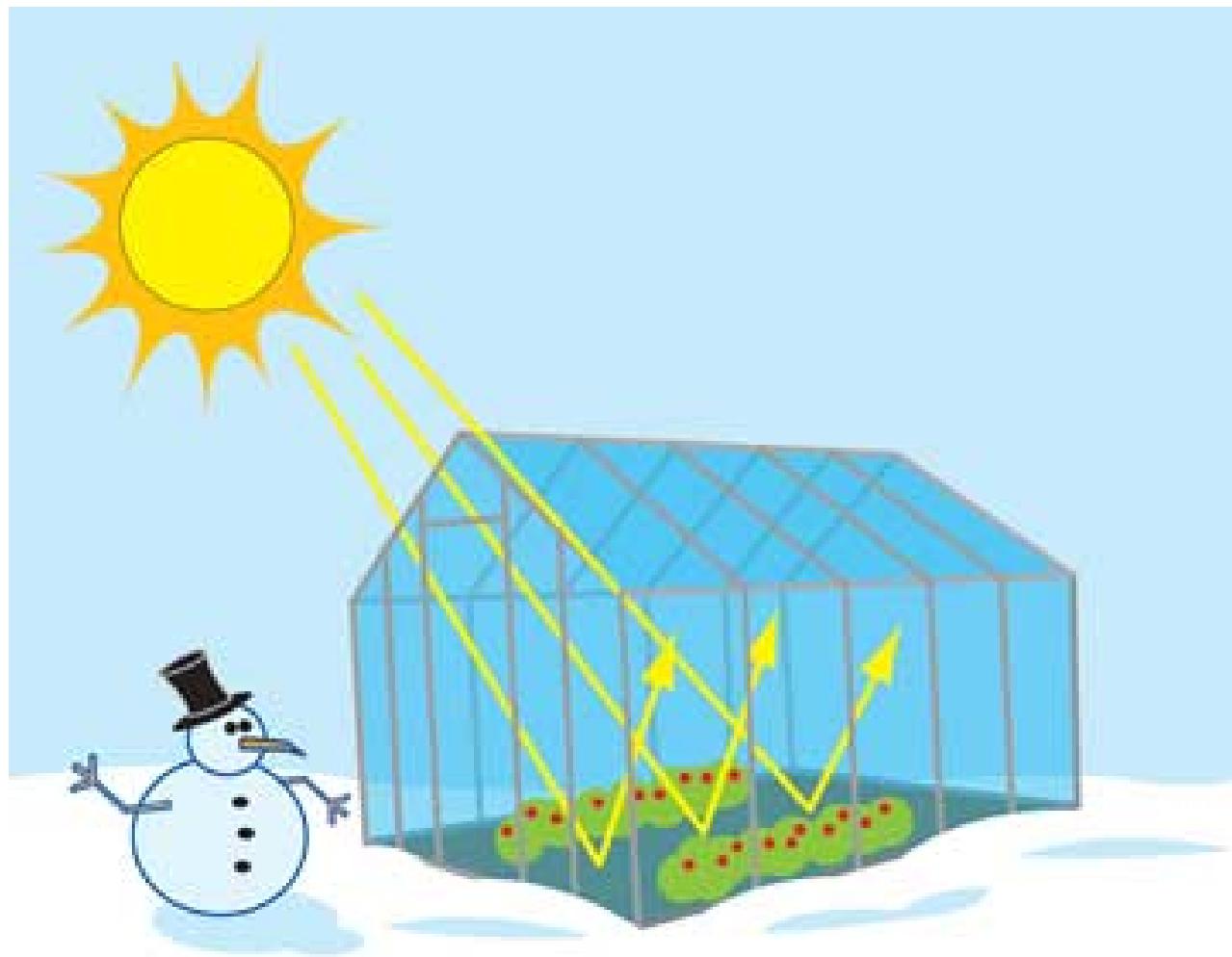
 Essential to all animals and plants

 Essential to several classes of animals and plants

 Believed essential to a variety of species

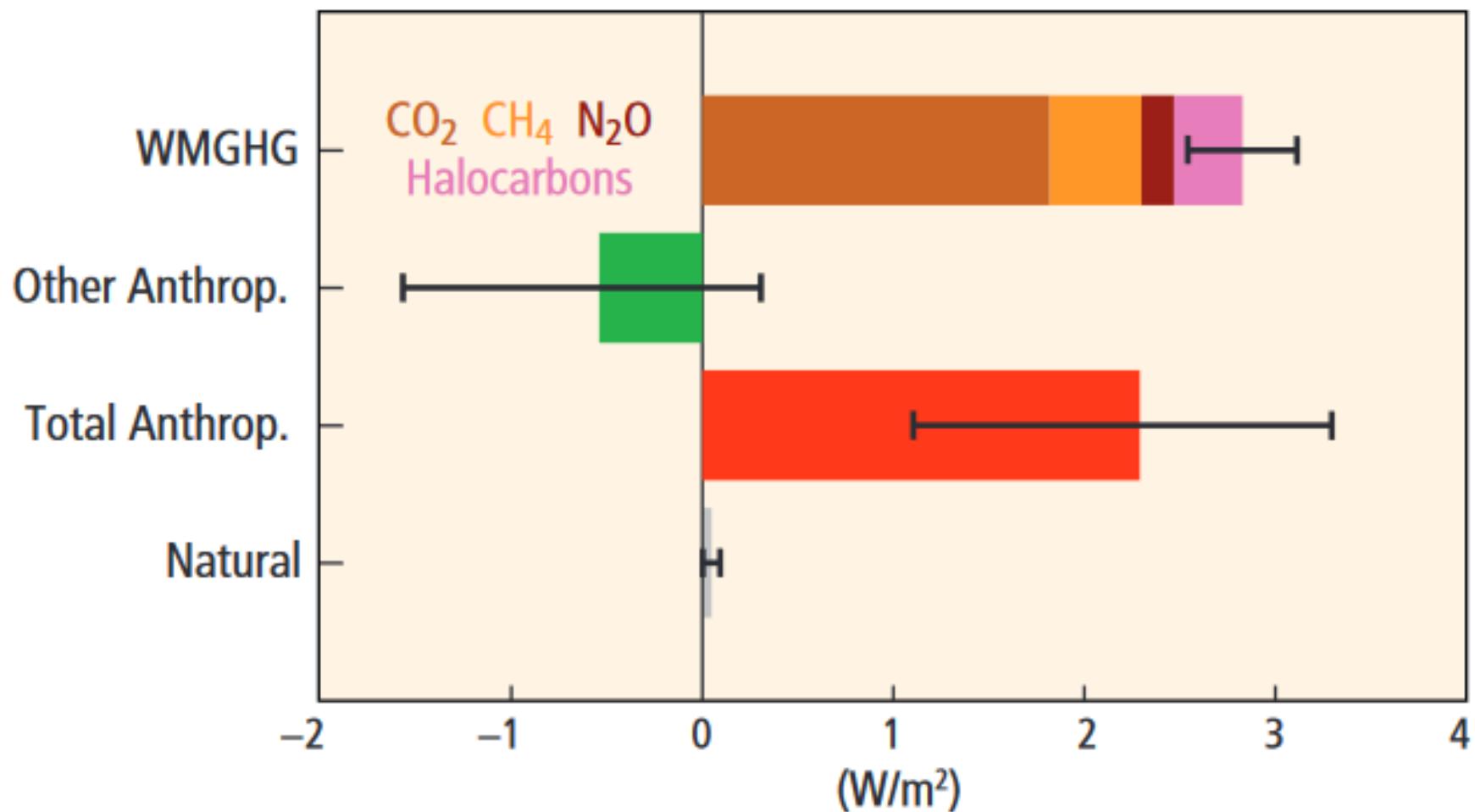
 Possible essential trace elements for some species



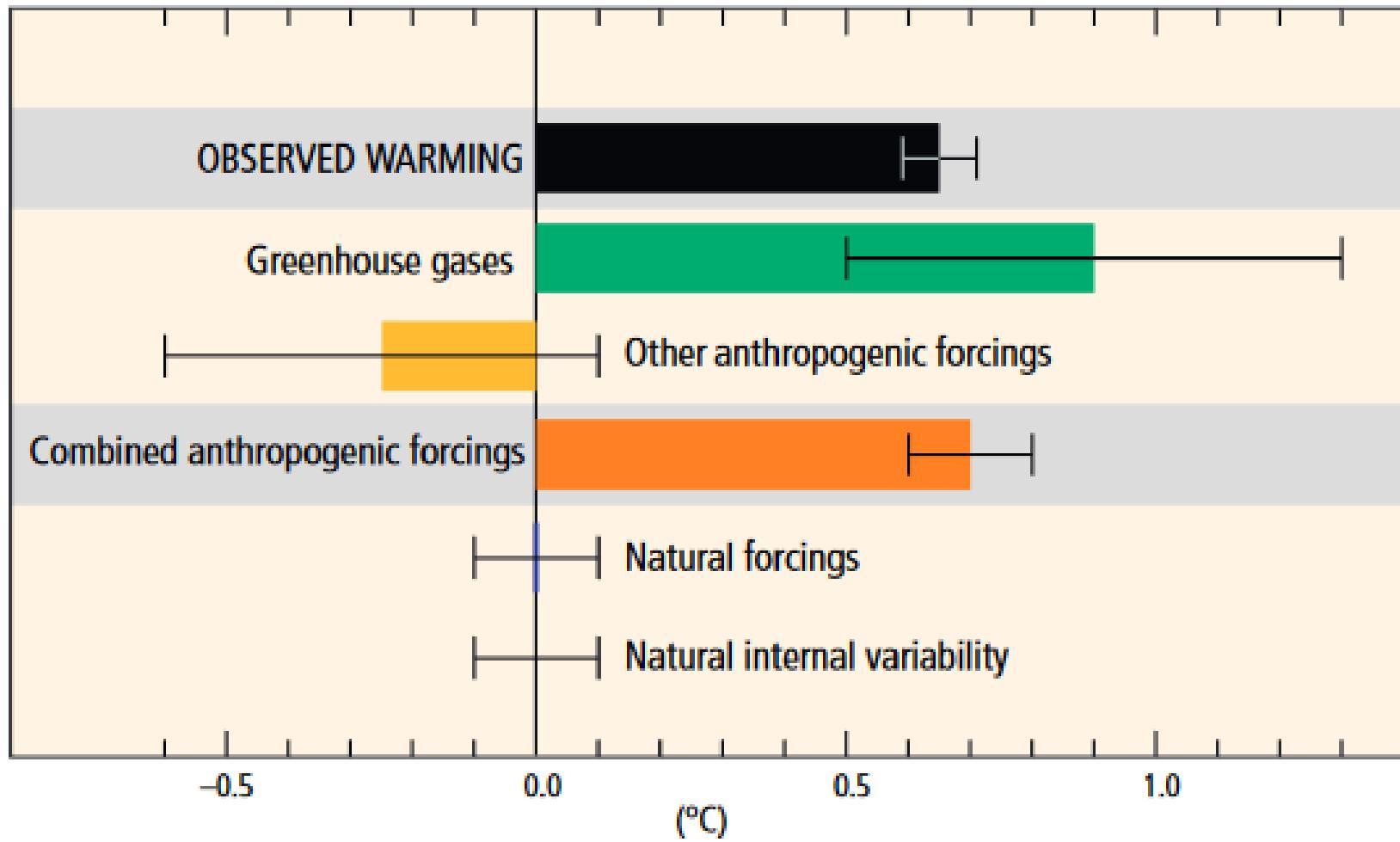


Source: NASA, Climate Kids

Radiative forcing in 2011 relative to 1750

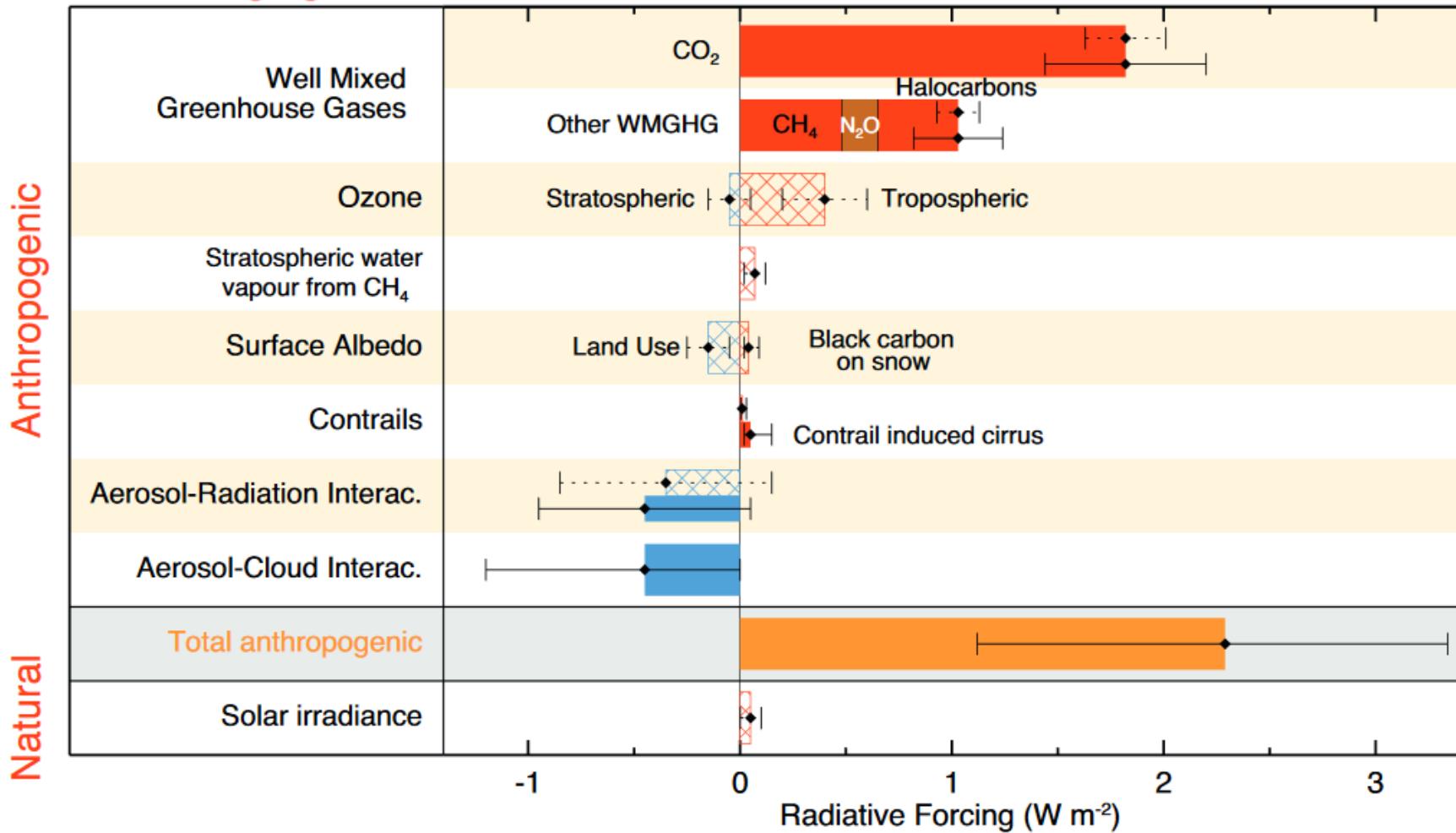


Contributions to observed surface temperature change over the period 1951–2010



Radiative forcing of climate between 1750 and 2011

Forcing agent



Greenhouse effect and greenhouse gases – good or bad?



Without greenhouse effect, the Earth would be **extremely cold**.



National Aeronautics and Space Administration



WATER VAPOR



Visit climatekids.nasa.gov

H₂O

This is water in gas form, like steam above a boiling pot or water evaporating off a lake. It forms clouds and rains back on Earth. This can cause a cooling effect.



National Aeronautics and Space Administration



CARBON DIOXIDE



Visit climatekids.nasa.gov

CO₂

Made up of carbon and oxygen, CO₂ is all around us naturally. It comes from decaying and living organisms, and from volcanoes.



National Aeronautics and Space Administration



OZONE



Visit climatekids.nasa.gov

O₃

Up in the atmosphere where the planes fly, the ozone layer blocks the sun's radiation, which helps protect us from the powerful rays.



National Aeronautics and Space Administration



METHANE



Visit climatekids.nasa.gov

CH₄

Methane, made of carbon and hydrogen, is a normal gas released from wetlands, growing rice, raising cattle, using natural gas, and mining coal.



National Aeronautics and Space Administration



NITROUS OXIDE



Visit climatekids.nasa.gov

N₂O

Nitrous oxide is a natural part of the nitrogen cycle. Bacteria in soil and the ocean make it.



National Aeronautics and Space Administration



CHLOROFLUOROCARBONS



Visit Climatekids.nasa.gov

CFCs

Fluorinated gases are not created in nature. They damage the protective ozone layer and are powerful greenhouse gases.



WATER VAPOR



Water vapor blocks heat from escaping, so it gets warmer. That makes even more water evaporate. Once this process happens, it can happen again more easily.



CARBON DIOXIDE



CO_2 is released when burning fossil fuels like coal and oil. It's the most important contributor to human-caused global warming.



OZONE



Close to the ground, ozone acts as a greenhouse gas and can be formed by burning gas in cars and factories.



METHANE



It traps a lot of heat. Scientists consider it the second most important contributor to human-caused global warming of all the greenhouse gases.



NITROUS OXIDE



N_2O is released by some types of factories, power plants, and plant fertilizer. It damages the protective ozone layer and is a powerful greenhouse gas.



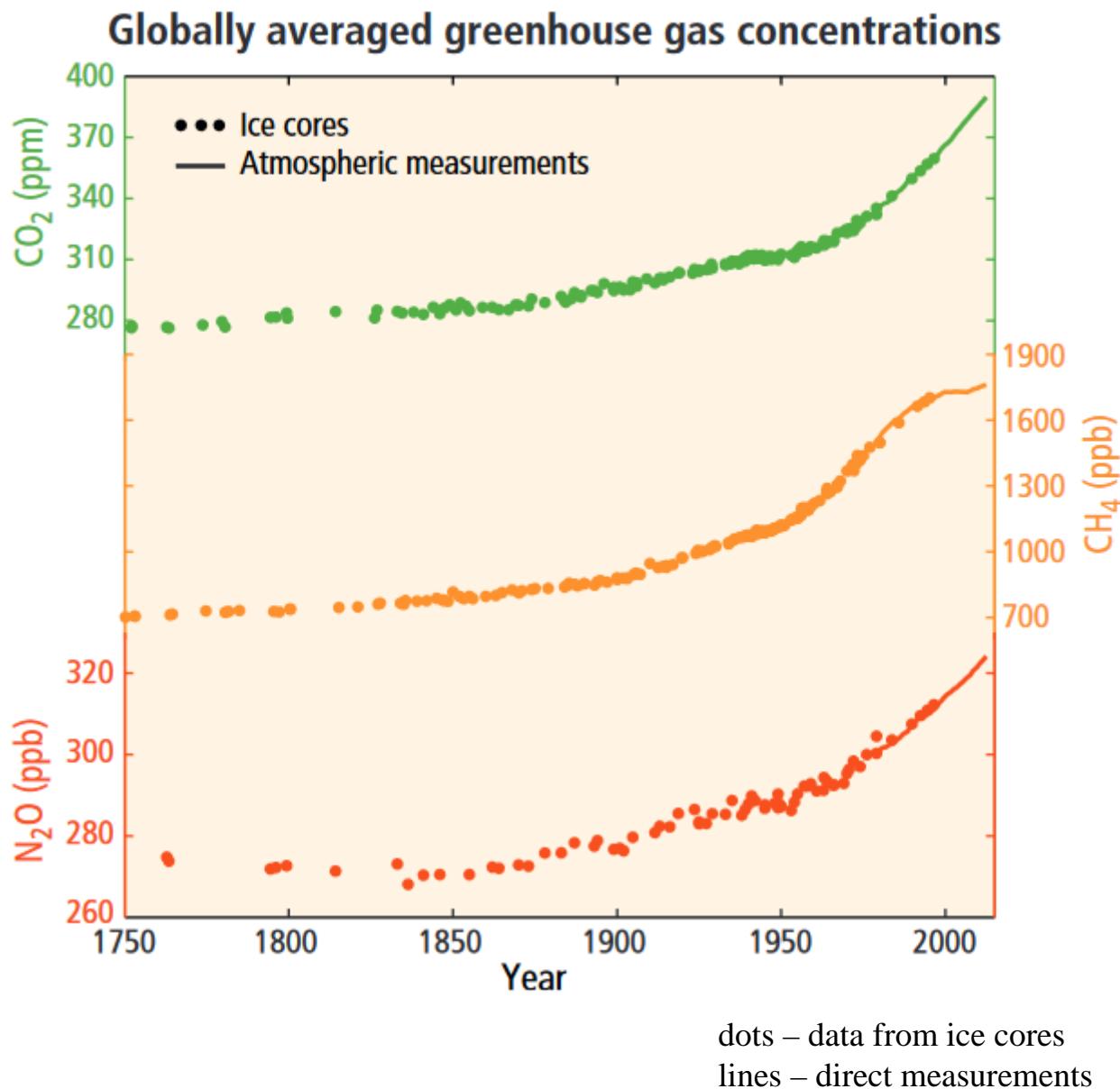
CHLOROFLUOROCARbons



You probably shouldn't have created me.

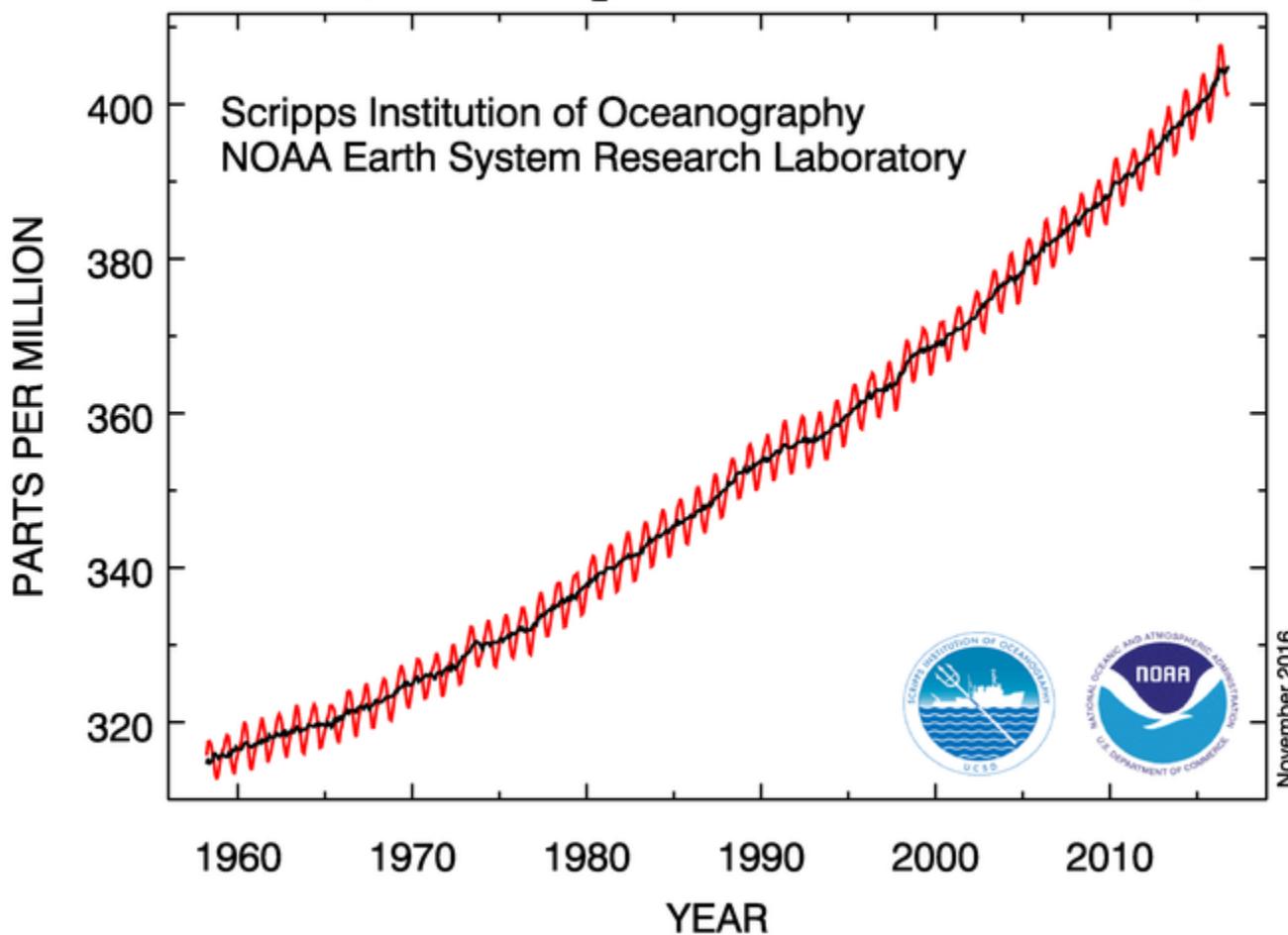


<https://climatekids.nasa.gov/time-machine/>

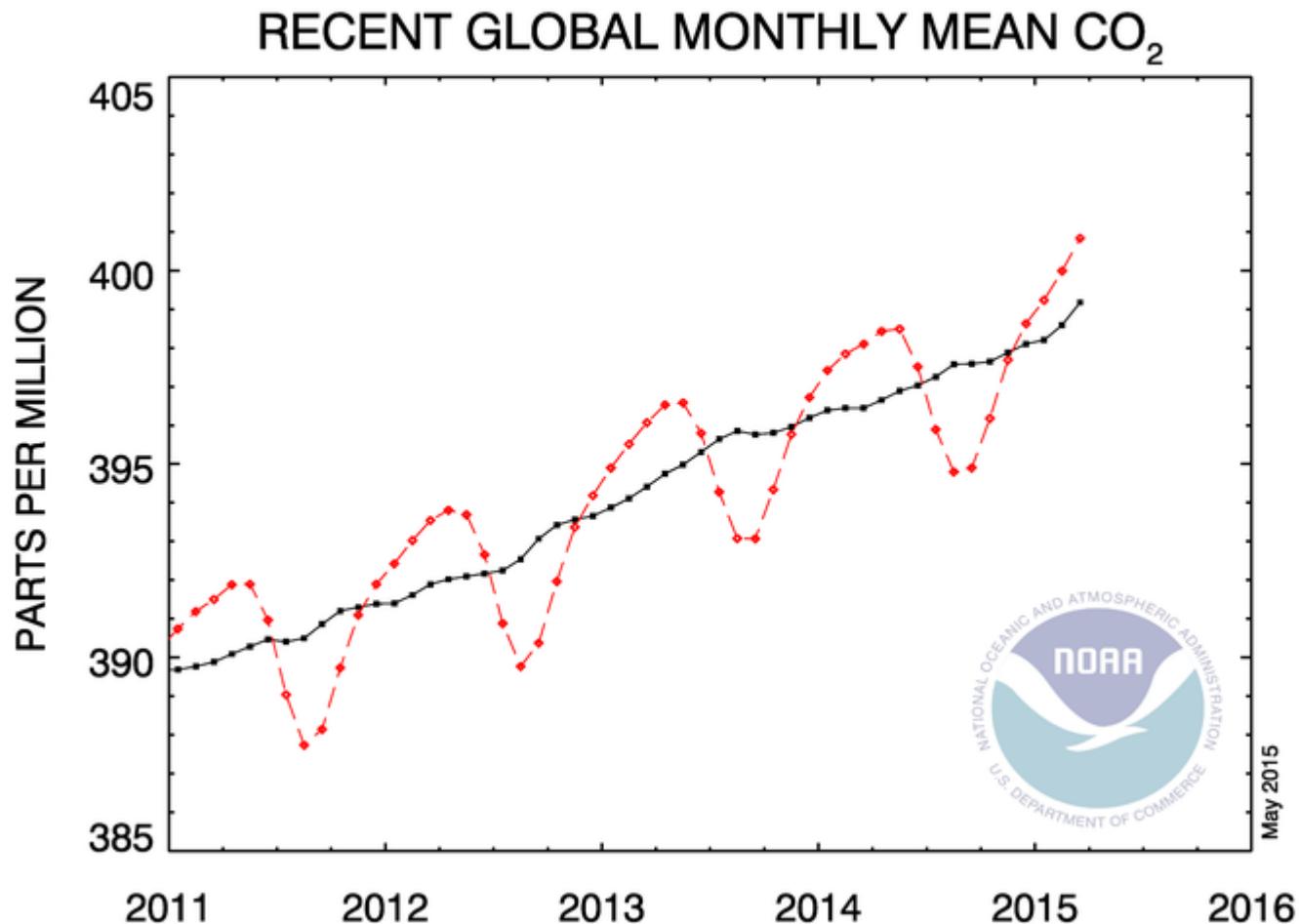


Source: IPCC, 2014

Atmospheric CO₂ at Mauna Loa Observatory



Source: NOAA

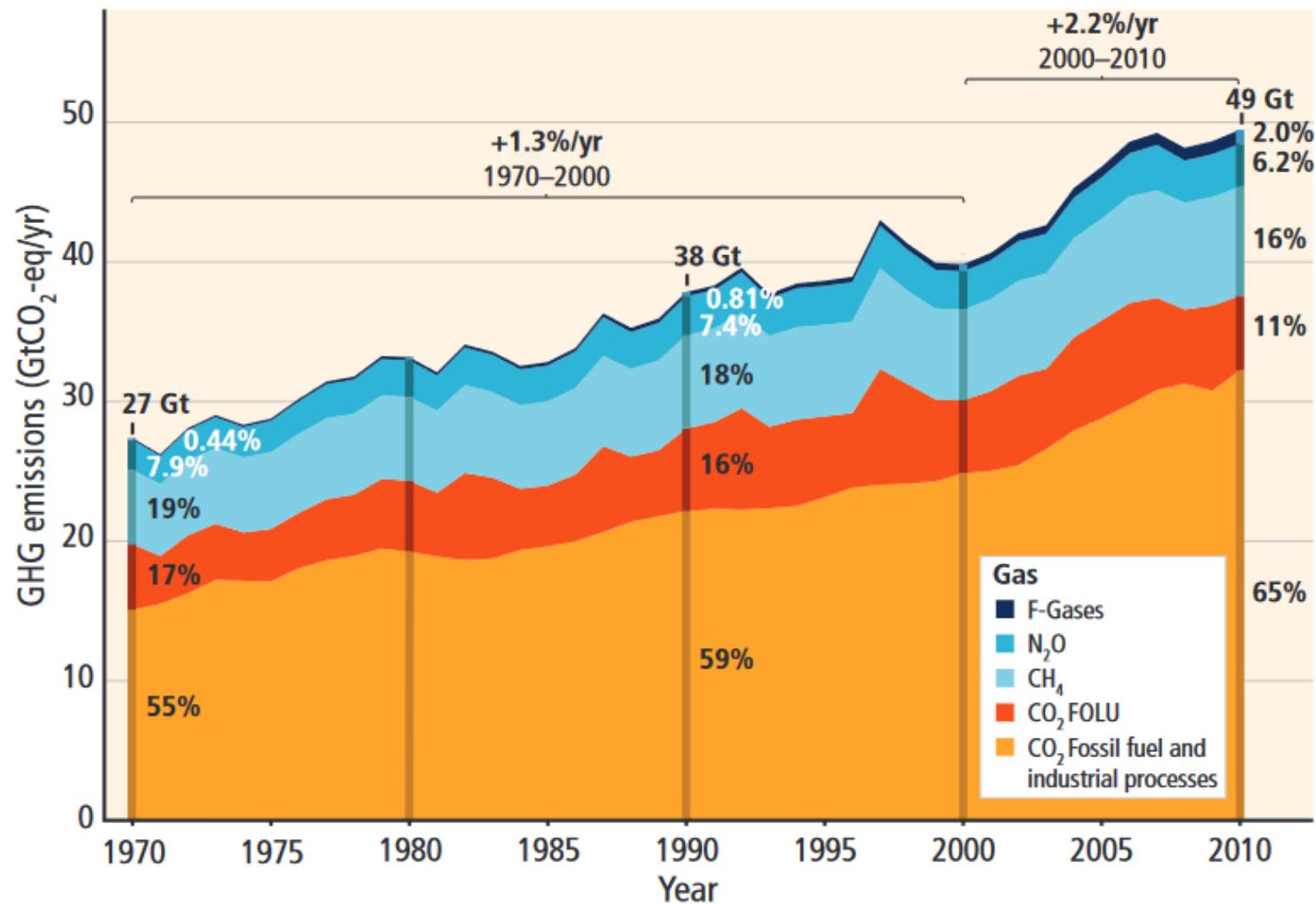


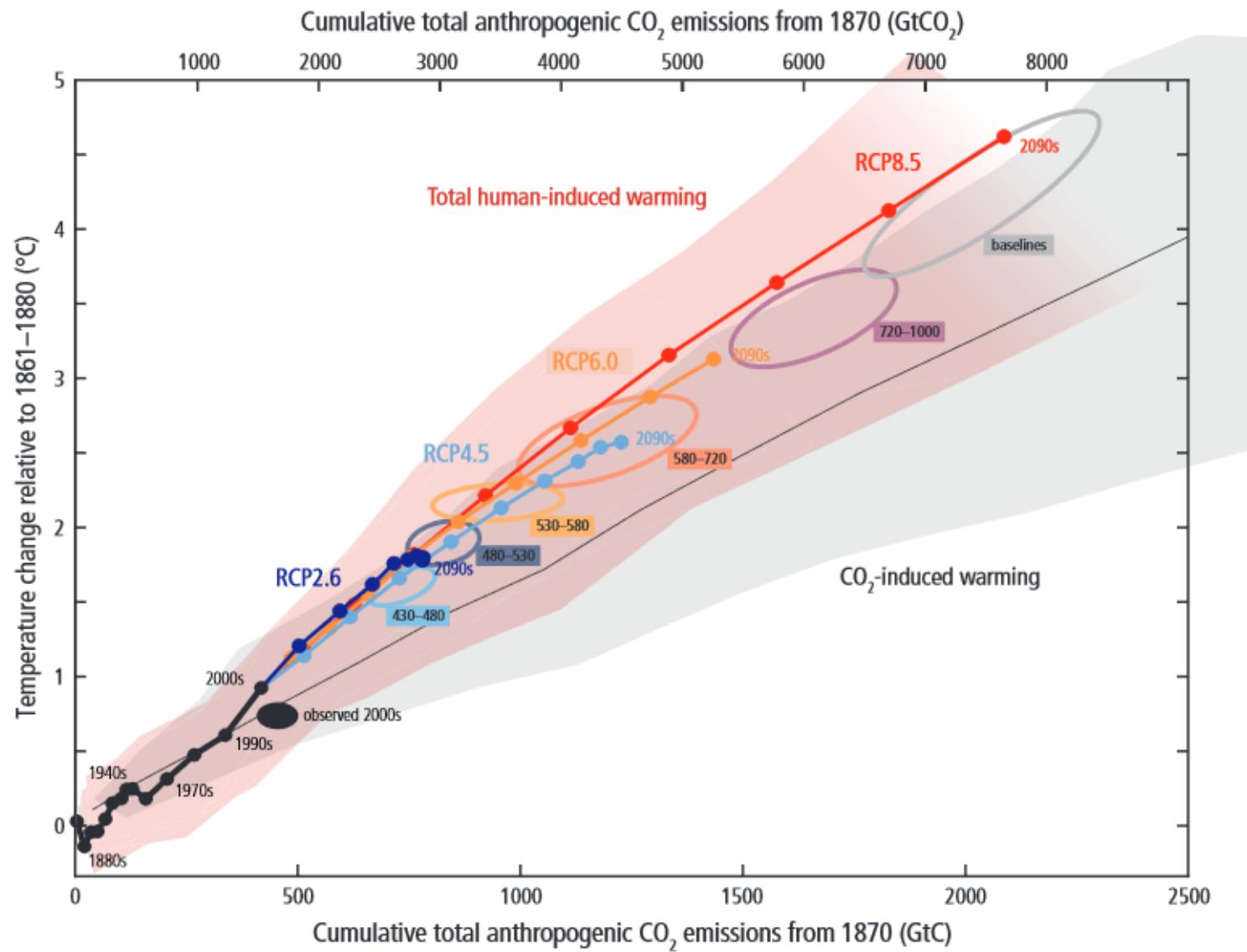
Detailed analysis show that the slope increases

<https://www.youtube.com/watch?v=x1SgmFa0r04>

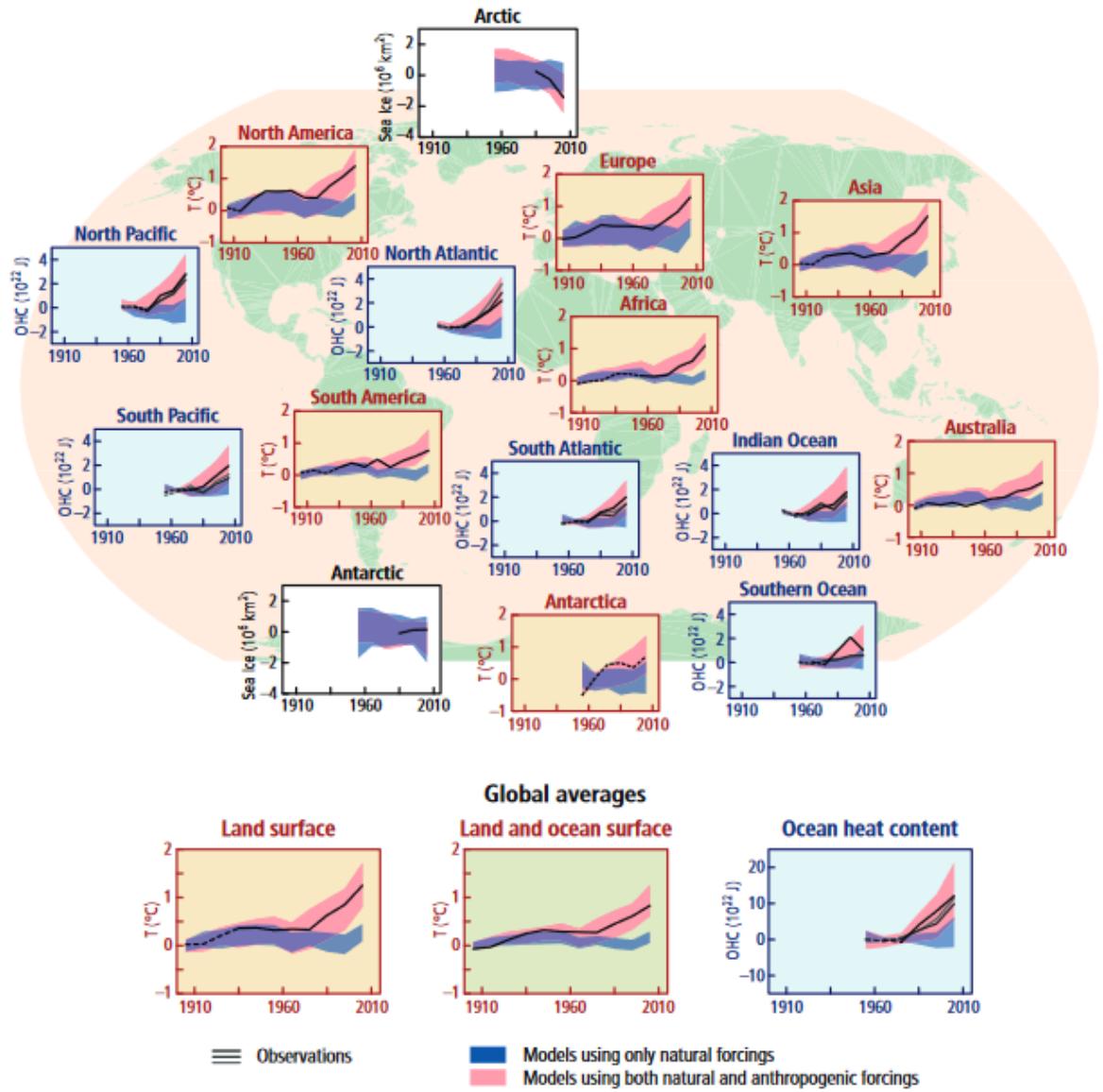
<https://www.youtube.com/watch?v=SHzRBMBVu-4>

Total annual anthropogenic GHG emissions by gases 1970–2010

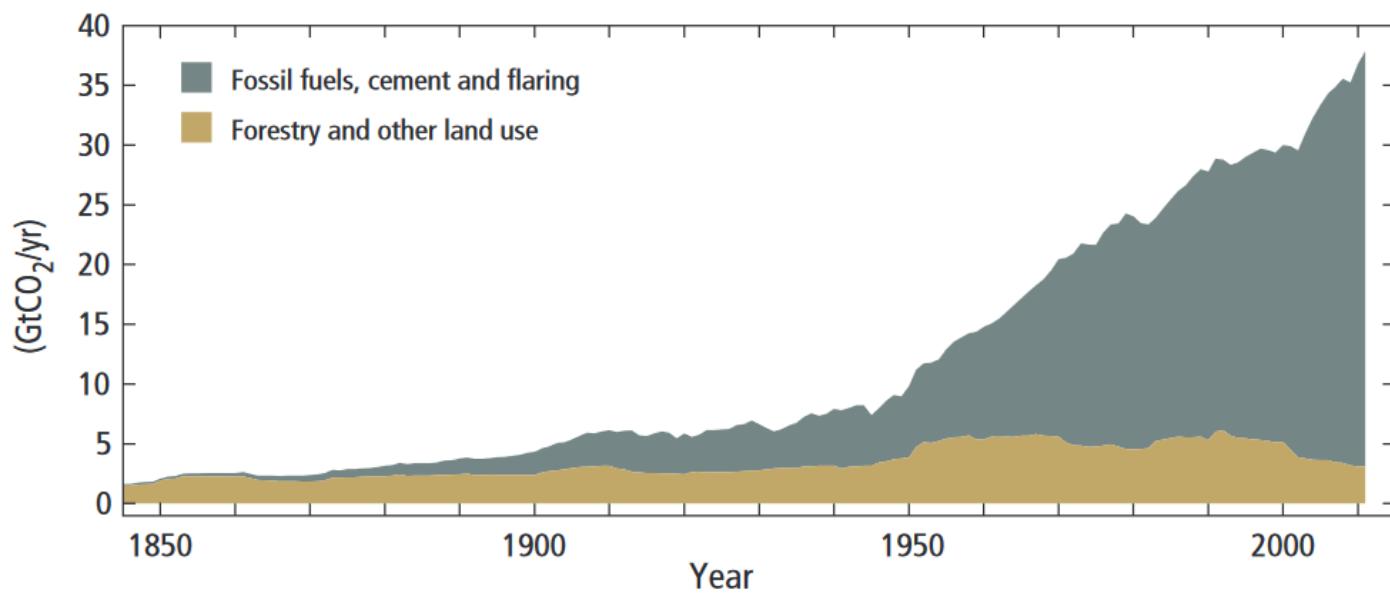




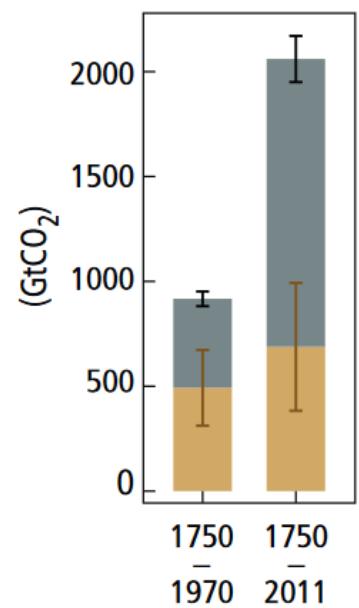
Source: IPCC, 2014



Global anthropogenic CO₂ emissions



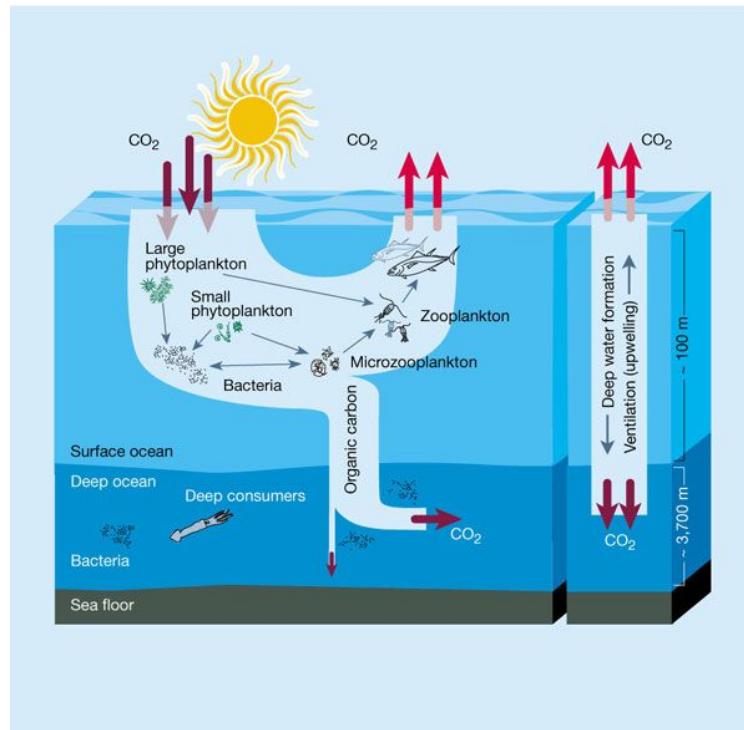
Cumulative CO₂ emissions



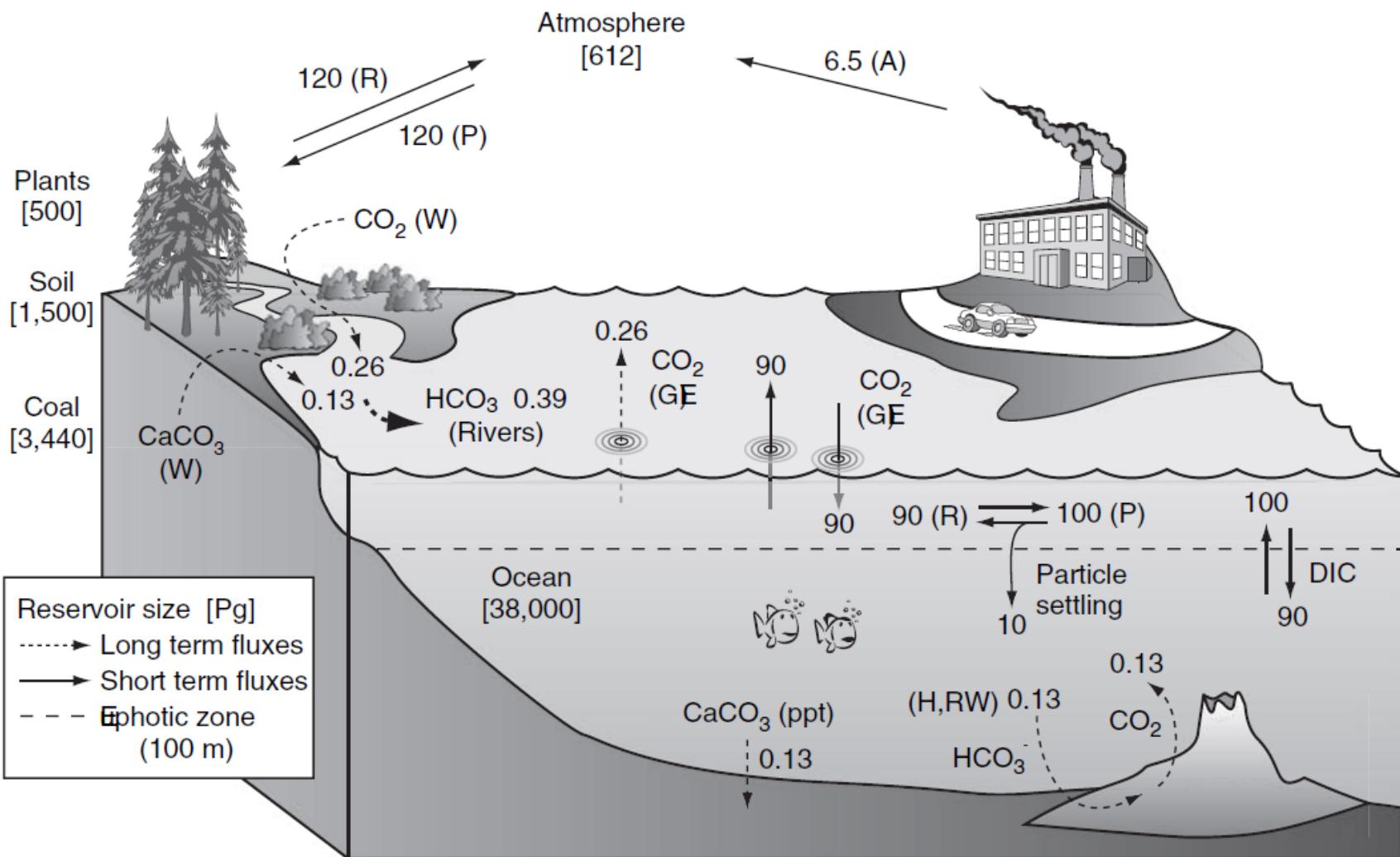
Source: IPCC, 2014

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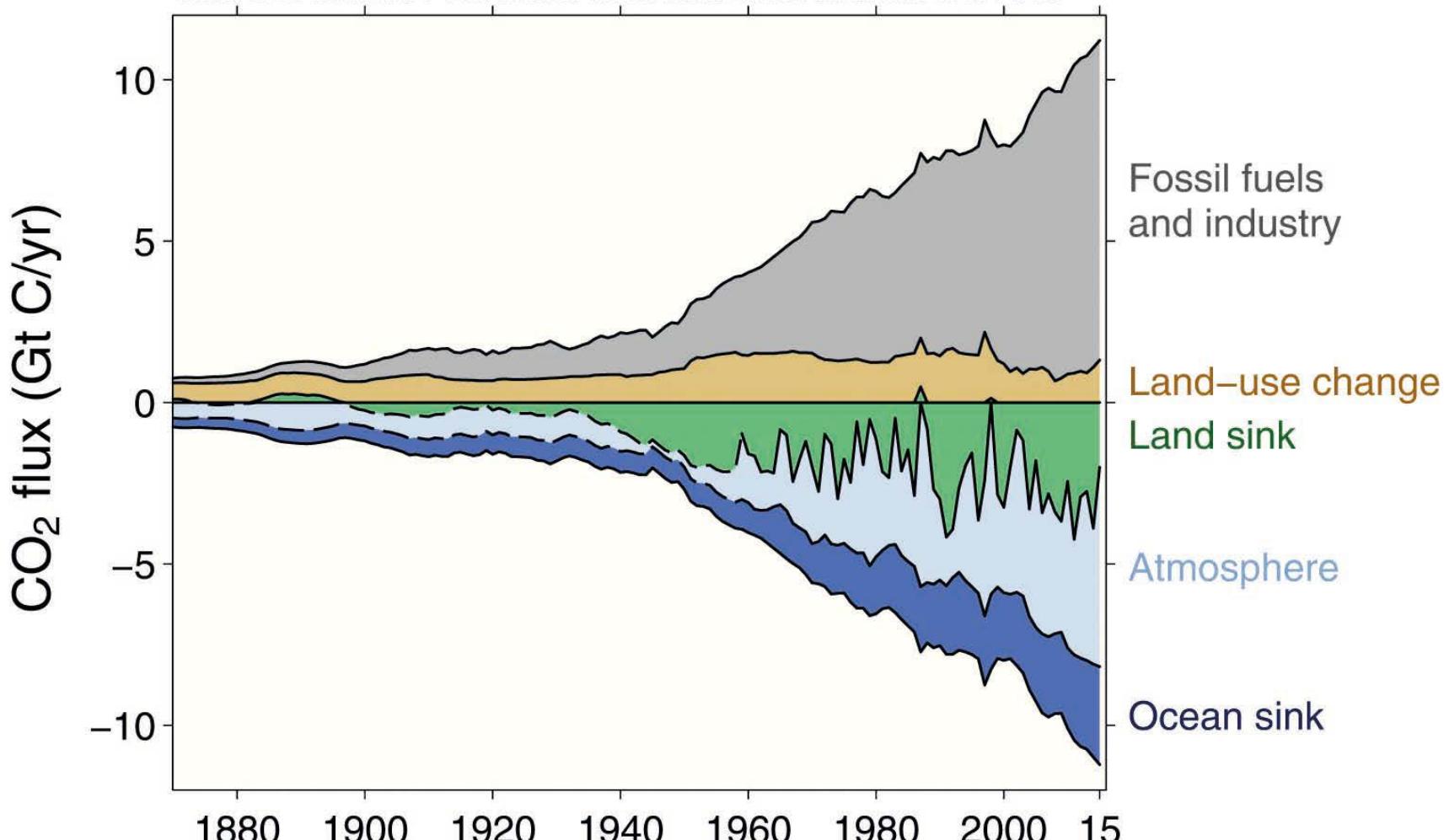


Source: Chisholm, 2000

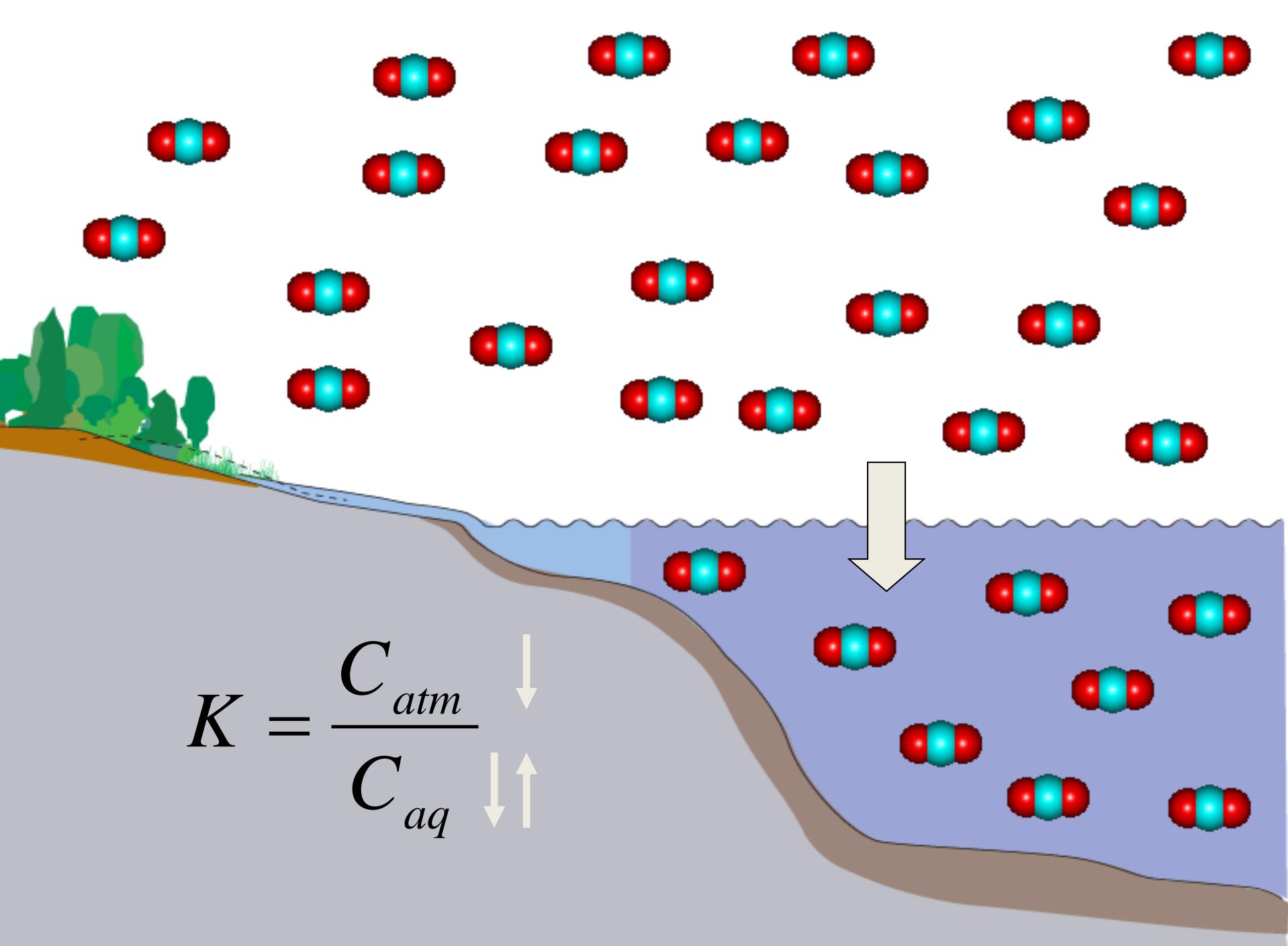


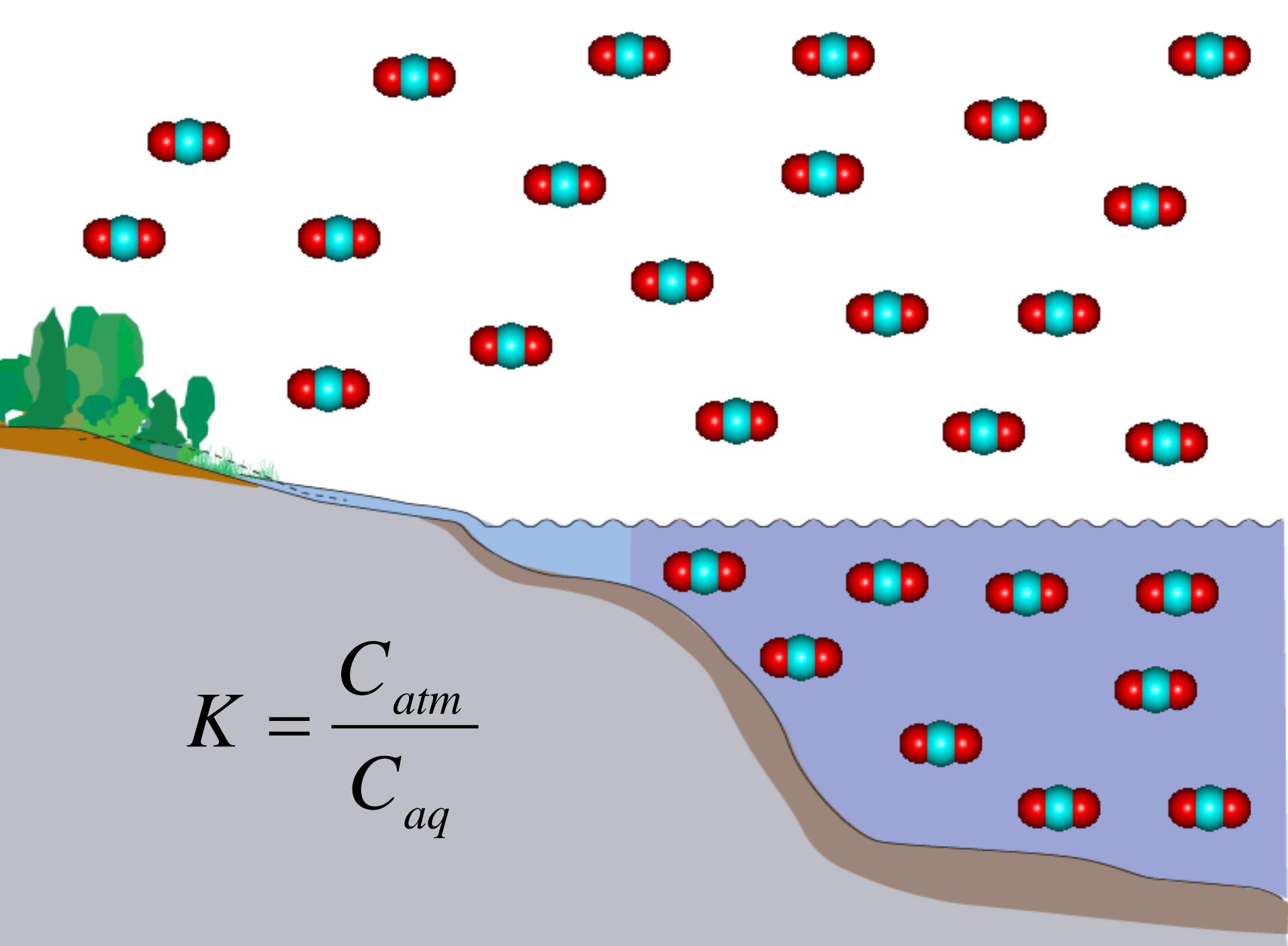
Source: Emerson & Hedges, 2008

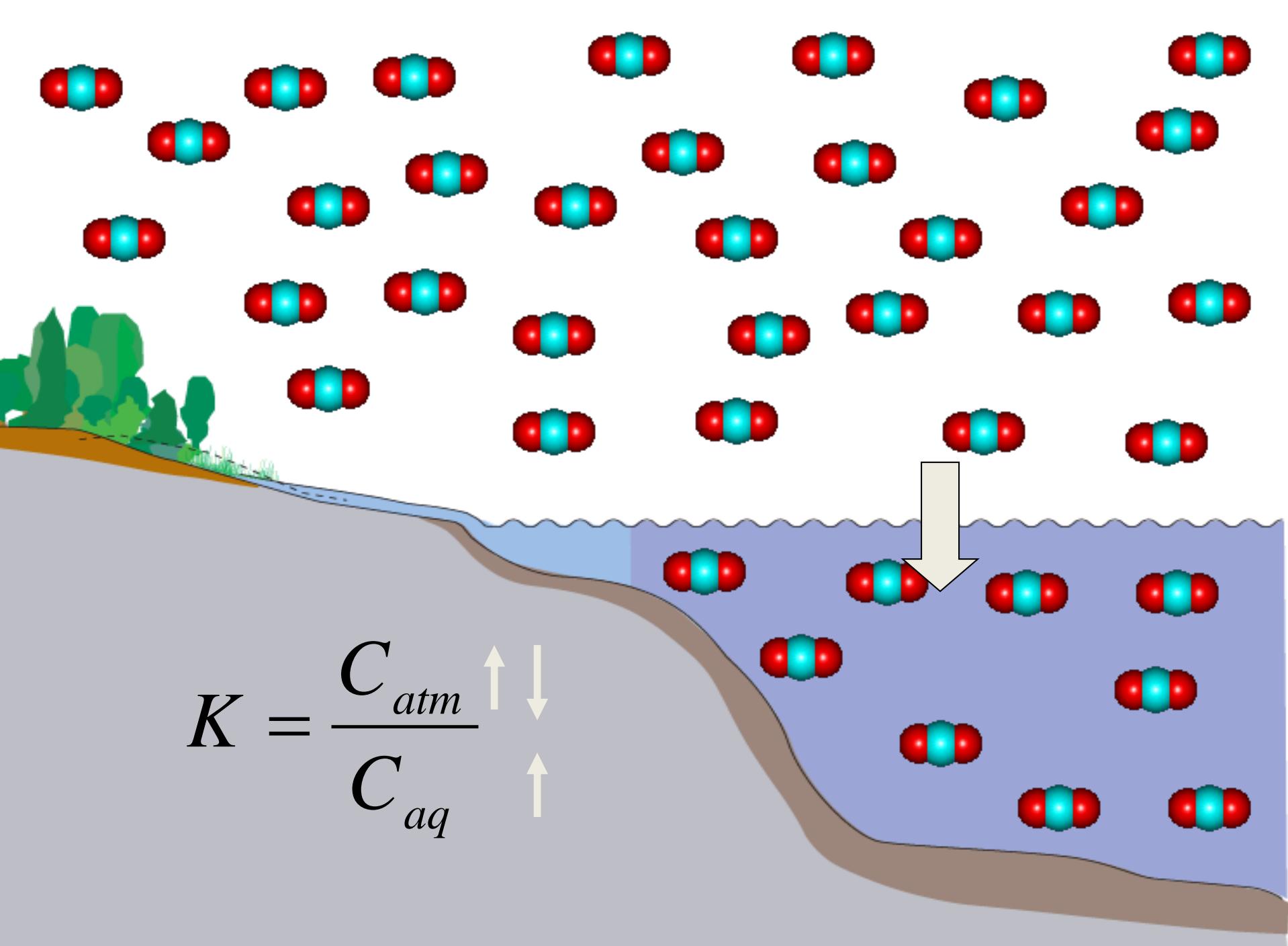
Data: CDIAC/NOAA-ESRL/GCP/Joos et al 2013/Khatiwala et al 2013



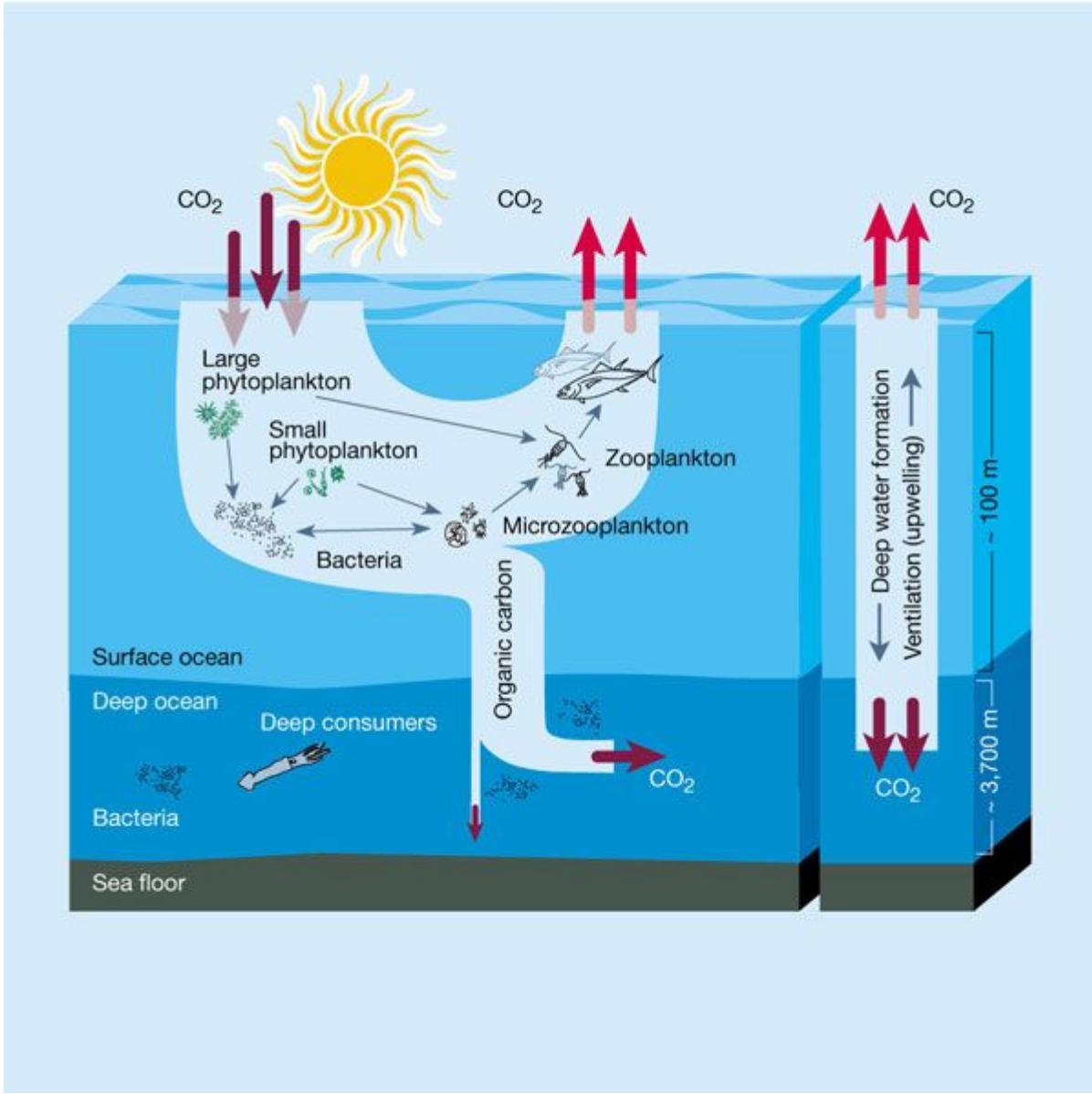
Global Carbon Project



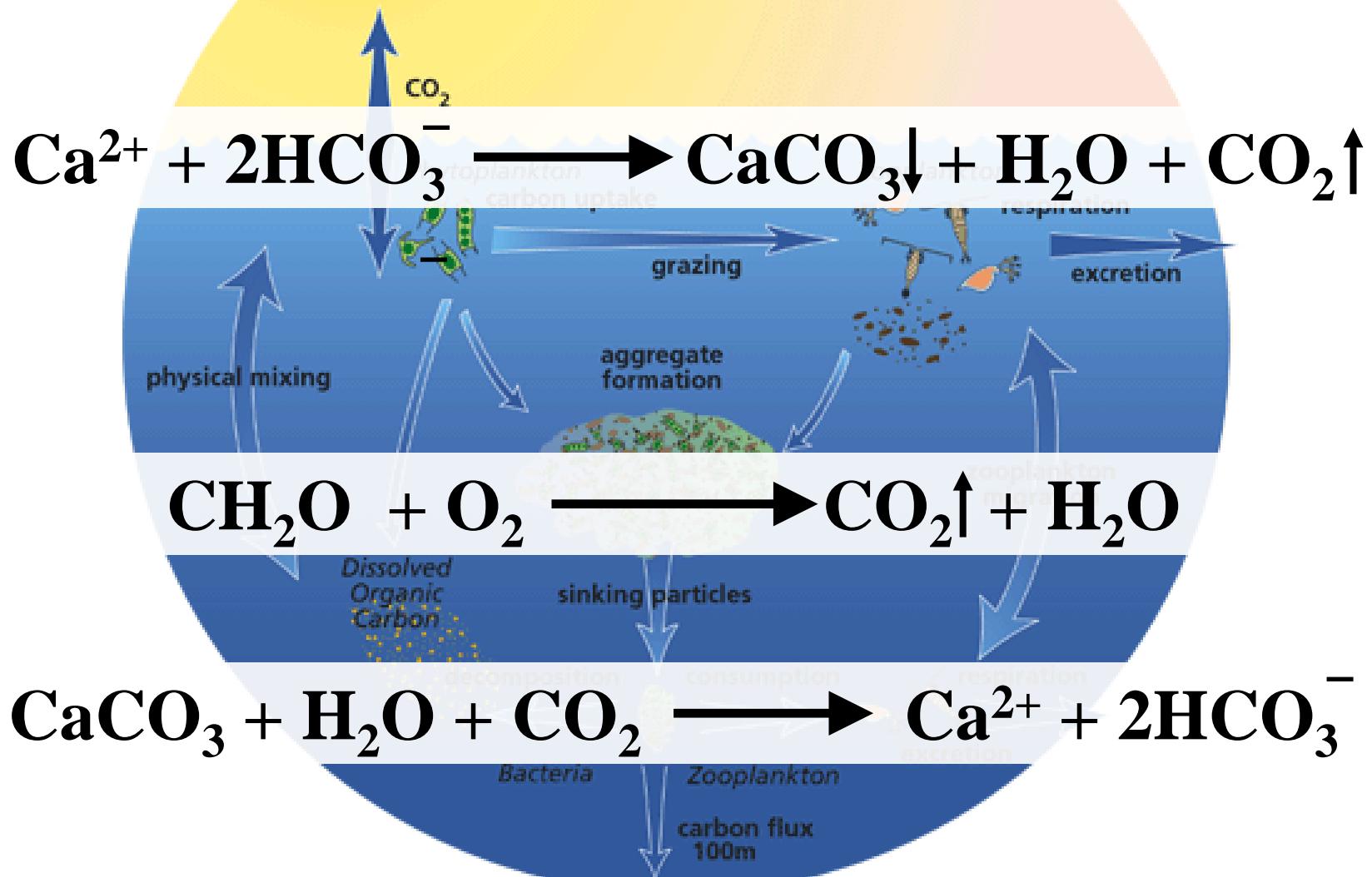


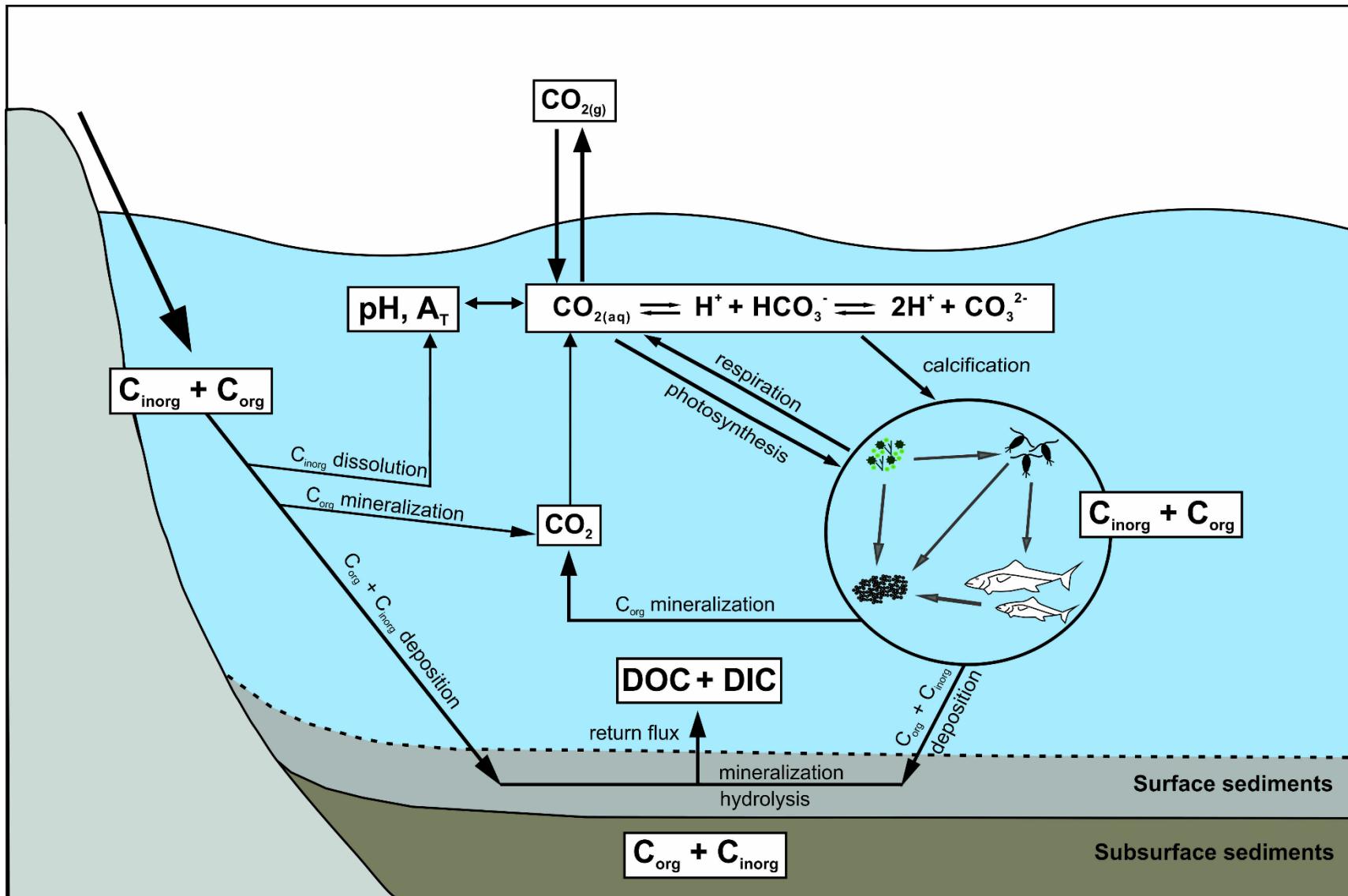


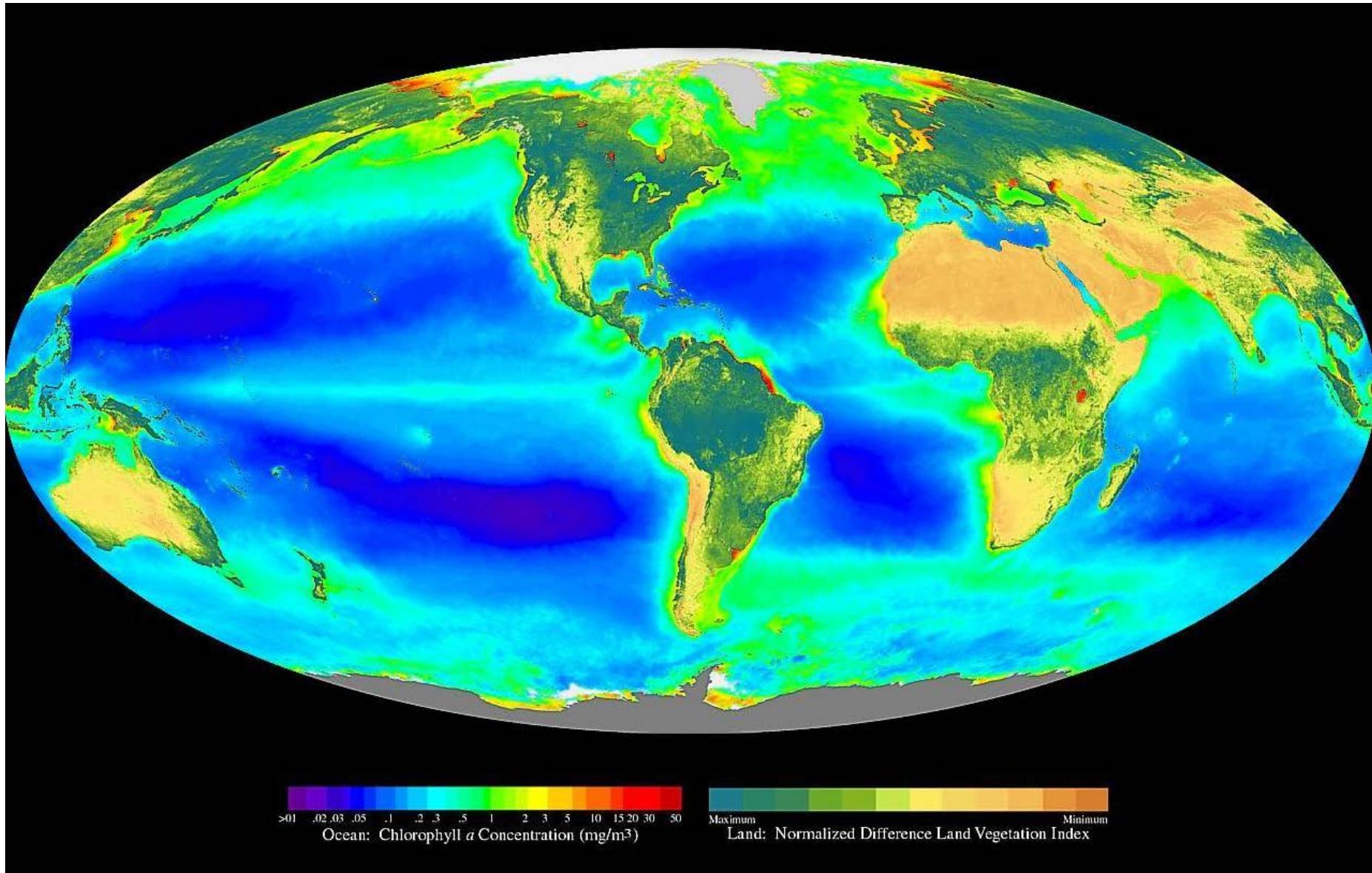
$$K = \frac{C_{atm}}{C_{aq}}$$



Source: Chisholm, 2000

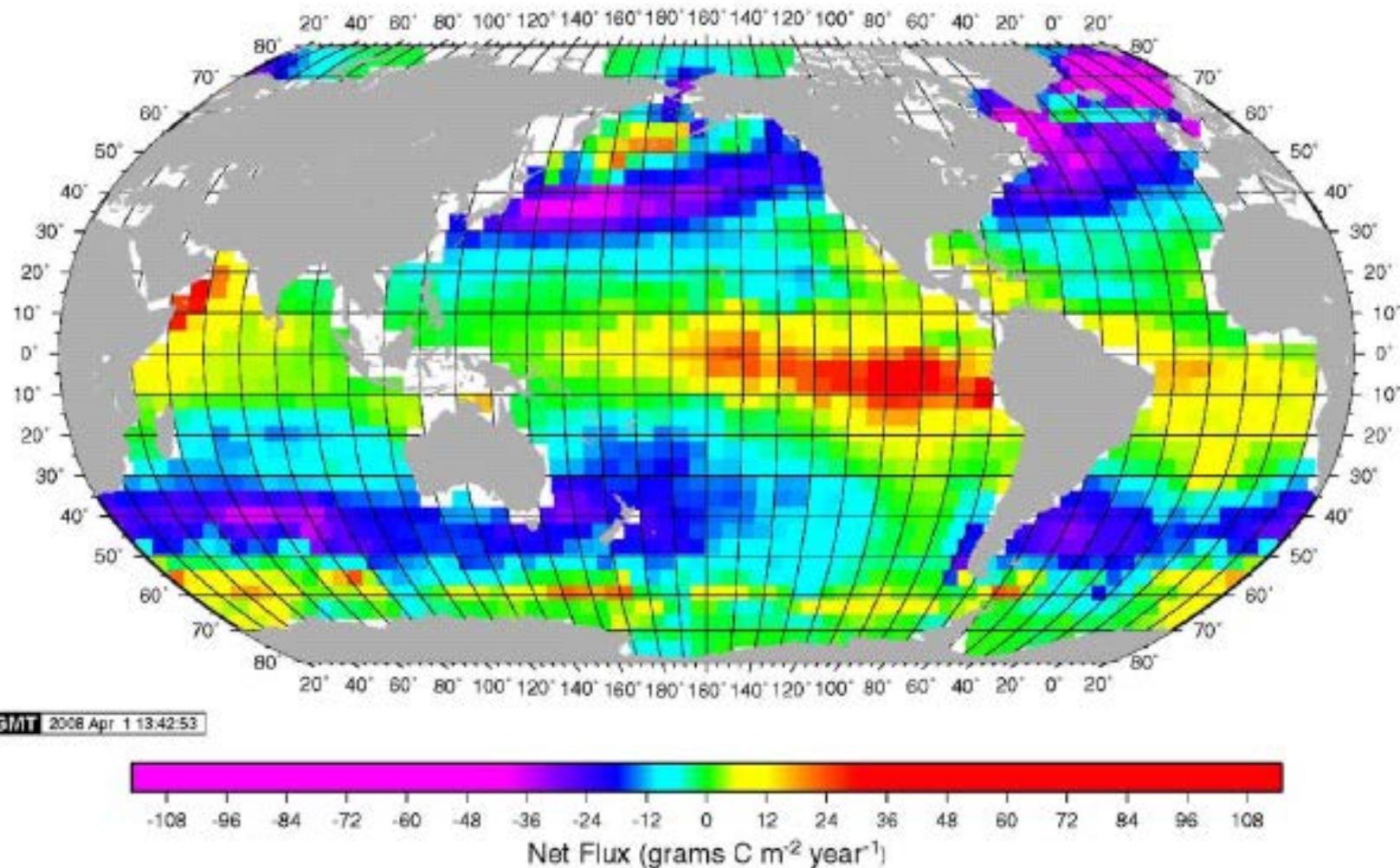




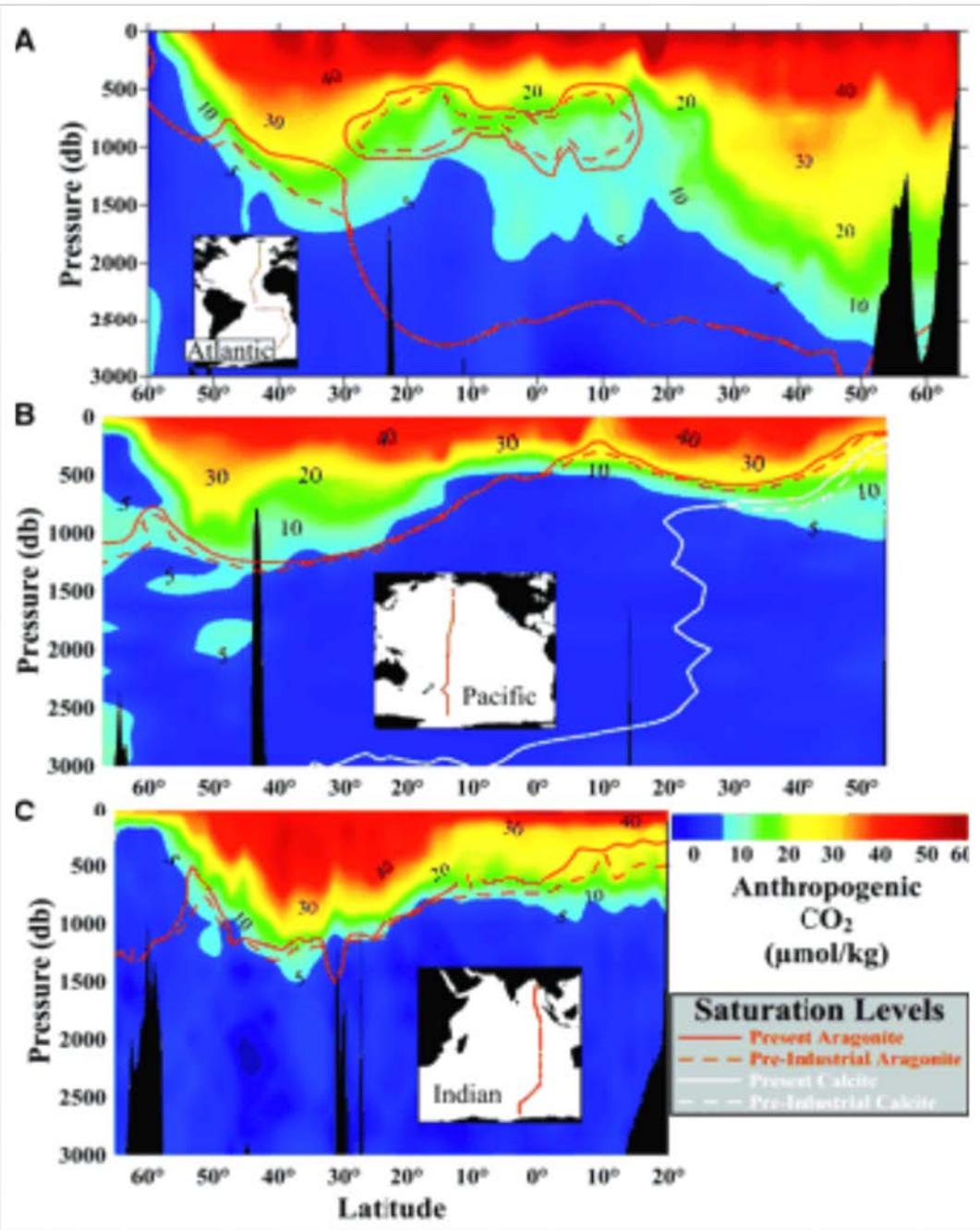


Source: NASA

Oceans absorb globally $1,6 \pm 0,9$ Pg C yr $^{-1}$

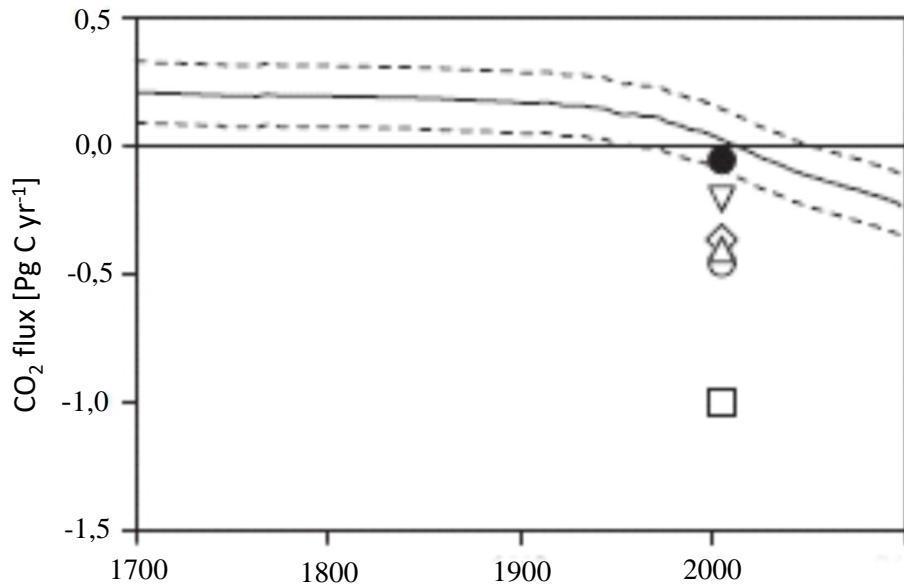


Source: Takahashi et al., 2009



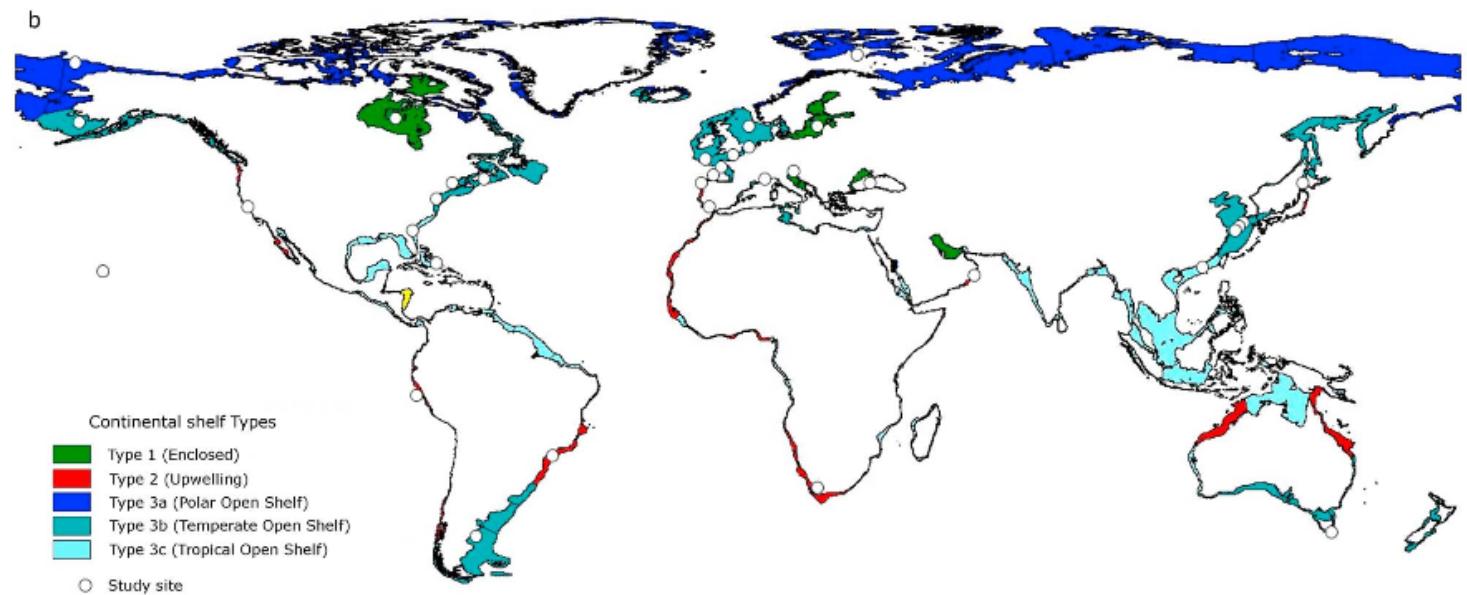
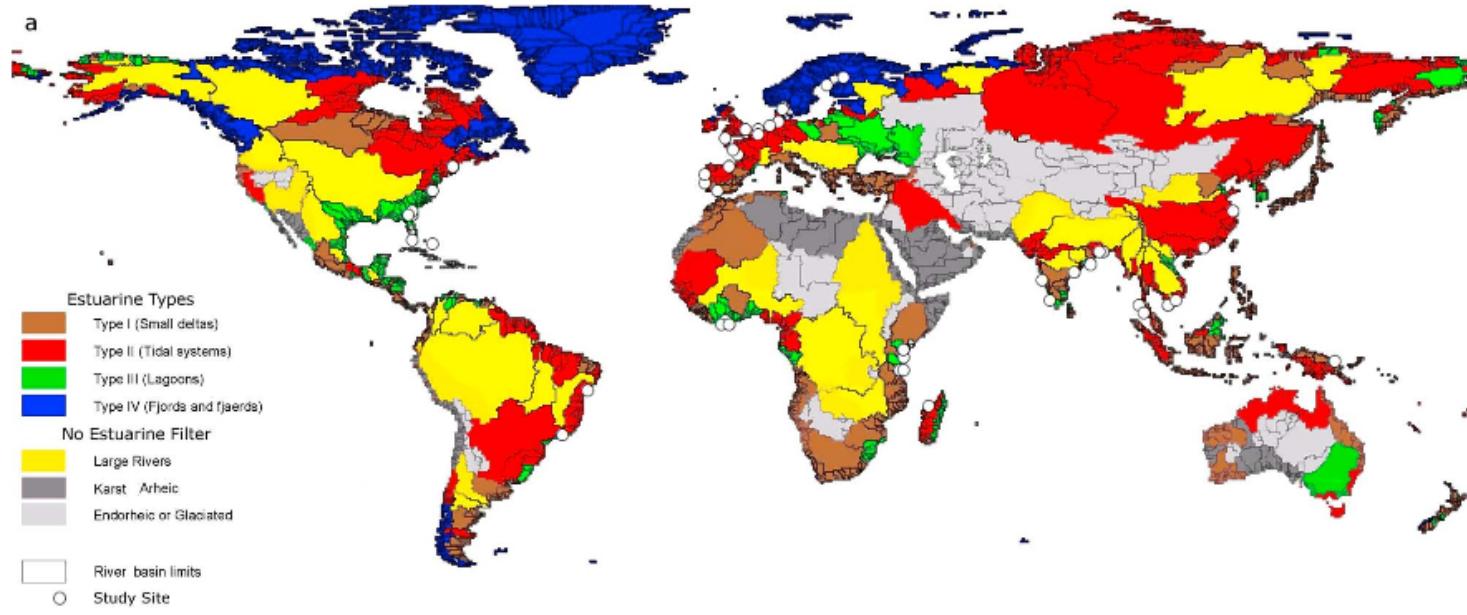
Source: Feely et al., 2004

Role of the shelf seas



- Andersson & Mackenzie, 2004
- Borges et al., 2005 (coastal + marginal seas)
- Borges et al., 2005 (coastal seas)
- ▽ Cai & Dai., 2004
- ◇ Chen, 2004
- △ Thomas et al., 2004
- Tsunogai et al., 1999





Source: Laruelle et al., 2010

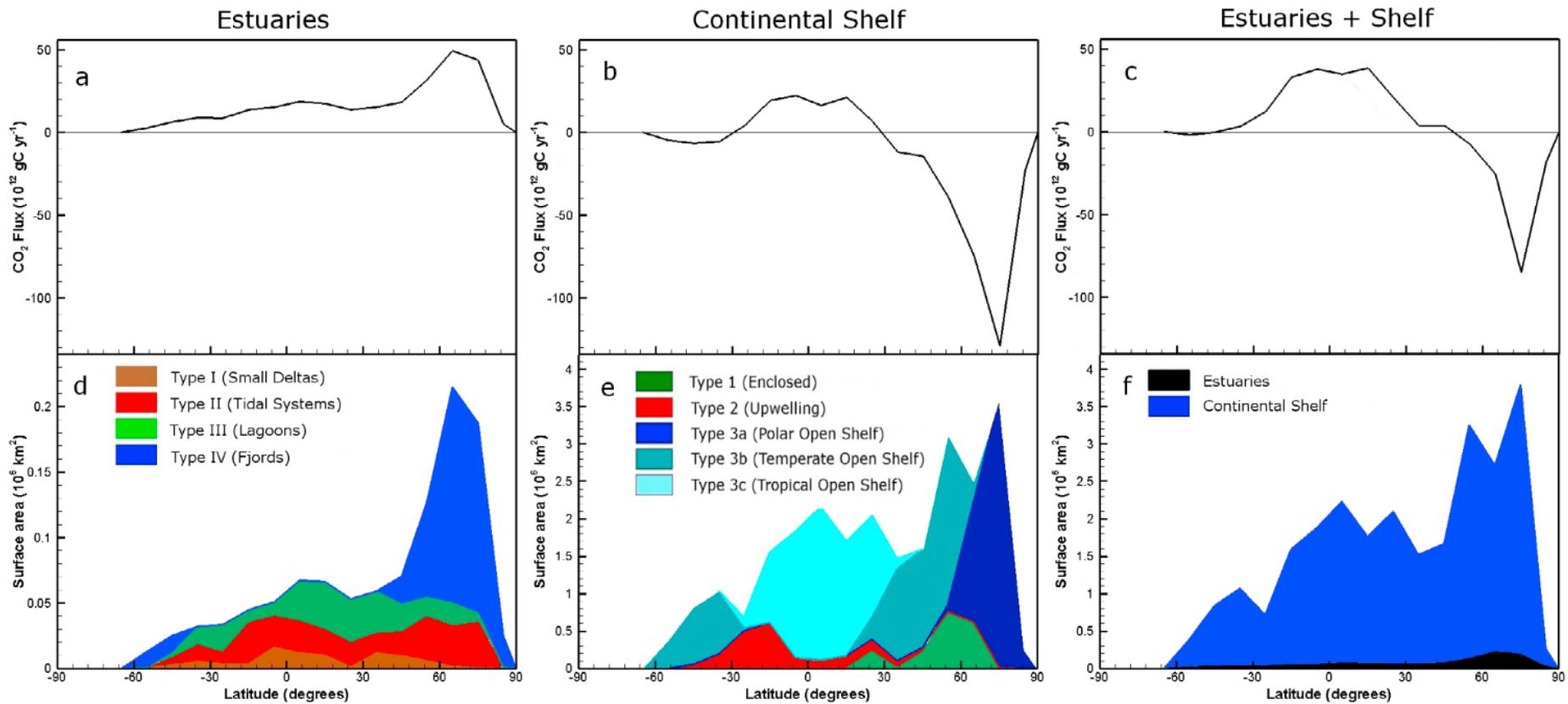


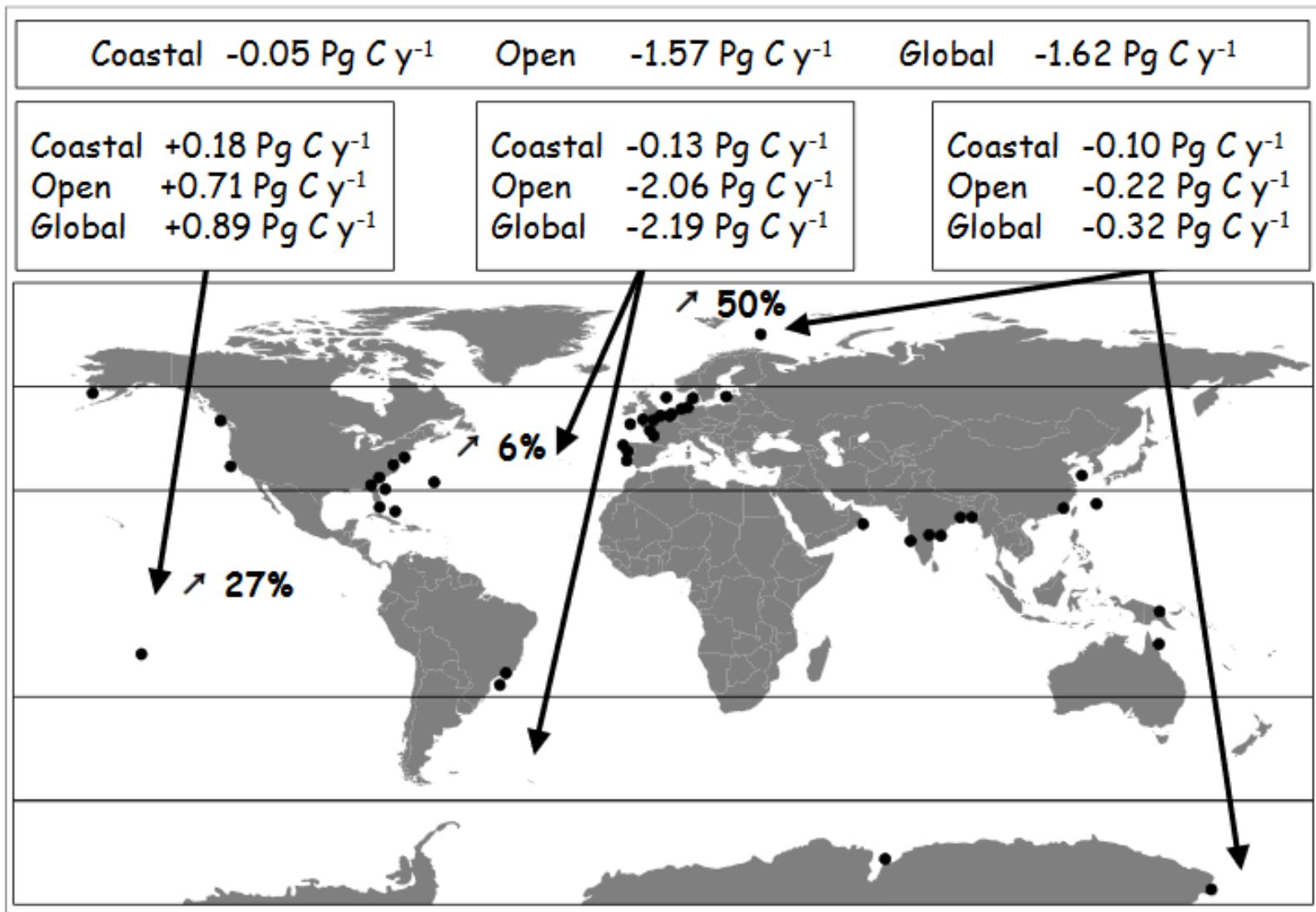
Table 1. Air-Water CO₂ Fluxes per Surface Area and Scaled Globally for Four Estuarine Types^a

	Surface Area (10 ⁶ km ²)	Air-Water CO ₂ Flux (molC m ⁻² yr ⁻¹)	Air-Water CO ₂ Flux (PgC yr ⁻¹)
Small deltas and estuaries (Type I)	0.084	25.7 ± 15.8	0.026 ± 0.016
Tidal systems and embayments (Type II)	0.276	28.5 ± 24.9	0.094 ± 0.082
Lagoons (Type III)	0.252	17.3 ± 16.6	0.052 ± 0.050
Fjords and fjärds (Type IV)	0.456	17.5 ± 14.0 ^b	0.096 ± 0.077
Total	1.067	21.0 ± 17.6	0.268 ± 0.225

Table 2. Air-Water CO₂ Fluxes per Surface Area and Scaled Globally for Different Types of Continental Shelves Along Three Climatic Zones^a

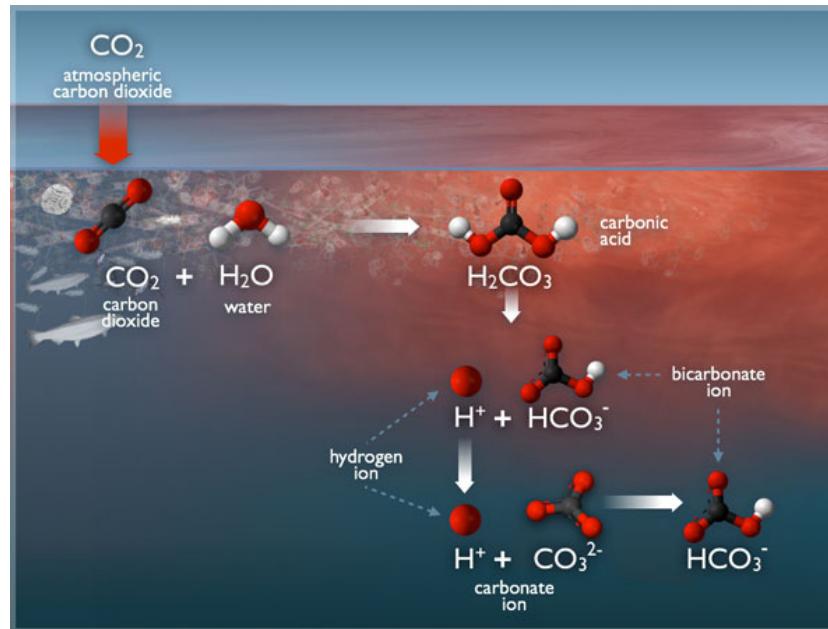
	Surface Area (10 ⁶ km ²)	Air-Water CO ₂ Flux (molC m ⁻² yr ⁻¹)	Air-Water CO ₂ Flux (PgC yr ⁻¹)
Polar (>60°)			
Enclosed	0.189	-0.8 ± 1.1	-0.002 ± 0.003
Open Shelf	5.477	-3.3 ± 1.7	-0.216 ± 0.111
Upwelling Pacific	0.086	3.2 ± 2.4	0.003 ± 0.002
Sub-total	5.752	-3.1 ± 1.7	-0.214 ± 0.116
Temperate (30°–60°)			
Enclosed	1.410	-0.8 ± 1.1	-0.014 ± 0.019
Open Shelf	7.170	-1.0 ± 1.0	-0.086 ± 0.087
Upwelling Pacific	0.293	3.2 ± 2.4	0.011 ± 0.008
Upwelling Atlantic	0.086	-1.6 ± 1.0	-0.002 ± 0.001
Upwelling Indian	0.123	0.9 ± 1.2 ^b	0.001 ± 0.002
Sub-total	9.082	-0.8 ± 1.1	-0.090 ± 0.117
Tropical (0–30°)			
Enclosed	0.231	-0.8 ± 1.1	-0.002 ± 0.003
Open Shelf	7.909	0.9 ± 1.0	0.083 ± 0.097
Upwelling Pacific	0.515	3.2 ± 2.4	0.020 ± 0.015
Upwelling Atlantic	0.715	-1.6 ± 1.0	-0.014 ± 0.009
Upwelling Indian	0.520	0.9 ± 1.2 ^b	0.006 ± 0.008
Sub-total	9.890	0.8 ± 1.1	0.093 ± 0.131
Total	24.724	-0.7 ± 1.2	-0.211 ± 0.364

Source: Laruelle et al., 2010



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Seawater acid-base system

Measurable parameters:

- C_T – total CO_2 concentration (DIC)

$$C_T = [\text{CO}_2]^* + [\text{HCO}_3^-] + [\text{CO}_3^{2-}]$$

- A_T – total alkalinity

$$A_T = [\text{HCO}_3^-] + 2[\text{CO}_3^{2-}] + [\text{B(OH)}_4^-] + [\text{OH}^-] + \dots - [\text{H}^+] - \dots$$

- $p\text{CO}_2$ – CO_2 partial pressure

- pH – spectrophotometric measurement with m-cresol purple, total scale

$$\text{pH}_T = -\log ([\text{H}^+]_F + [\text{HSO}_4^-]) = -\log [\text{H}^+]_T$$

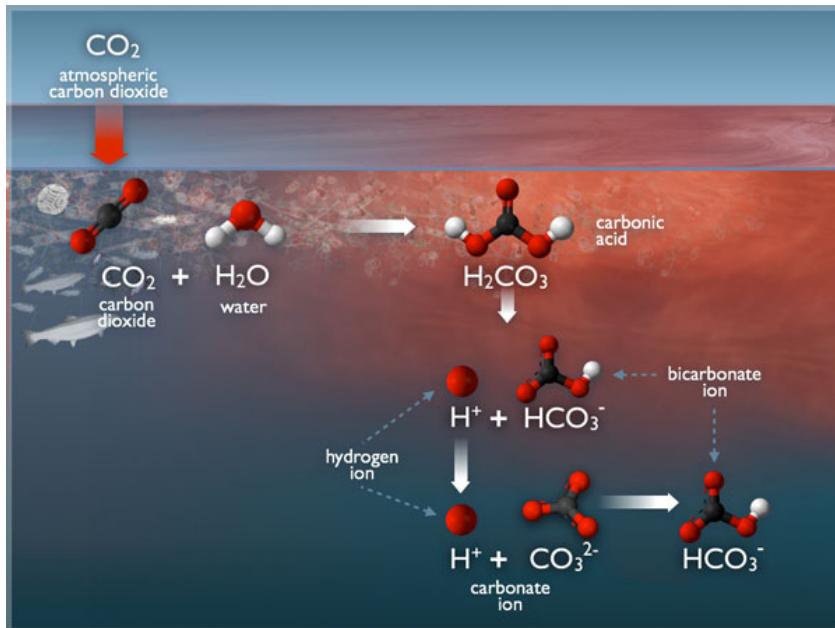
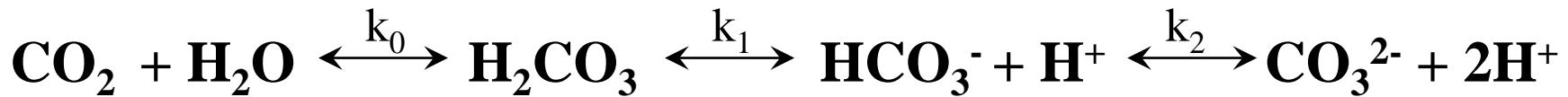
It is possible to calculate 2 parameters when the following is known:

- other 2 parameters
- temperature & salinity
- equilibrium constants for each of the acid dissociation reactions
- total concentrations for each non- CO_2 substances

The pair used in the calculations:

- C_T & A_T – recommended, used in biogeochemical modelling

CO_2 system



$$k_0 = \frac{[\text{H}_2\text{CO}_3]}{\text{pCO}_2}$$

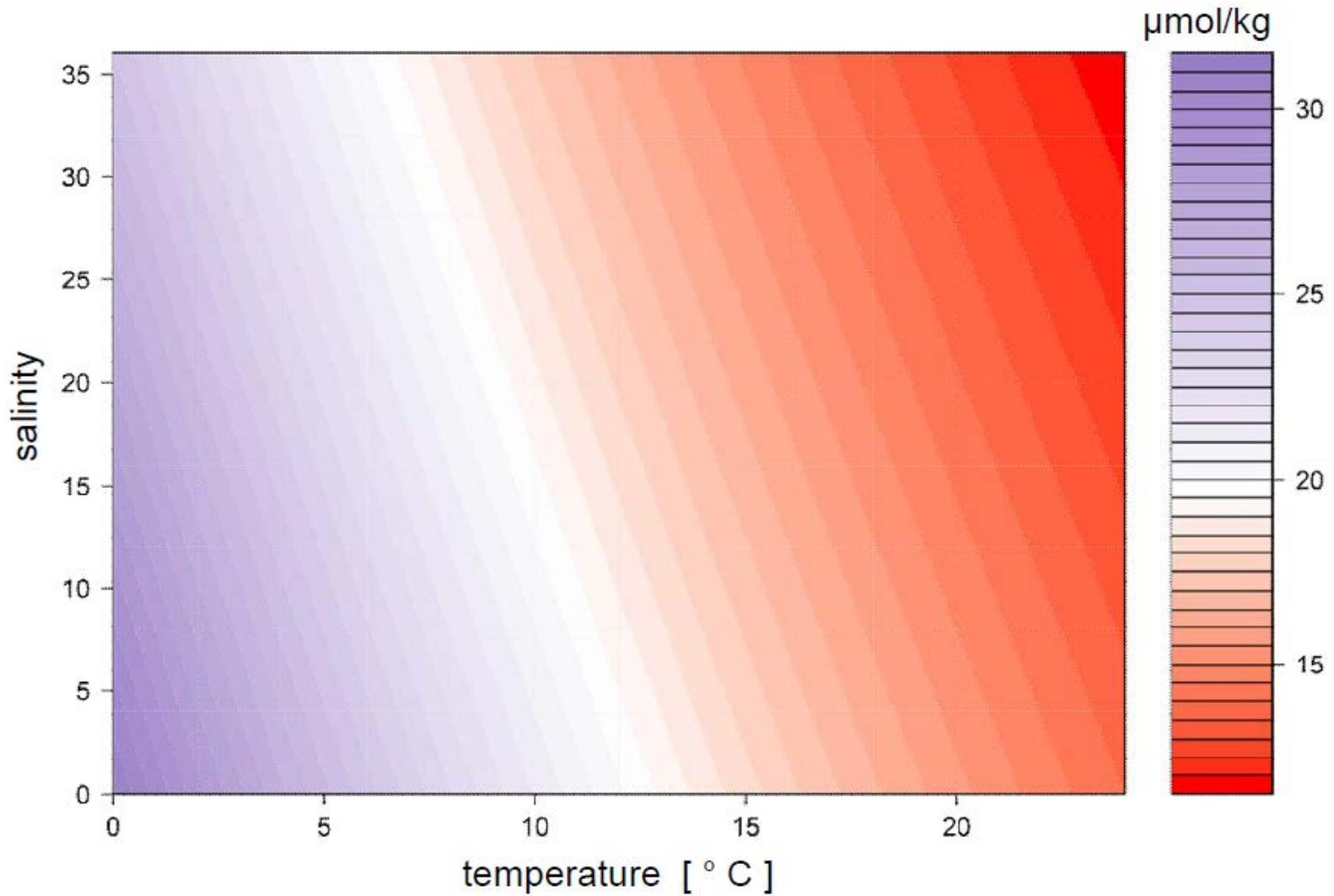
$$k_1 = \frac{[\text{H}^+][\text{HCO}_3^-]}{[\text{H}_2\text{CO}_3]}$$

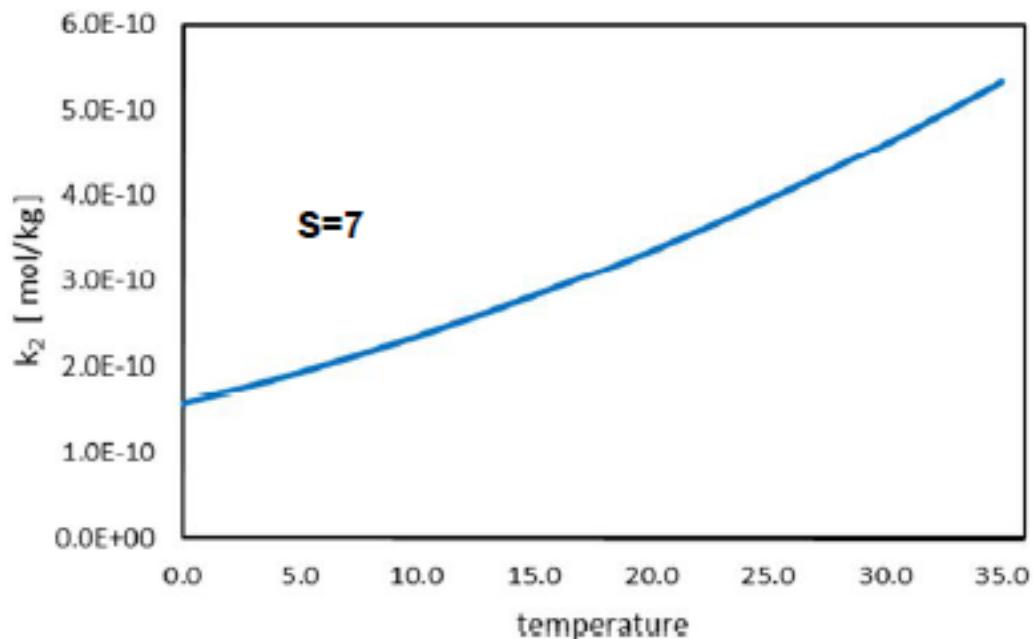
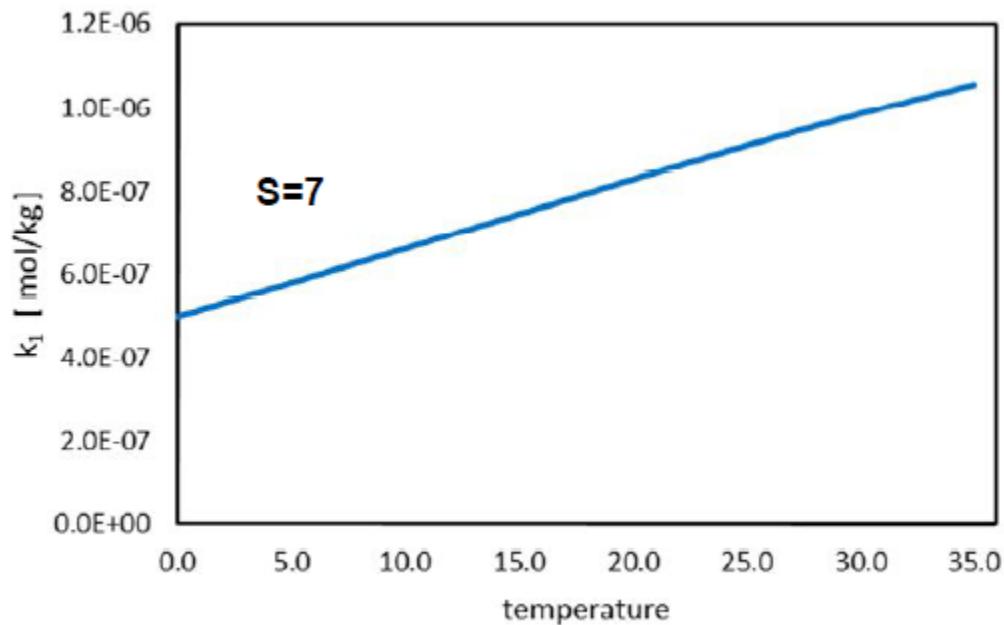
$$k_2 = \frac{[\text{H}^+][\text{CO}_3^{2-}]}{[\text{HCO}_3^-]}$$

Source: www.whoi.edu



Concentration of CO_2^* as a function of S and T





Source: B. Schneider

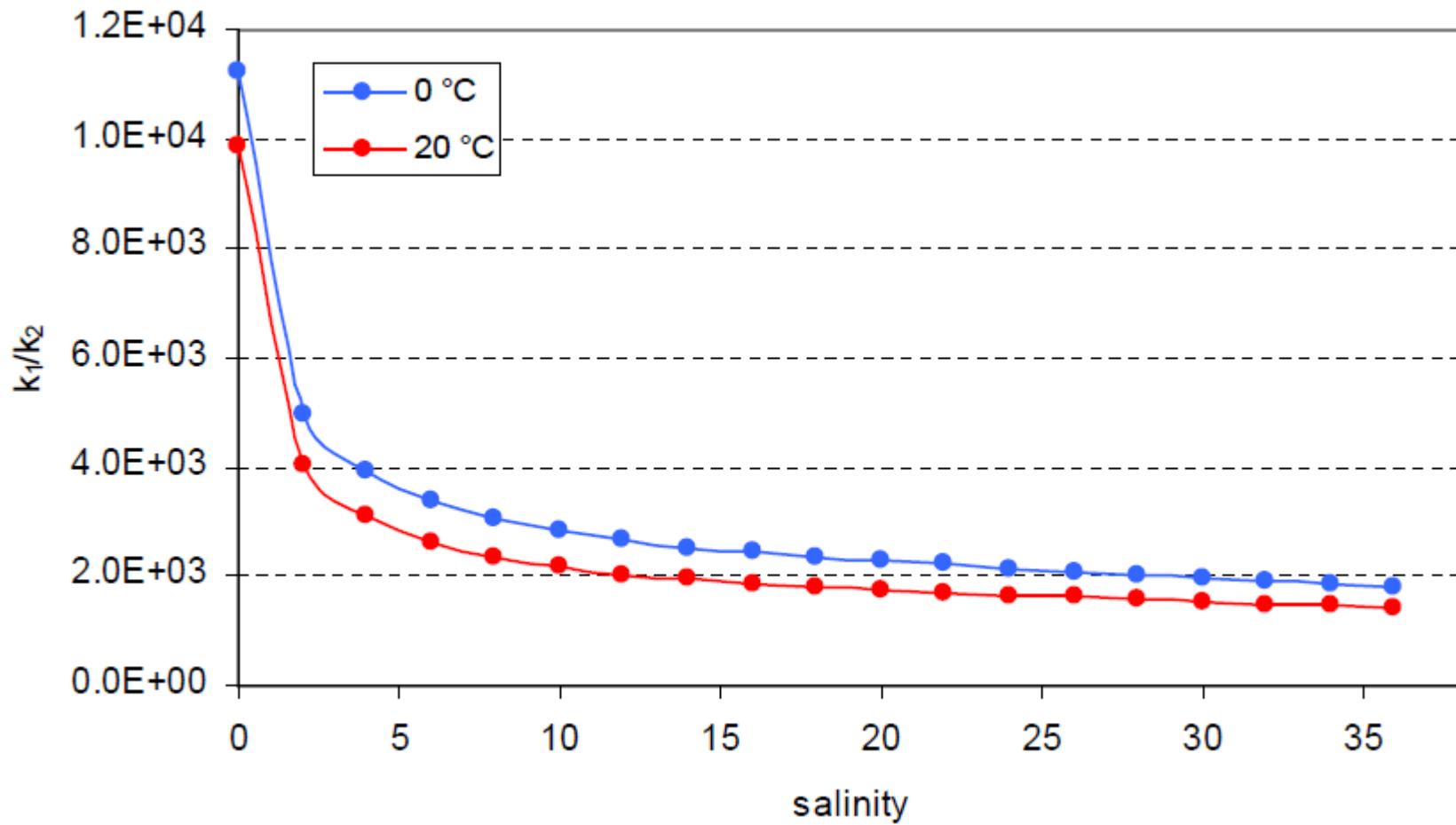
Buffer reaction



$$k_1 = \frac{[\text{H}^+][\text{HCO}_3^-]}{[\text{CO}_2^*]}$$

$$k_2 = \frac{[\text{H}^+][\text{CO}_3^{2-}]}{[\text{HCO}_3^-]}$$

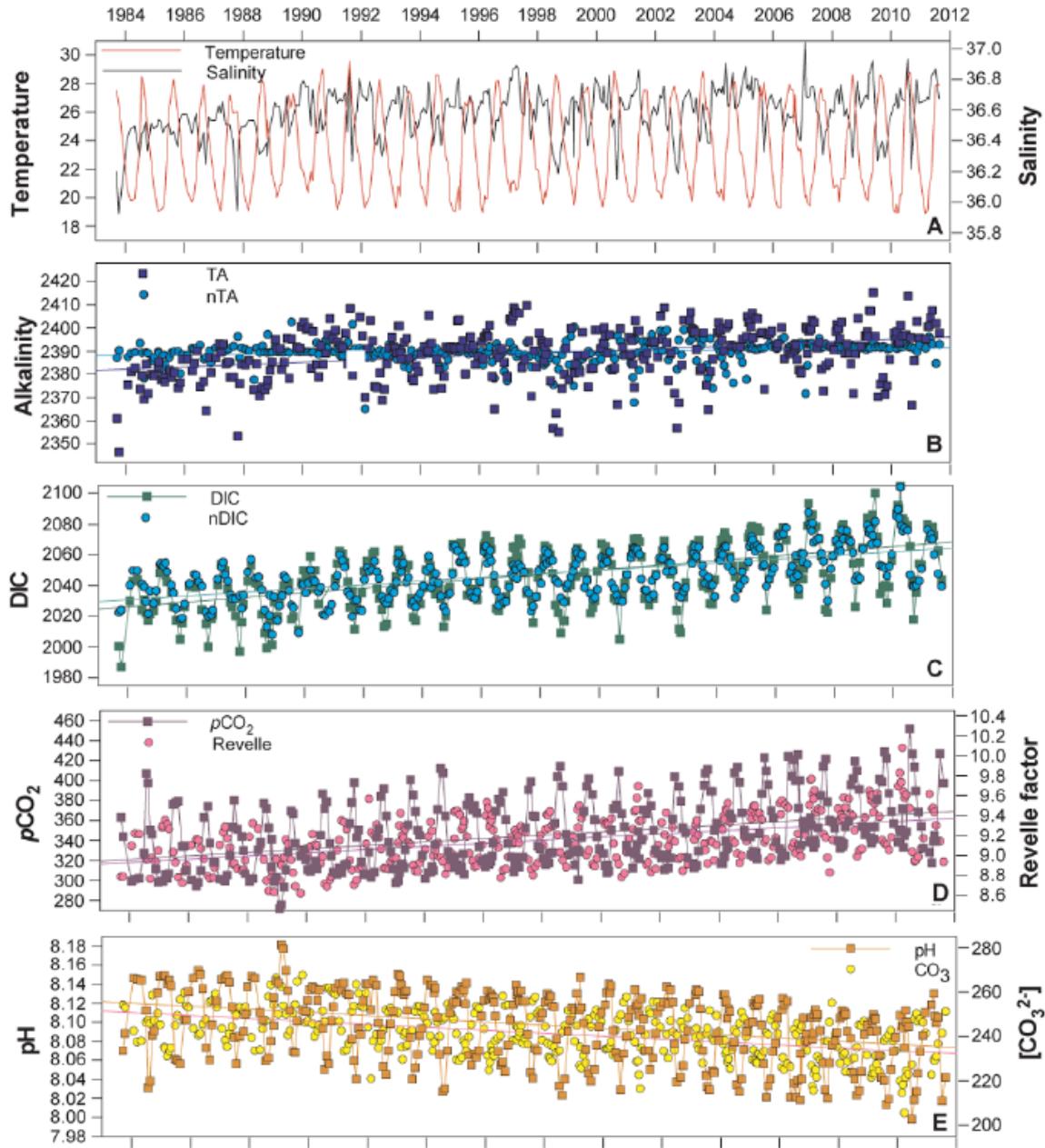
$$\frac{k_1}{k_2} = \frac{[\text{HCO}_3^-]^2}{[\text{CO}_2^*][\text{CO}_3^{2-}]}$$



$$\frac{k_1}{k_2} = \frac{[\text{HCO}_3^-]^2}{[\text{CO}_2^*][\text{CO}_3^{2-}]}$$

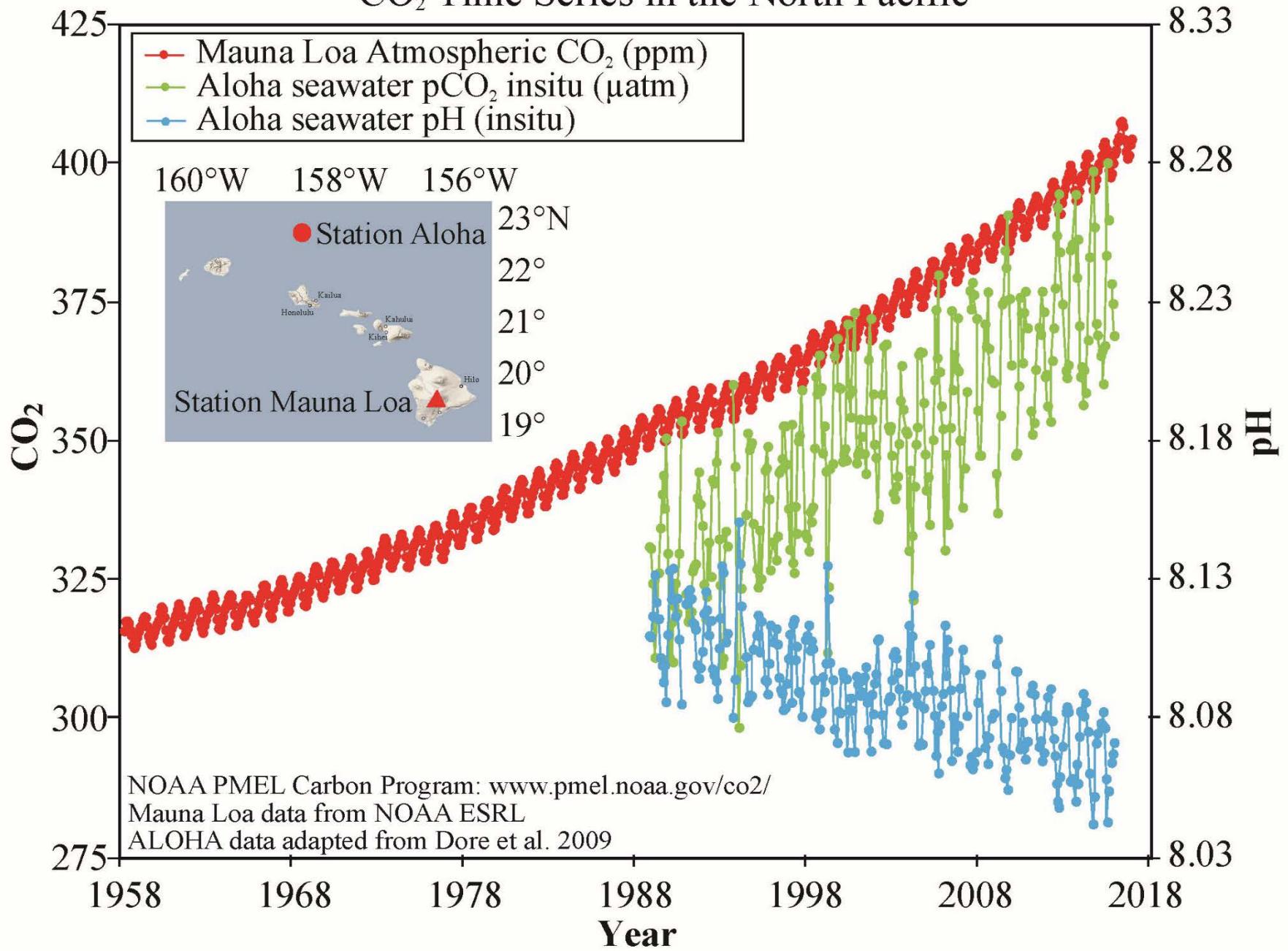
Source: B. Schneider

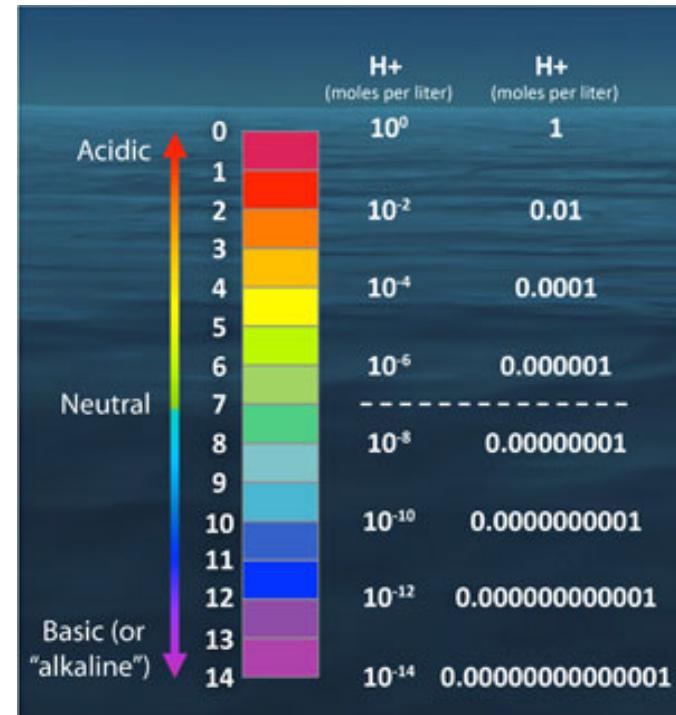
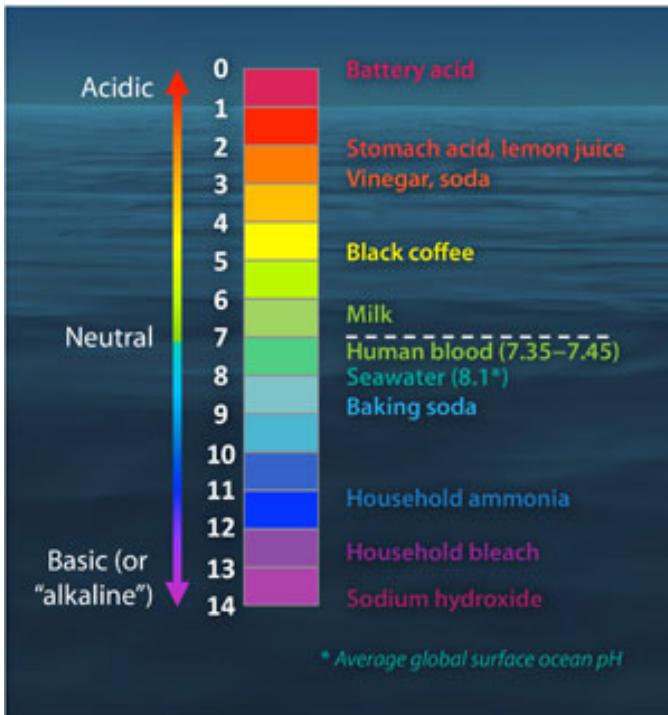
BATS - Bermuda Atlantic Time-series Study



Source: Bates et al., 2012

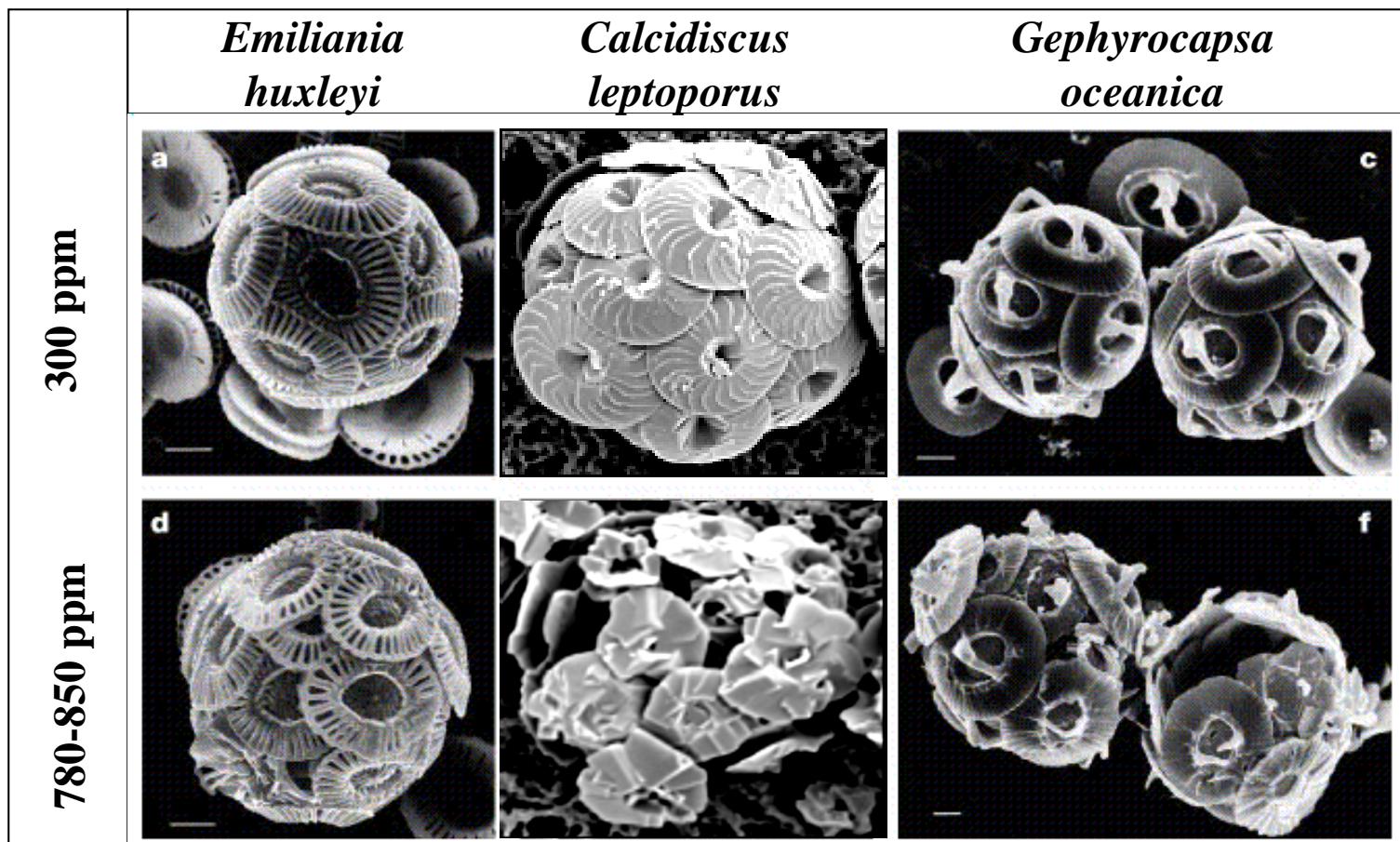
CO₂ Time Series in the North Pacific

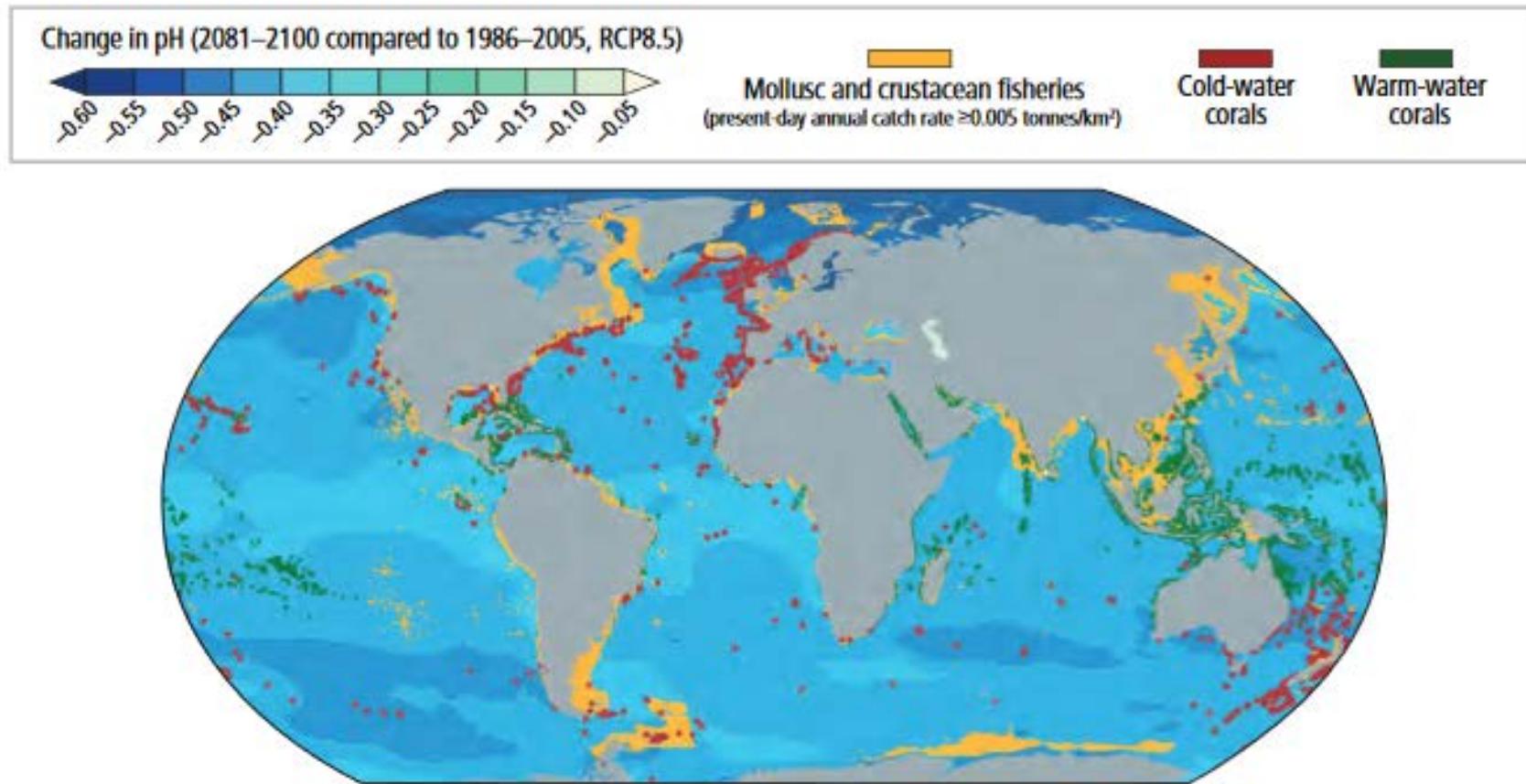




pH	H ⁺ (moles per liter)	change in acidity
7.2	6.3×10^{-8}	+900%
7.3	5.0×10^{-8}	+694%
7.4	4.0×10^{-8}	+531%
7.5	3.2×10^{-8}	+401%
7.6	2.5×10^{-8}	+298%
7.7	2.0×10^{-8}	+216%
7.8	1.6×10^{-8}	+151%
7.9	1.3×10^{-8}	+100%
8.0	1.0×10^{-8}	+58%
8.1	7.9×10^{-9}	+26%
8.2	6.3×10^{-9}	

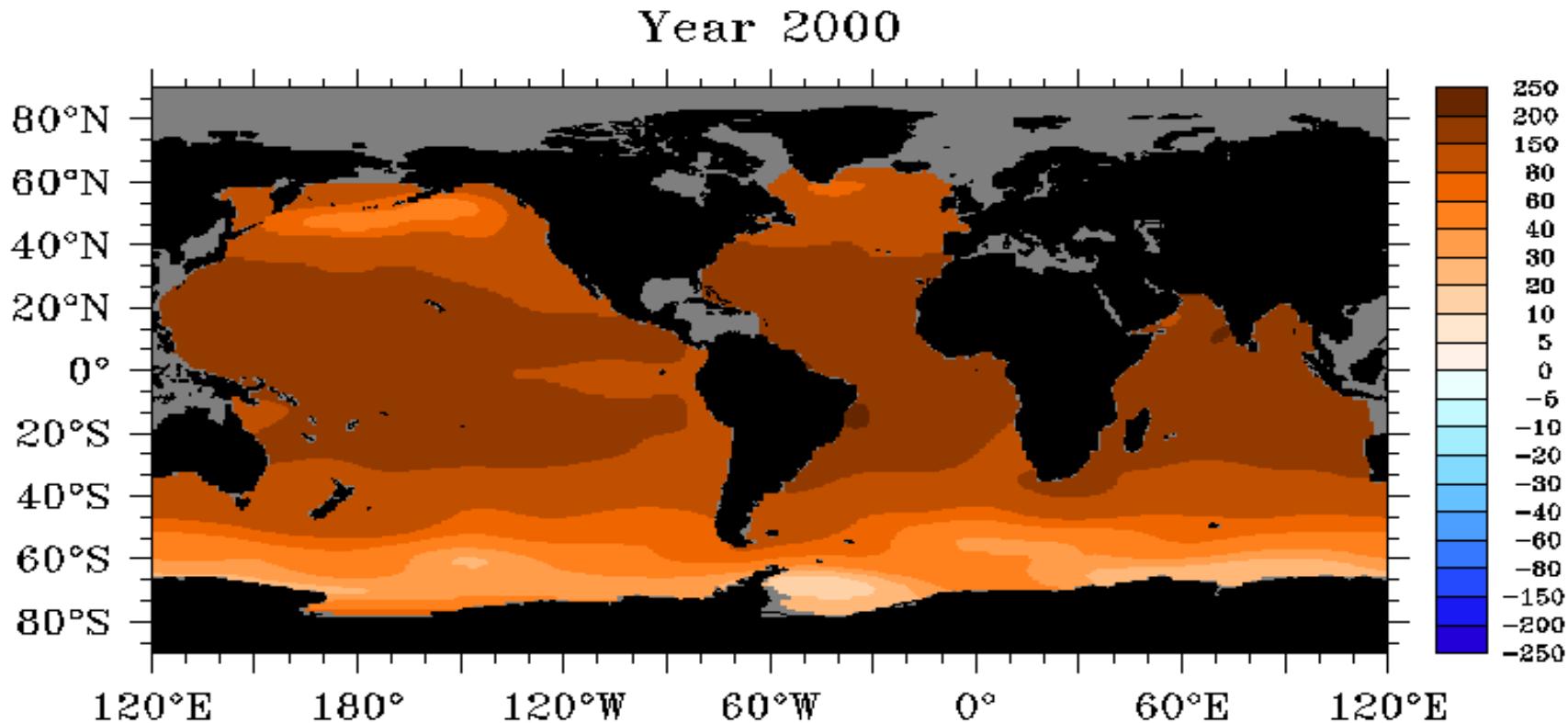
Source: NOAA





Source: IPCC, 2014

Global aragonite saturation 2099



Calciferous organisms



CaCO_3

production period

coccolithophores

calcite

days

foraminifera

calcite

weeks

pteropods

aragonite

months

corals

aragonite

months/years

doubling of CO_2 in
the atmosphere

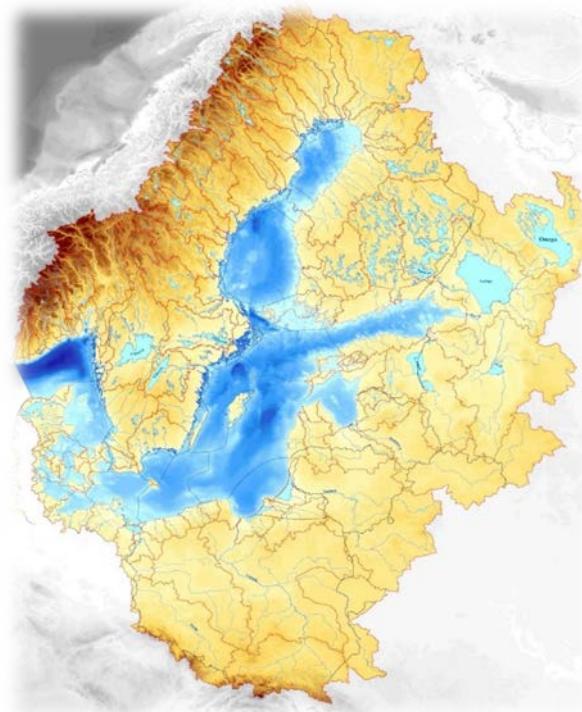


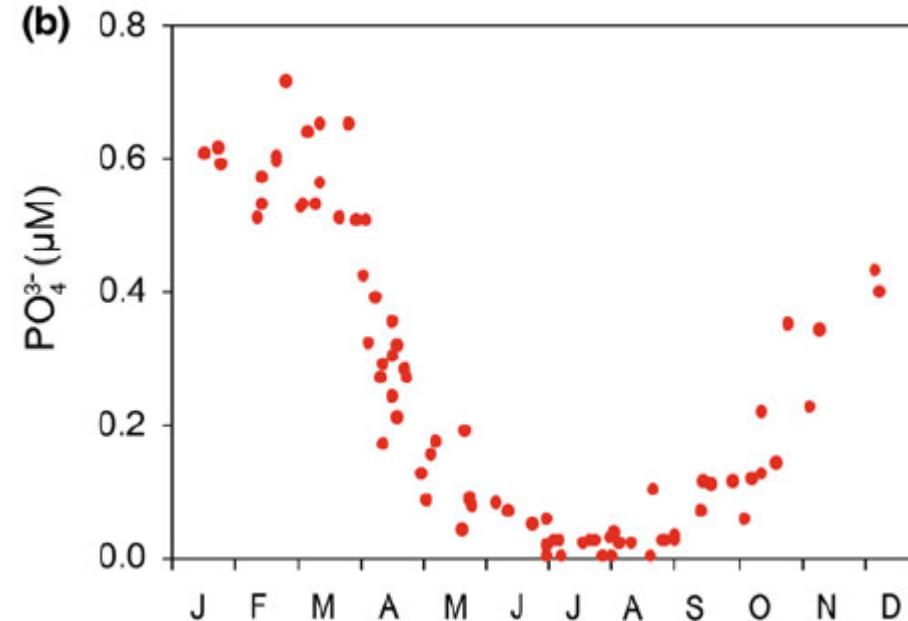
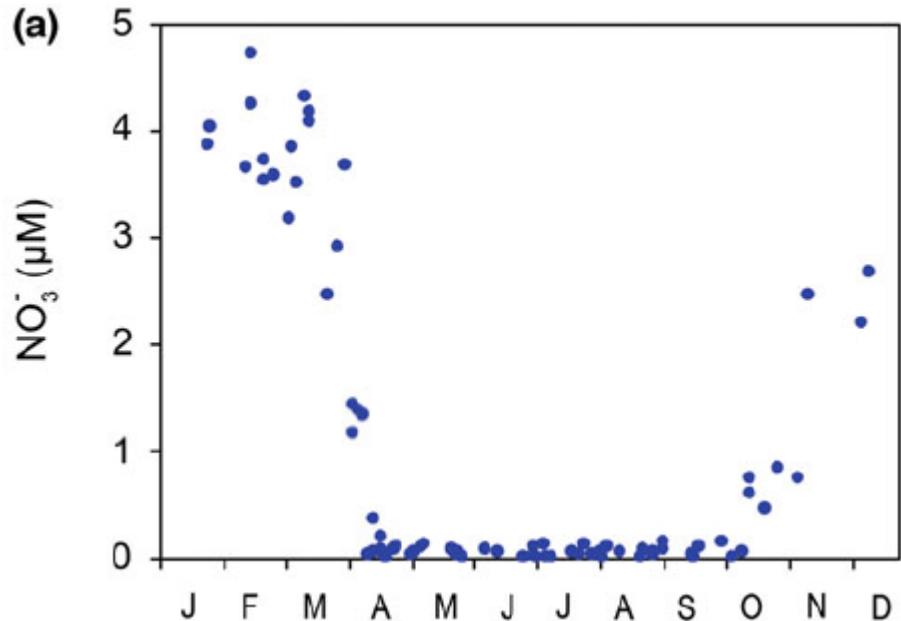
decrease in calcification
by 20-40 %



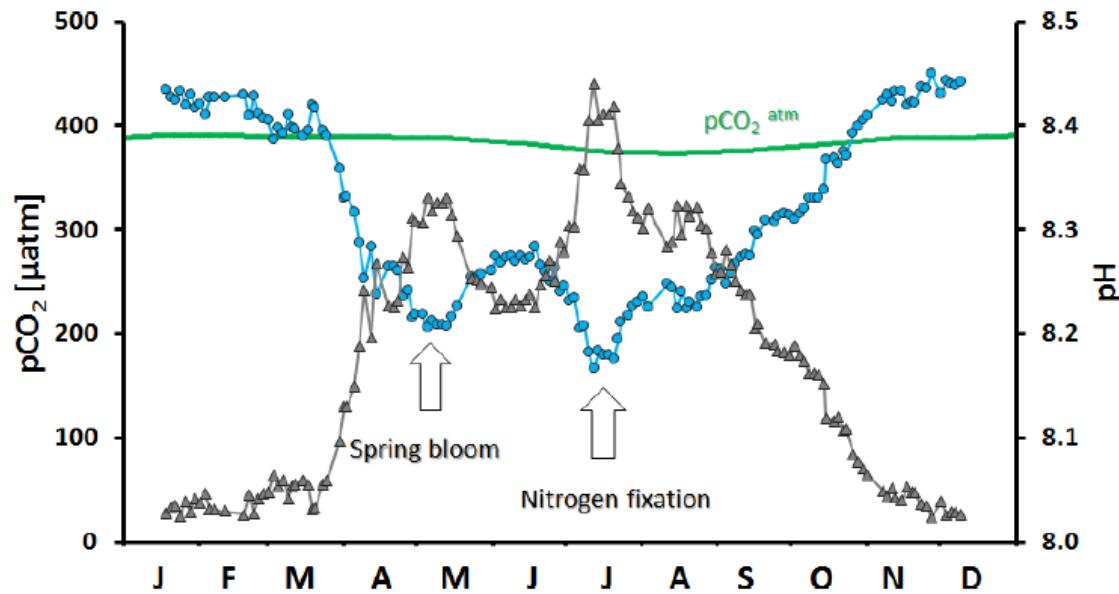
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Source: Biological Oceanography of the Baltic Sea



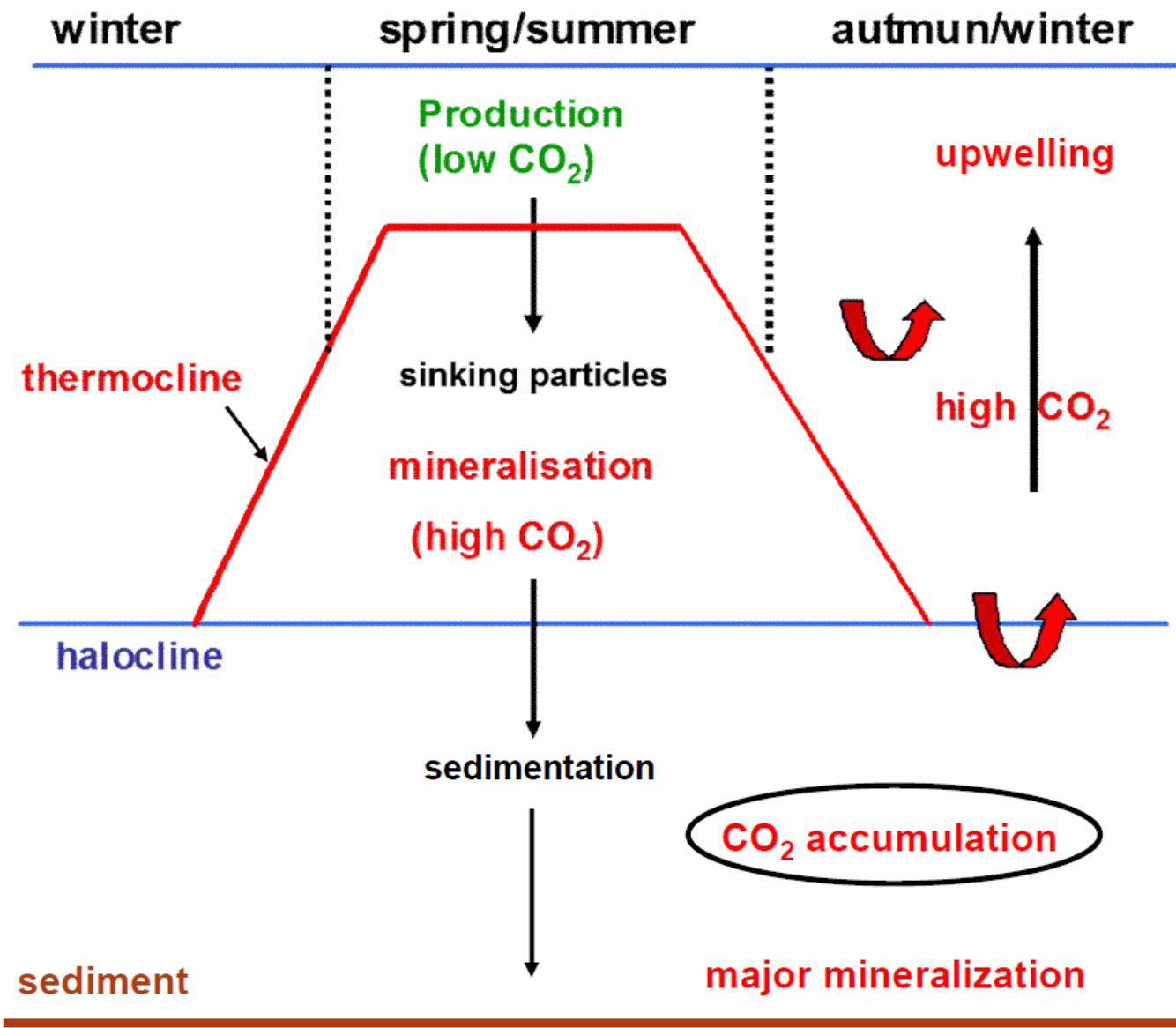
Source: Kuliński et al., 2017 modified after Schneider , 2011

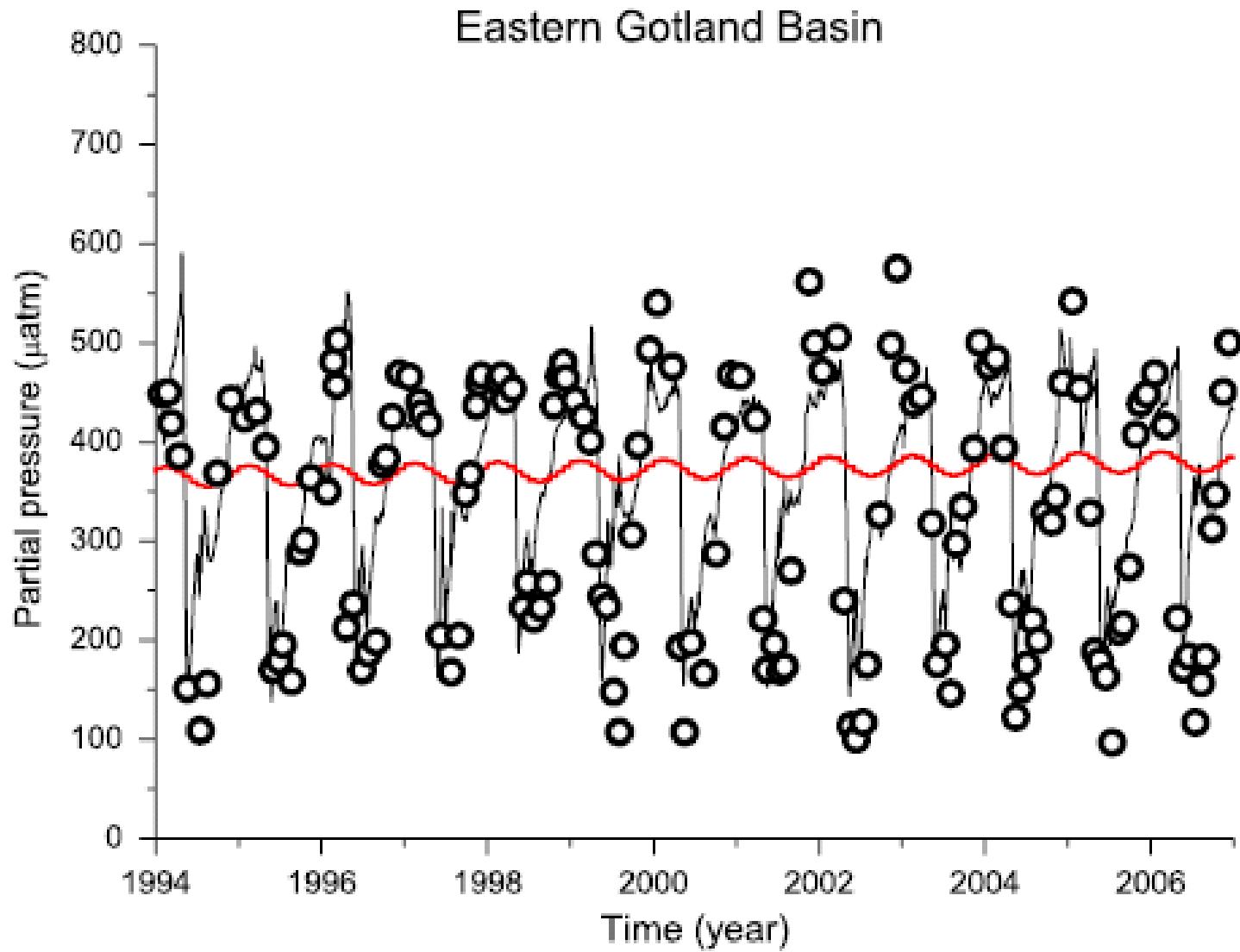


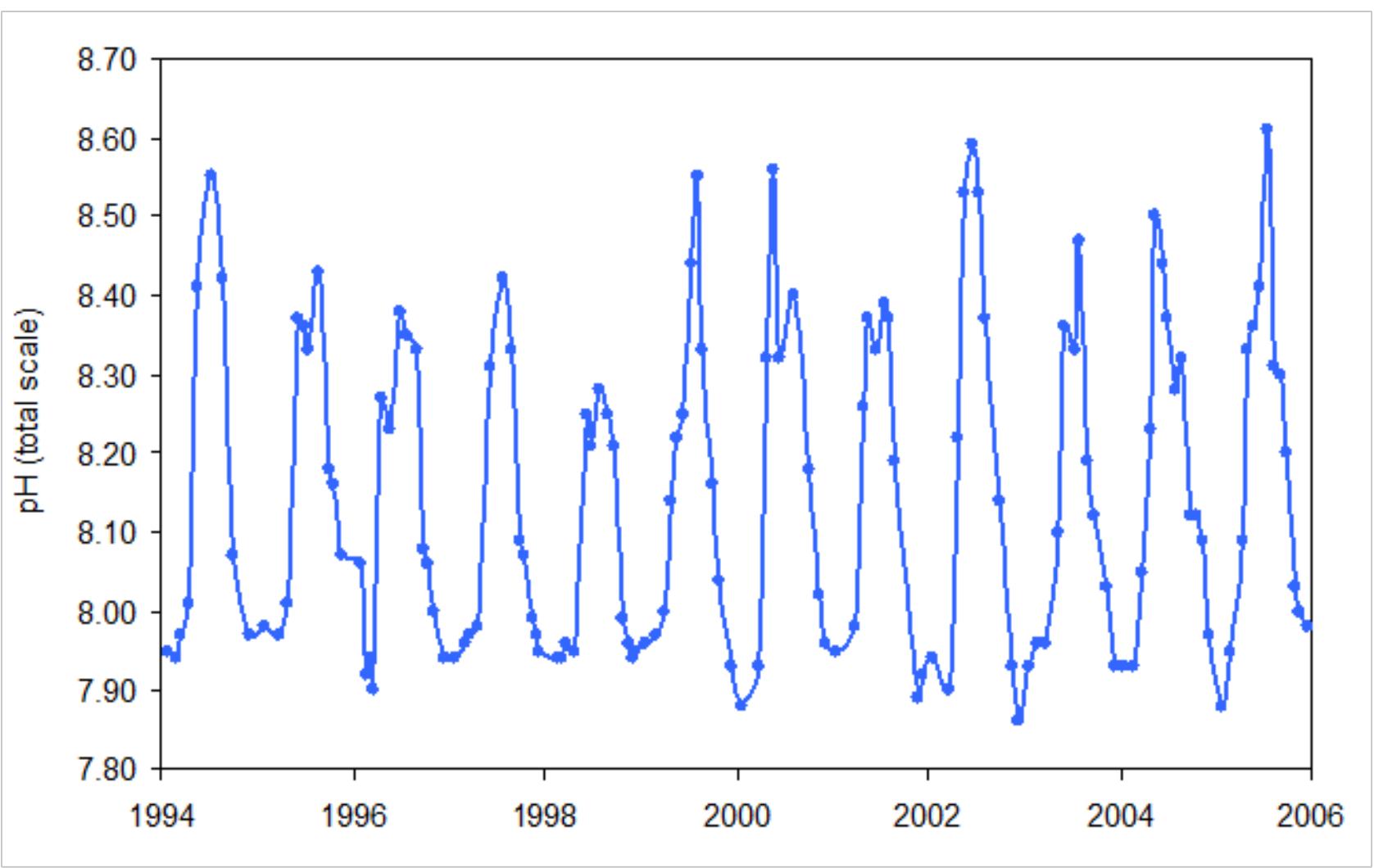
Source: NASA



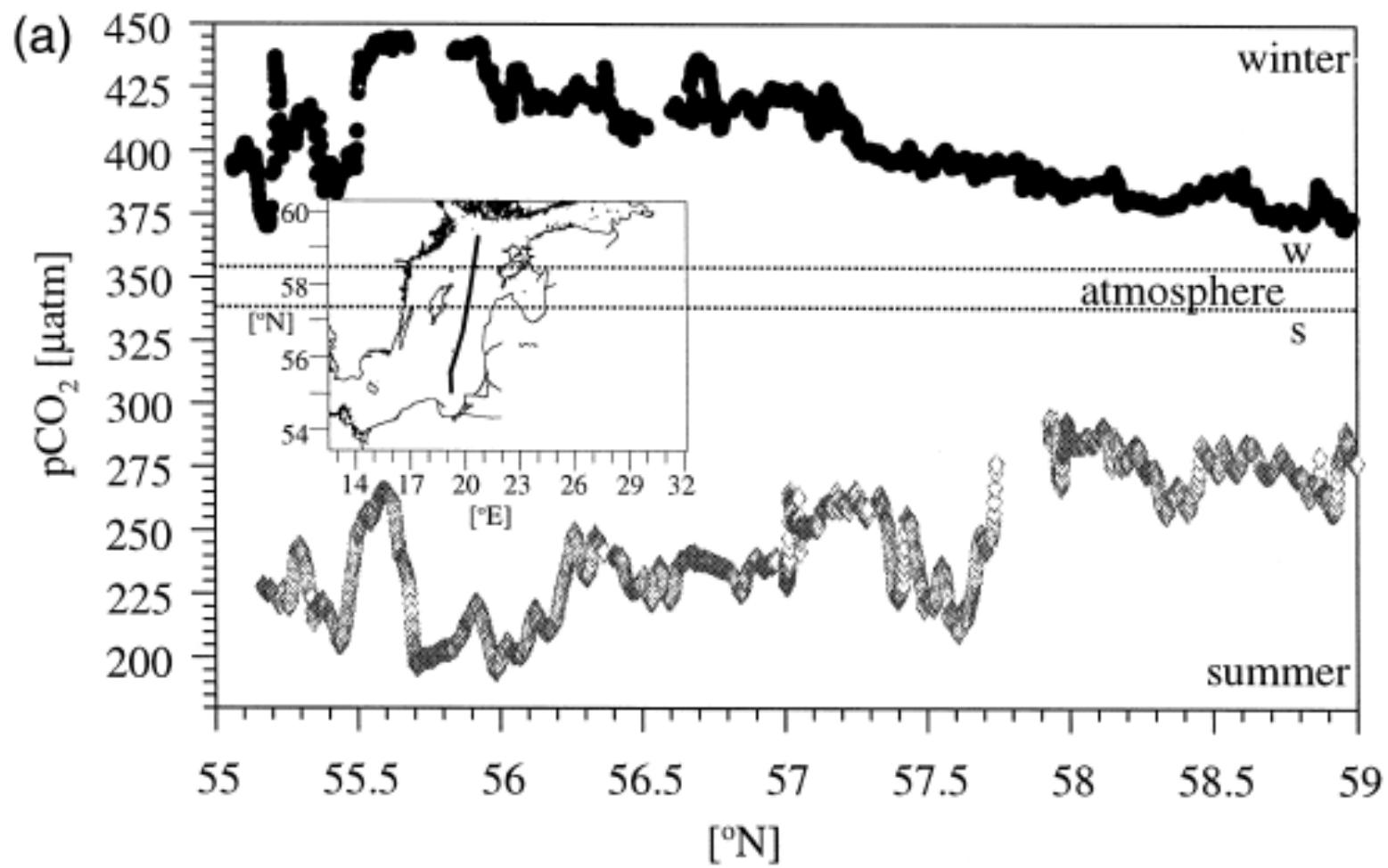
Source: PAP





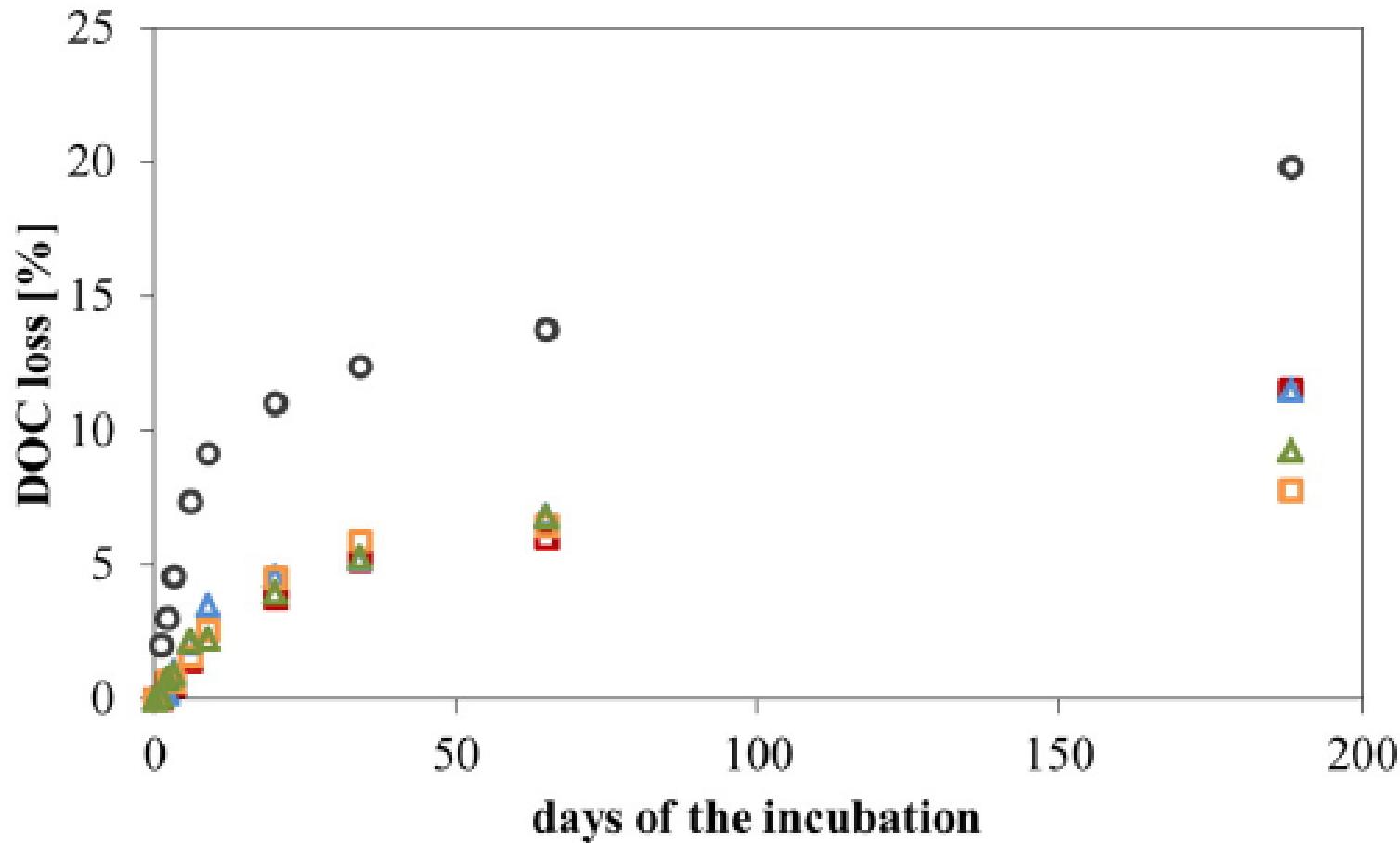


Source: SMHI

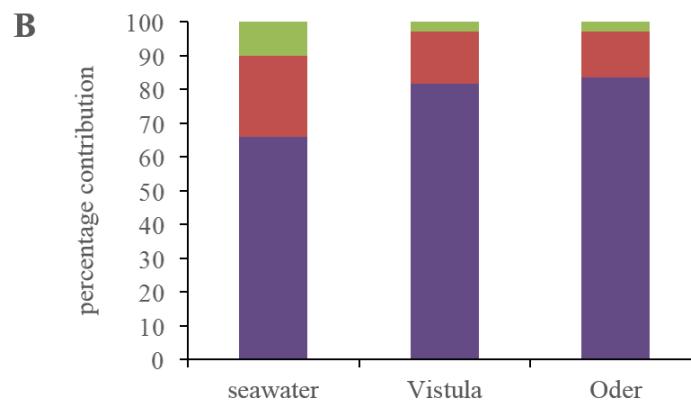
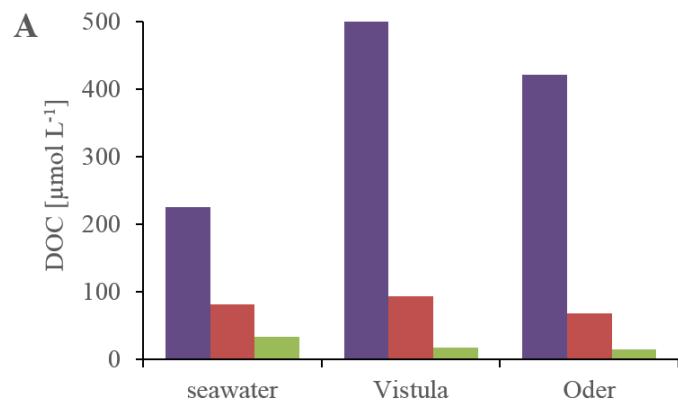


Source: Thomas & Schneider, 2009

$$DOC_{(t)} = \boxed{DOC_{L(t=0)} \cdot e^{-k_{b(L)} \cdot t}} + \boxed{DOC_{SL(t=0)} \cdot e^{-k_{b(SL)} \cdot t}} + \boxed{DOC_R}$$



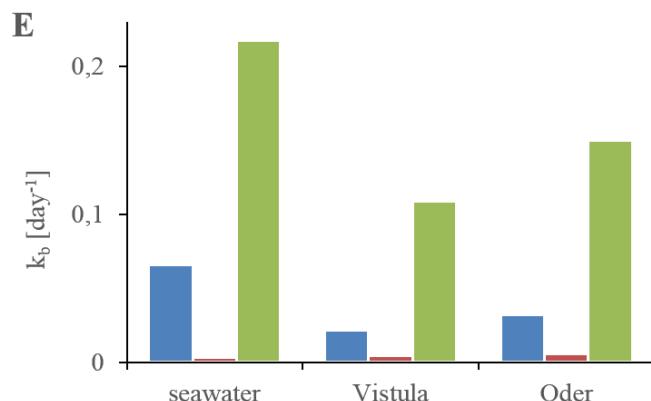
$$DOC_{(t)} = DOC_{L(t=0)} \cdot e^{-k_b(L) \cdot t} + DOC_{SL(t=0)} \cdot e^{-k_b(SL) \cdot t} + DOC_R$$



■ DOC_R

■ DOC_{SL}

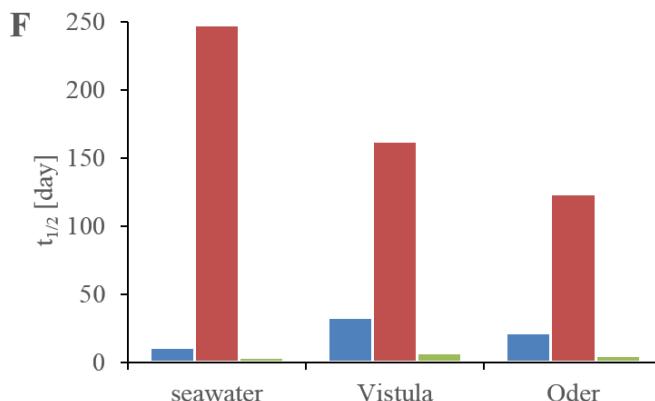
■ DOC_L



■ DOC_B

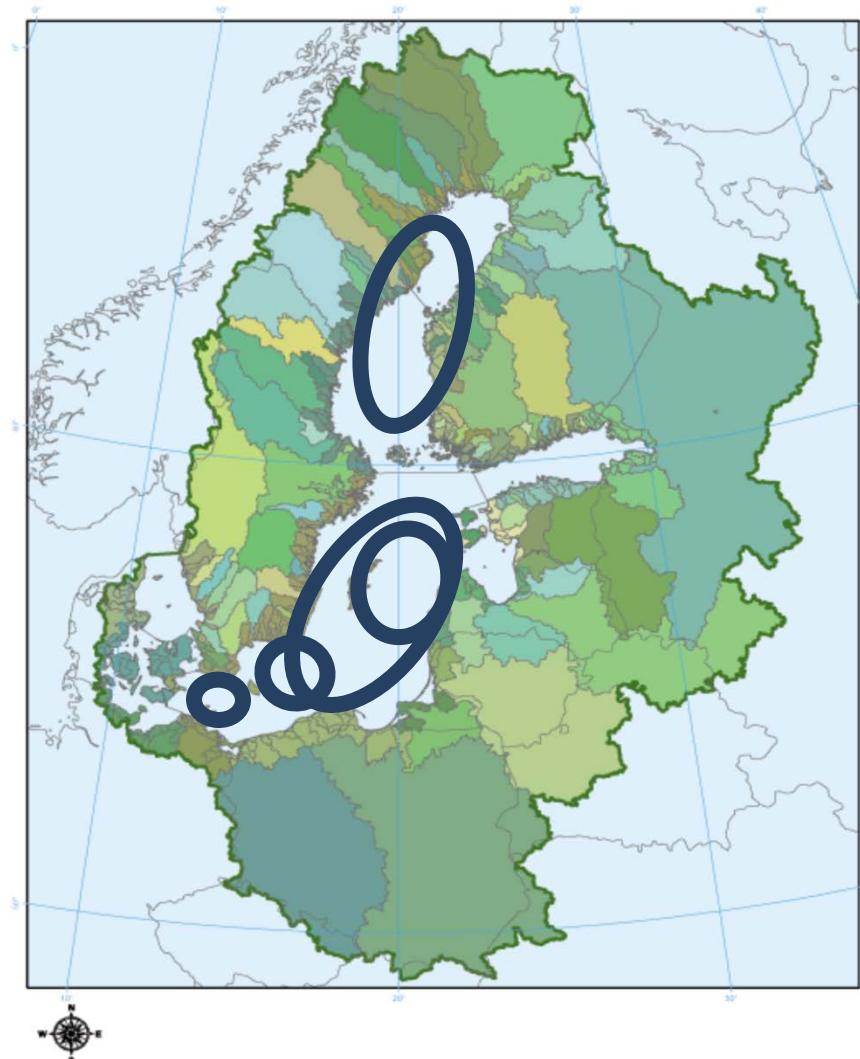
■ DOC_{SL}

■ DOC_L

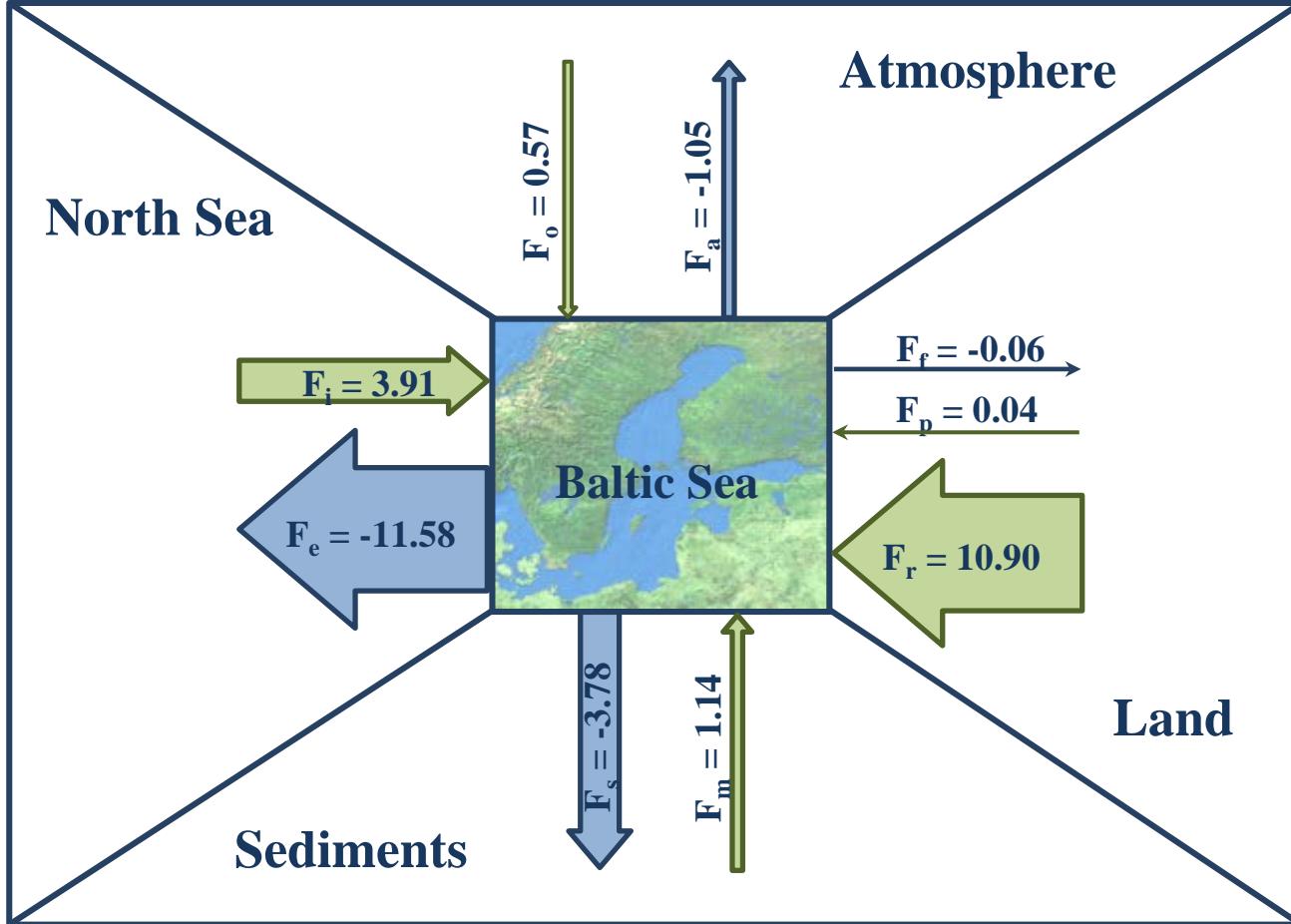


Baltic Sea sink or source???

- **10.8 g C m⁻² yr⁻¹**
(Thomas et al., 2003)
- **36.0 g C m⁻² yr⁻¹**
(Kuss et al., 2006)
- **-35.4 g C m⁻² yr⁻¹**
(Algesten et al., 2006)
- **-19.7 g C m⁻² yr⁻¹**
(Wesslander et al., 2010)
- **-28.1 g C m⁻² yr⁻¹**
(Wesslander et al., 2010)



HELCOM, 2007



Values are in $Tg (10^{12} g) C yr^{-1}$

Net CO_2 emission to the atmosphere?

$$-1.05 \pm 1.71 \text{ Tg C yr}^{-1}$$

$$-3.0 \pm 4.88 \text{ g C m}^{-2} \text{ yr}^{-1}$$

Source: Kuliński & Pempkowiak, 2011

River run-off
IC: 62%
OC: 38%

Import from the North Sea
IC: 95%
OC: 5%

Export to the North Sea
IC: 83%
OC: 17%

Return flux from the sediments
IC: 91%
OC: 9%

Seawater acid-base system

Measurable parameters:

- C_T – total CO_2 concentration (DIC)

$$C_T = [\text{CO}_2]^* + [\text{HCO}_3^-] + [\text{CO}_3^{2-}]$$

- A_T – total alkalinity

$$A_T = [\text{HCO}_3^-] + 2[\text{CO}_3^{2-}] + [\text{B(OH)}_4^-] + [\text{OH}^-] + \dots - [\text{H}^+] - \dots$$

- $p\text{CO}_2$ – CO_2 partial pressure

- pH – spectrophotometric measurement with m-cresol purple, total scale

$$\text{pH}_T = -\log ([\text{H}^+]_F + [\text{HSO}_4^-]) = -\log [\text{H}^+]_T$$

It is possible to calculate 2 parameters when the following is known:

- other 2 parameters
- temperature & salinity
- equilibrium constants for each of the acid dissociation reactions
- total concentrations for each non- CO_2 substances

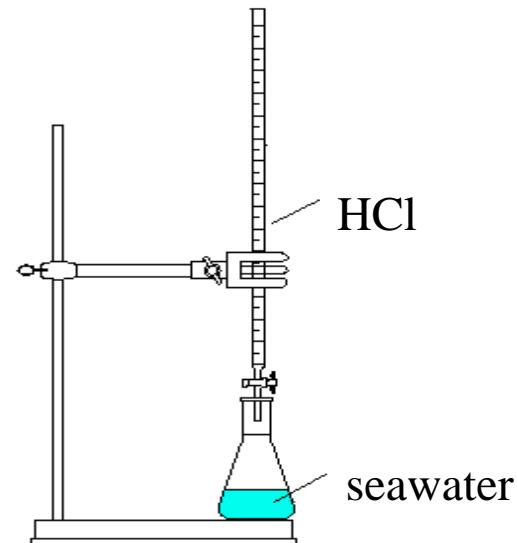
The pair used in the calculations:

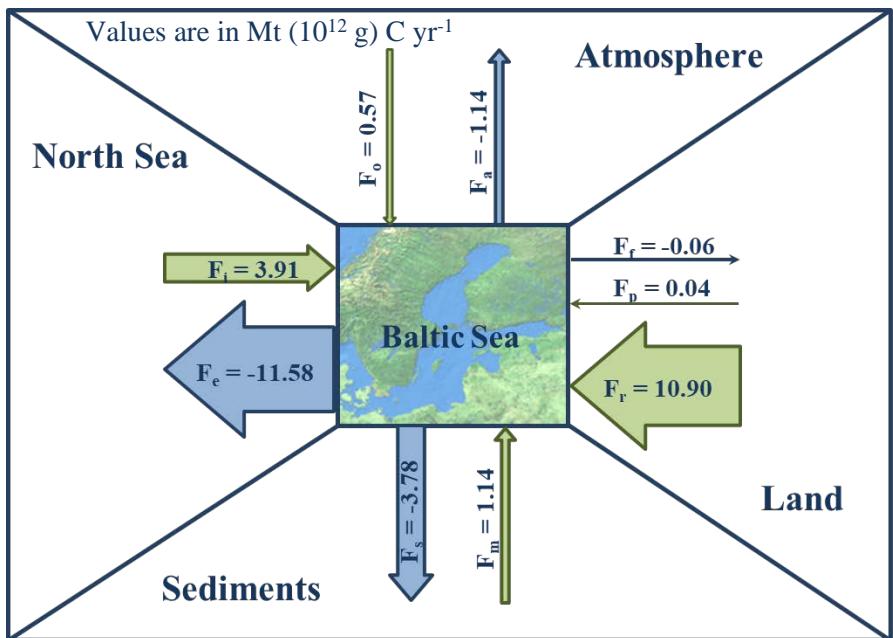
- C_T & A_T – recommended, used in biogeochemical modelling

Total alkalinity

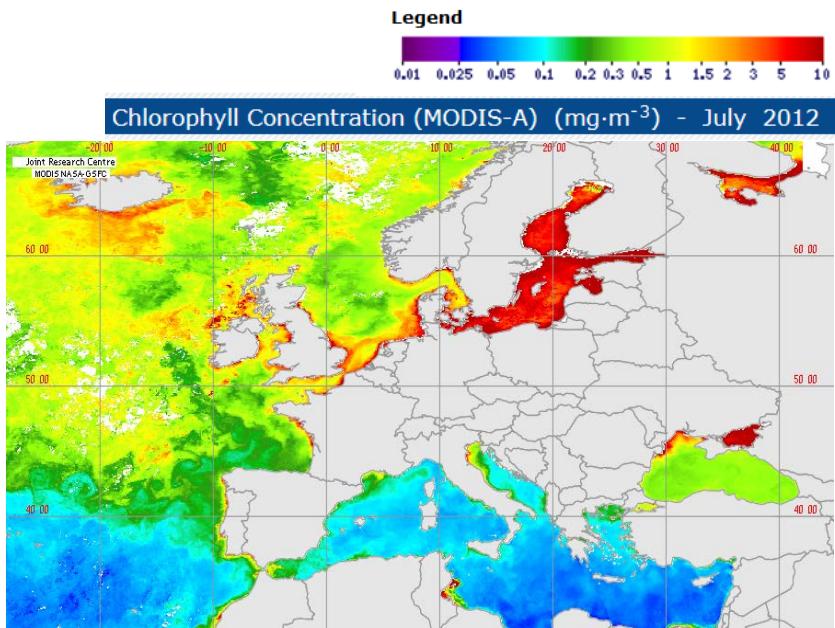
The total alkalinity of seawater is defined as the excess of proton acceptors (bases formed from weak acids with a dissociation constant $K \leq 10^{-4.5}$ at 25°C) over proton donors (acids with $K > 10^{-4.5}$) and expressed as a hydrogen ion equivalent in one kilogram of sample (Dickson, 1981):

$$\begin{aligned} A_T = & [HCO_3^-] + 2[CO_3^{2-}] + [B(OH)_4^-] + [OH^-] + [HPO_4^{2-}] \\ & + 2[PO_4^{3-}] + [SiO(OH)_3^-] + [NH_3] + [HS^-] + \dots + \text{minor bases} \\ - & [H^+]_{\text{wolny}} - [HSO_4^-] - [HF] - [H_3PO_4] - \dots - \text{minor acids} \end{aligned}$$

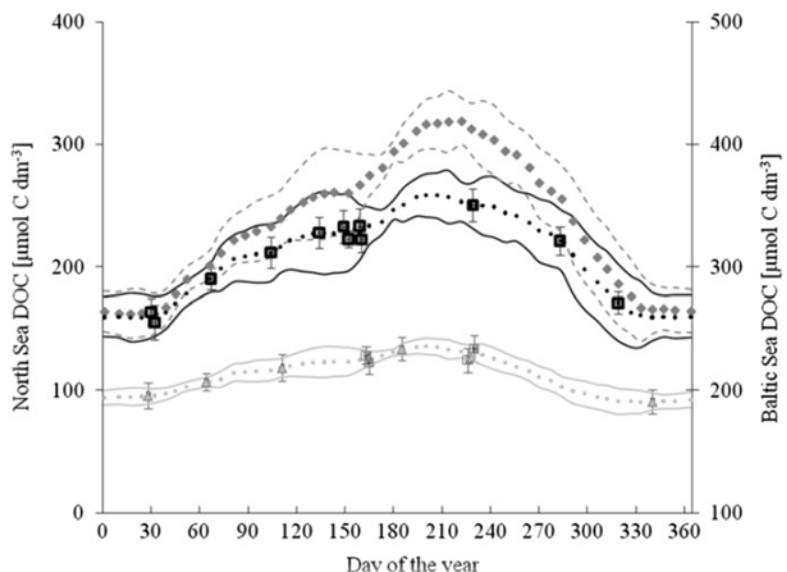




Source: Kuliński & Pempkowiak, 2011



Source: HELCOM

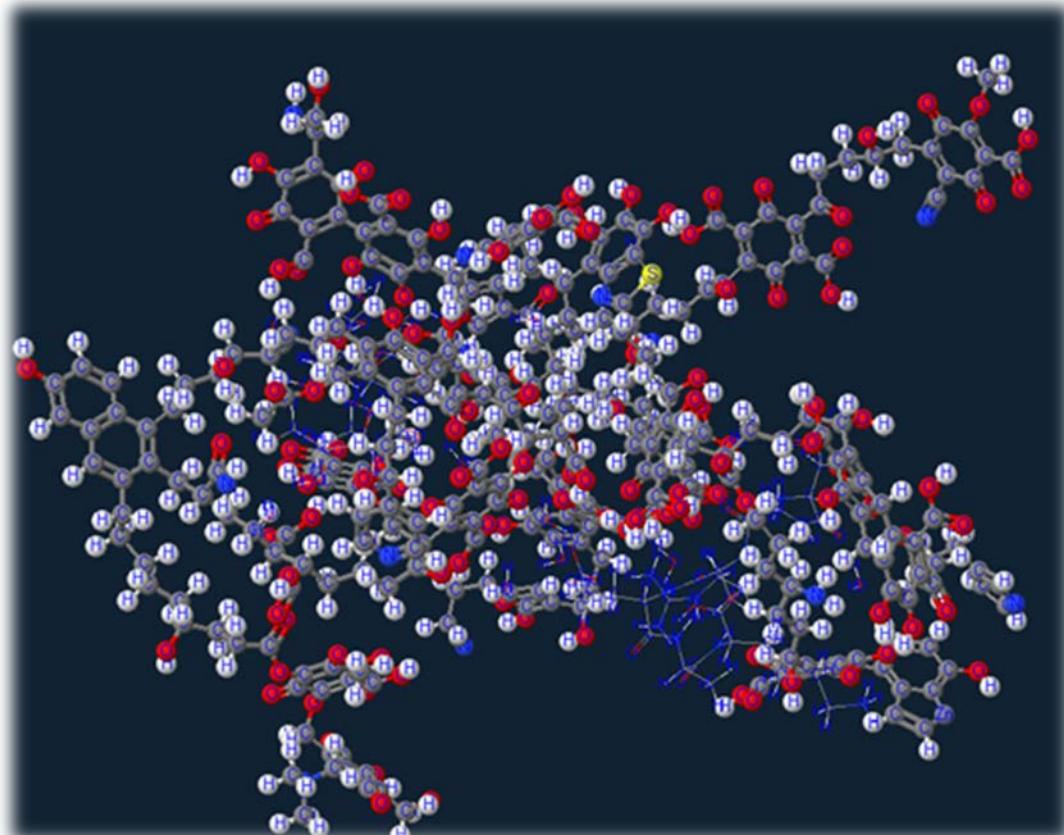


Source: Kuliński et al., 2011



Functional groups in DOM

Group	Structure	Exchange H ?
Alcohol	$\begin{array}{c} \\ -C-O-H \\ \end{array}$	Yes
Phenol	$\begin{array}{c} \text{C}_6\text{H}_5 \\ \\ O-H \end{array}$	Yes
Ether	$\begin{array}{c} \\ -C-O-C- \\ \end{array}$	
Aldehyde	$\begin{array}{c} \\ -C-C(=O)-H \\ \end{array}$	No
Ketone	$\begin{array}{c} O \\ \\ -C-C-C- \\ \end{array}$	
Carboxyl	$\begin{array}{c} O \\ \\ -C-C(=O)-O-H \\ \end{array}$	Yes
Ester	$\begin{array}{c} O \\ \\ -C-C(=O)-O-C- \\ \end{array}$	
Amine	$\begin{array}{c} \\ -C-N \\ \end{array}$	Yes
Amide	$\begin{array}{c} O \\ \\ -C-C(=O)-N \\ \end{array}$	Yes



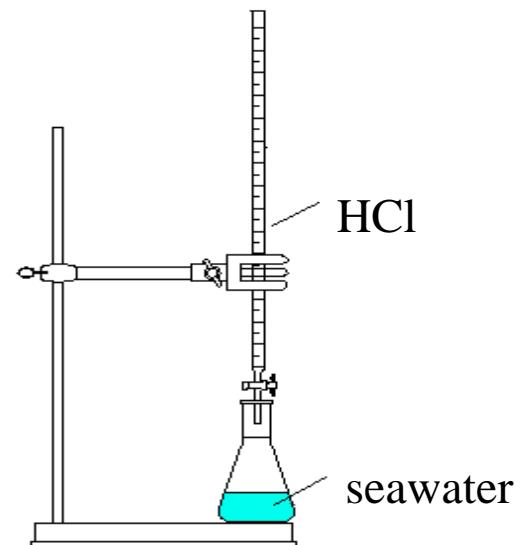
hypothetical structure of humic-like substances

Total alkalinity

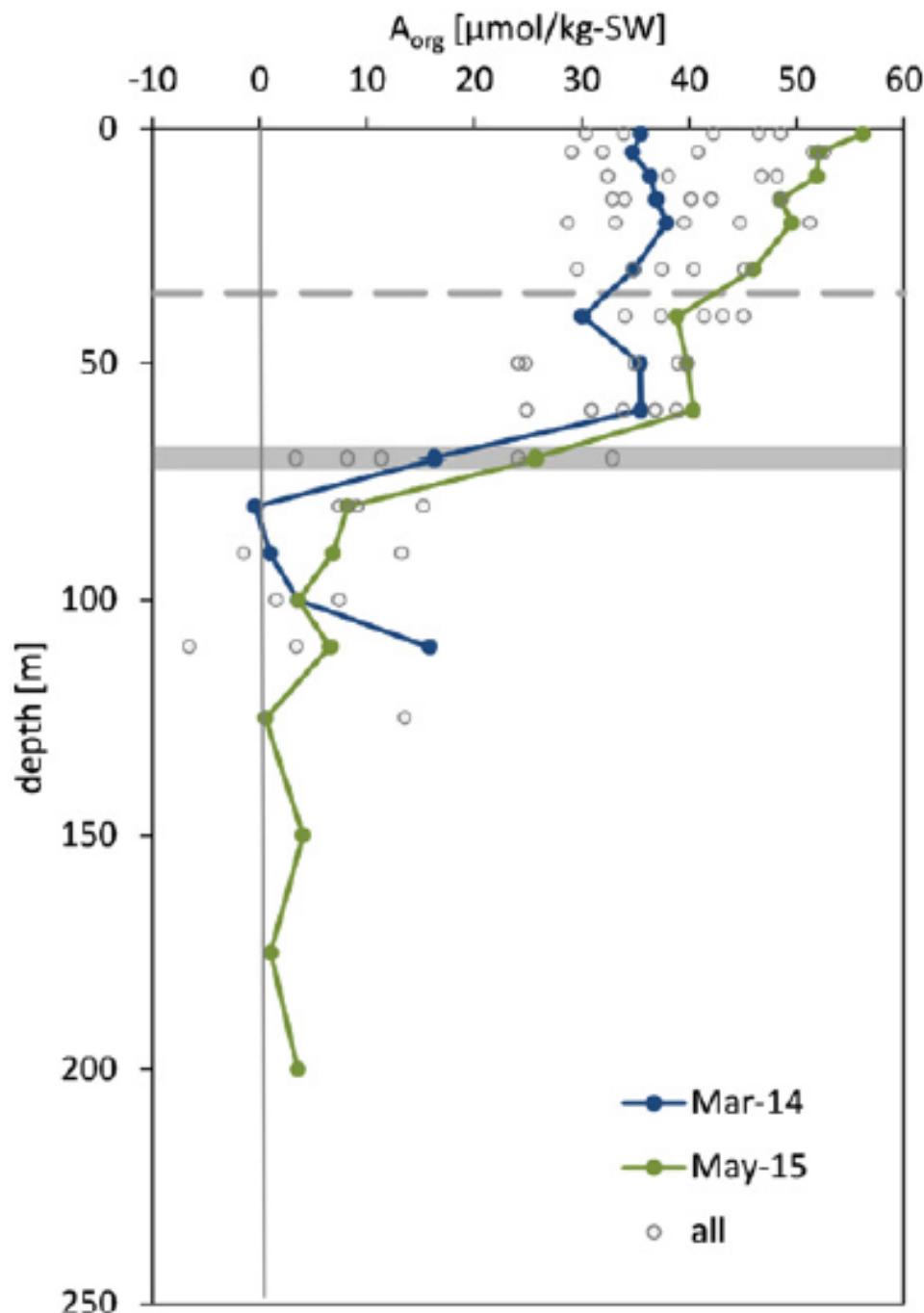
The total alkalinity of seawater is defined as the excess of proton acceptors (bases formed from weak acids with a dissociation constant $K \leq 10^{-4.5}$ at 25°C) over proton donors (acids with $K > 10^{-4.5}$) and expressed as a hydrogen ion equivalent in one kilogram of sample (Dickson, 1981):

$$\begin{aligned} A_T = & [HCO_3^-] + 2[CO_3^{2-}] + [B(OH)_4^-] + [OH^-] + [HPO_4^{2-}] \\ & + 2[PO_4^{3-}] + [SiO(OH)_3^-] + [NH_3] + [HS^-] + \dots + \text{Organics} \\ & - [H^+]_{\text{wolny}} - [HSO_4^-] - [HF] - [H_3PO_4] - \dots \end{aligned}$$

$$A_T = A_{\text{inorganic}} + A_{\text{org}}$$



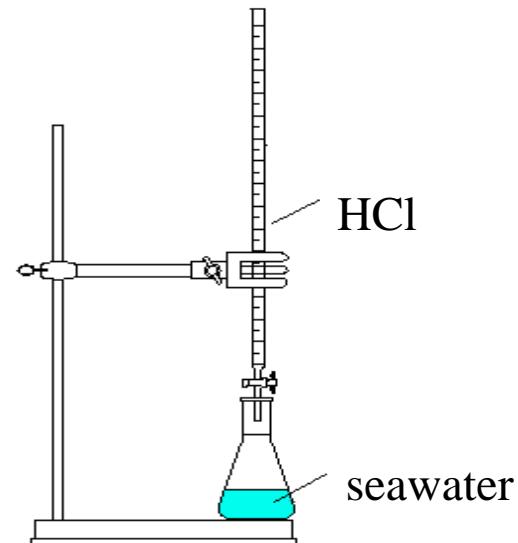
$$f = 0.12$$
$$pK_{DOM} = 7.34$$

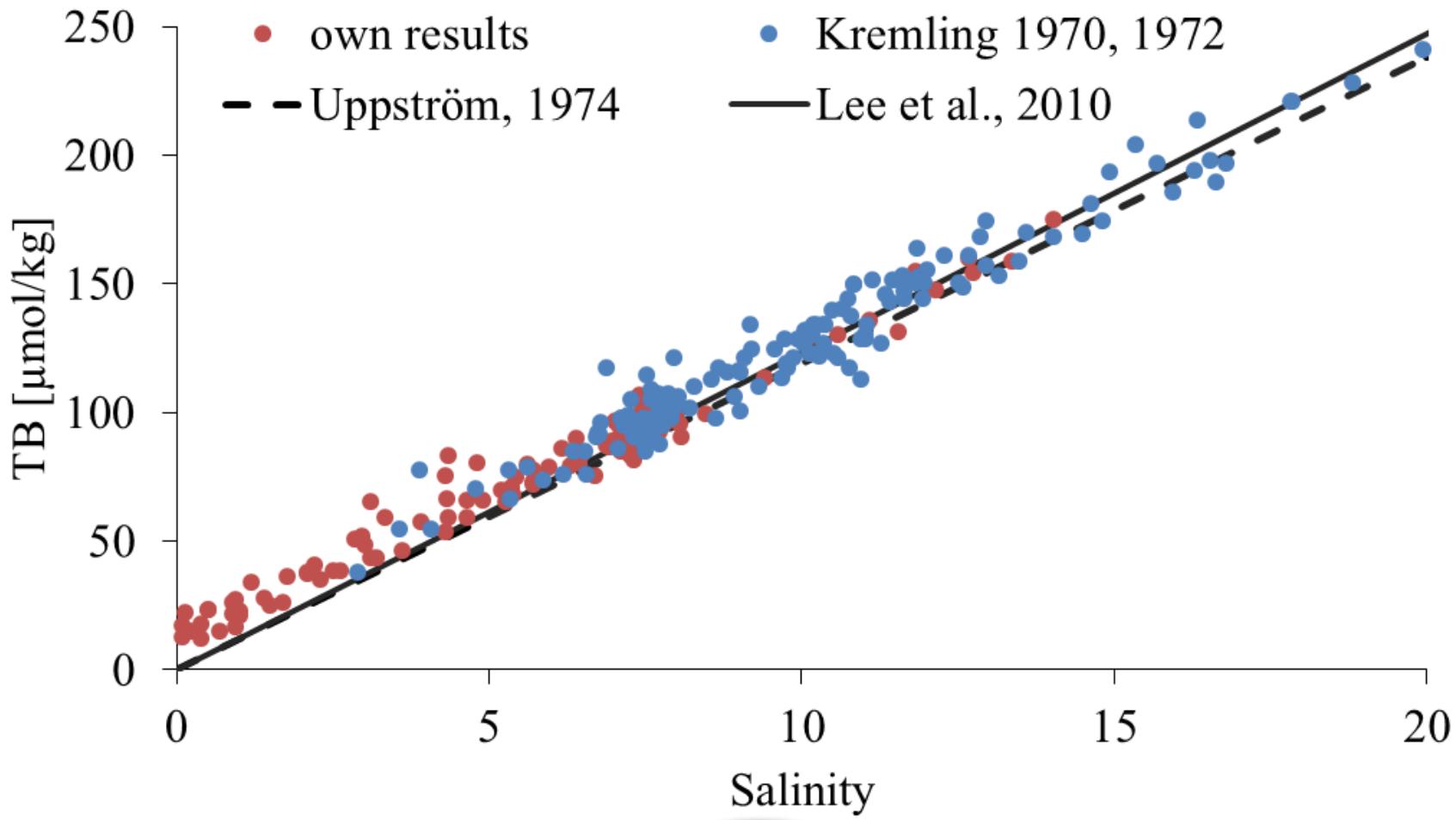


Total alkalinity

The total alkalinity of seawater is defined as the excess of proton acceptors (bases formed from weak acids with a dissociation constant $K \leq 10^{-4.5}$ at 25°C) over proton donors (acids with $K > 10^{-4.5}$) and expressed as a hydrogen ion equivalent in one kilogram of sample (Dickson, 1981):

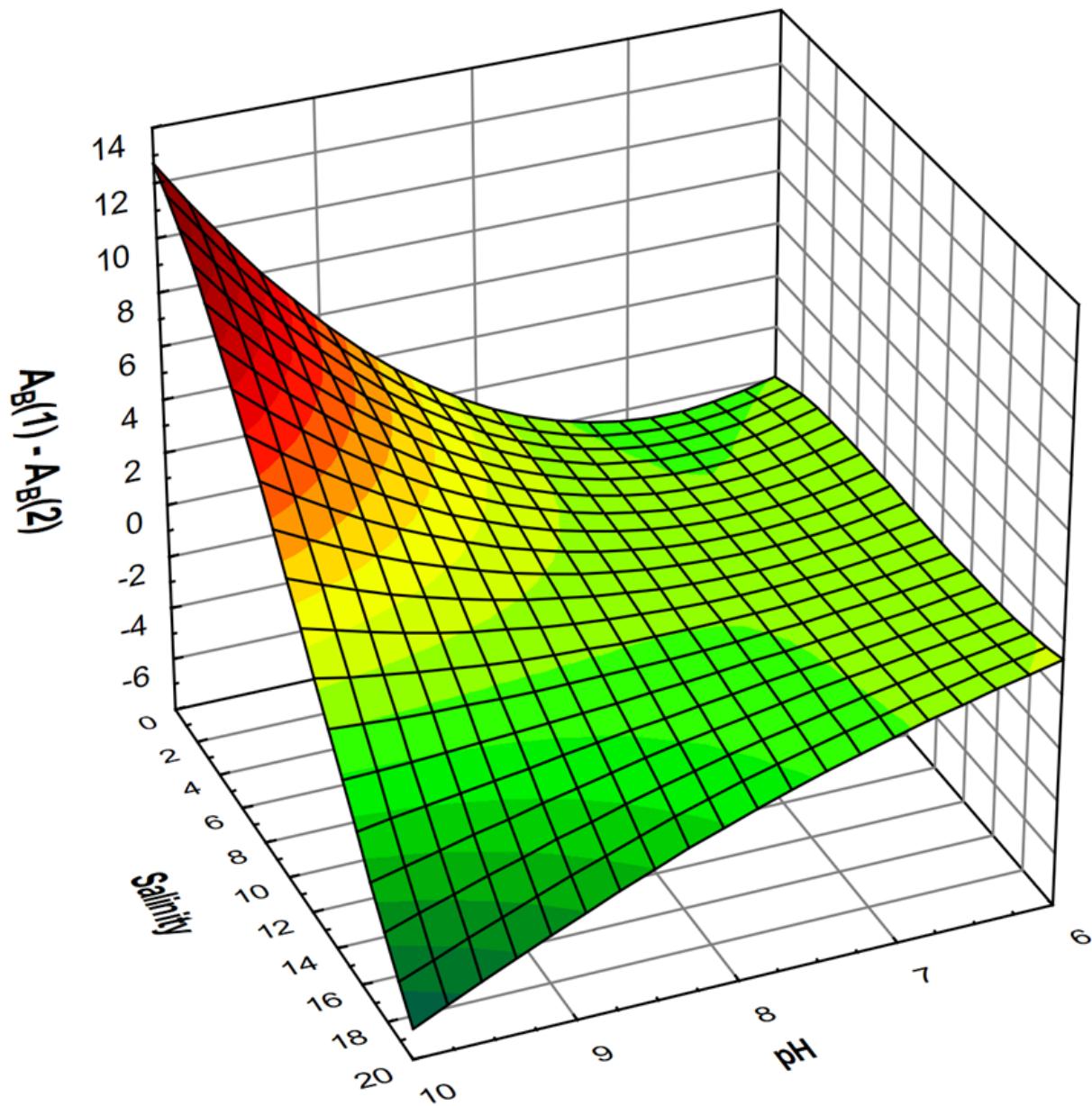
$$\begin{aligned} A_T = & [HCO_3^-] + 2[CO_3^{2-}] + [B(OH)_4^-] + [OH^-] + [HPO_4^{2-}] \\ & + 2[PO_4^{3-}] + [SiO(OH)_3^-] + [NH_3] + [HS^-] + \dots + \text{minor bases} \\ - & [H^+]_{\text{wolny}} - [HSO_4^-] - [HF] - [H_3PO_4] - \dots - \text{minor acids} \end{aligned}$$



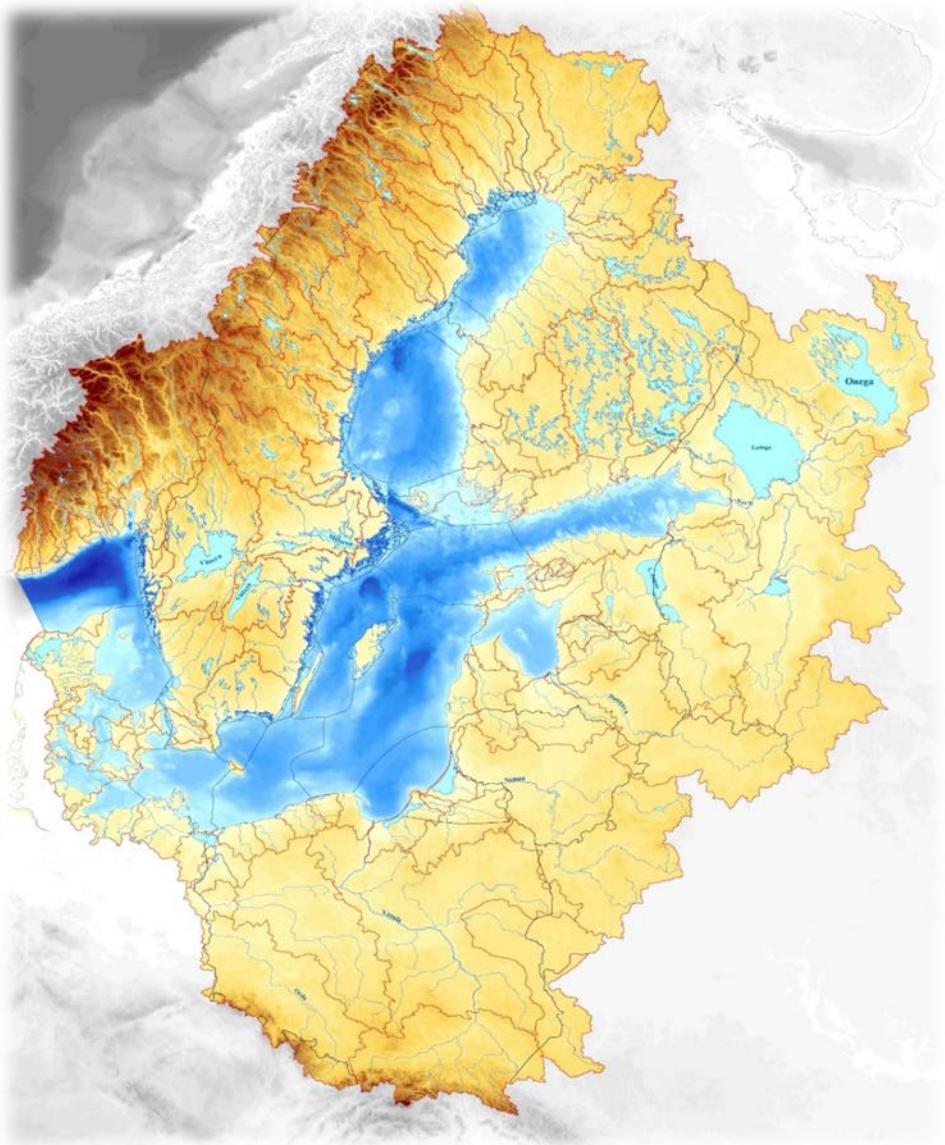


Source: Kuliński et al., 2018

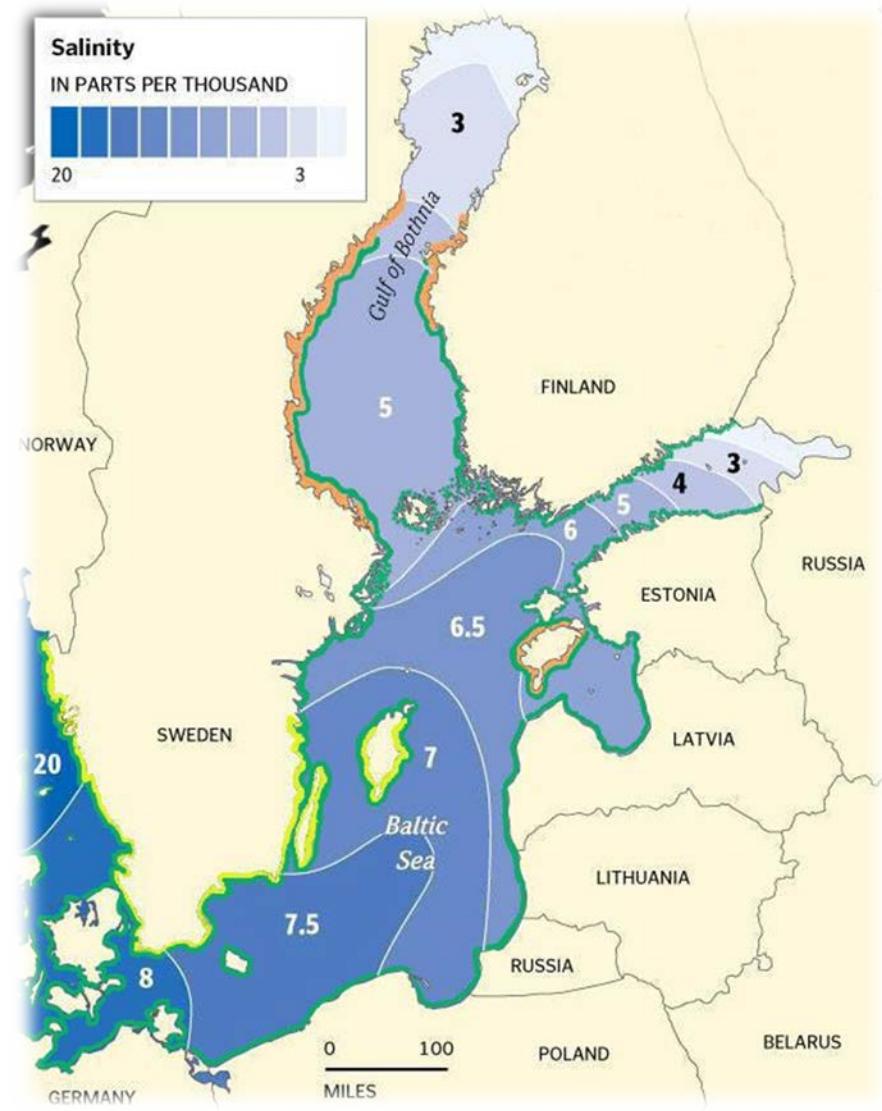
$- [\text{H}^+]_{\text{wolny}} - [\text{HSO}_4^-] - [\text{HF}] - [\text{H}_3\text{PO}_4] - \dots$



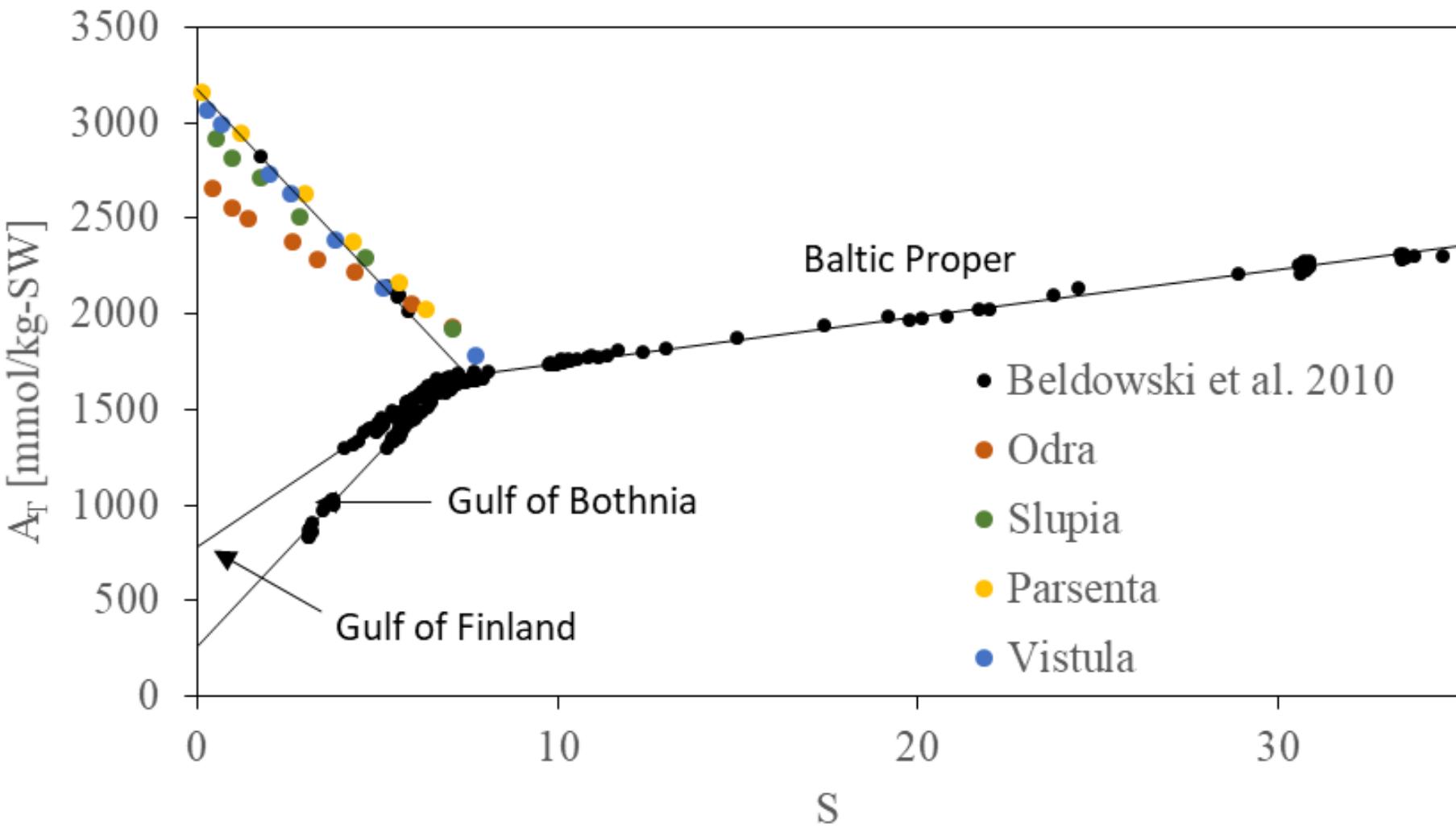
Source: Kuliński et al., 2018



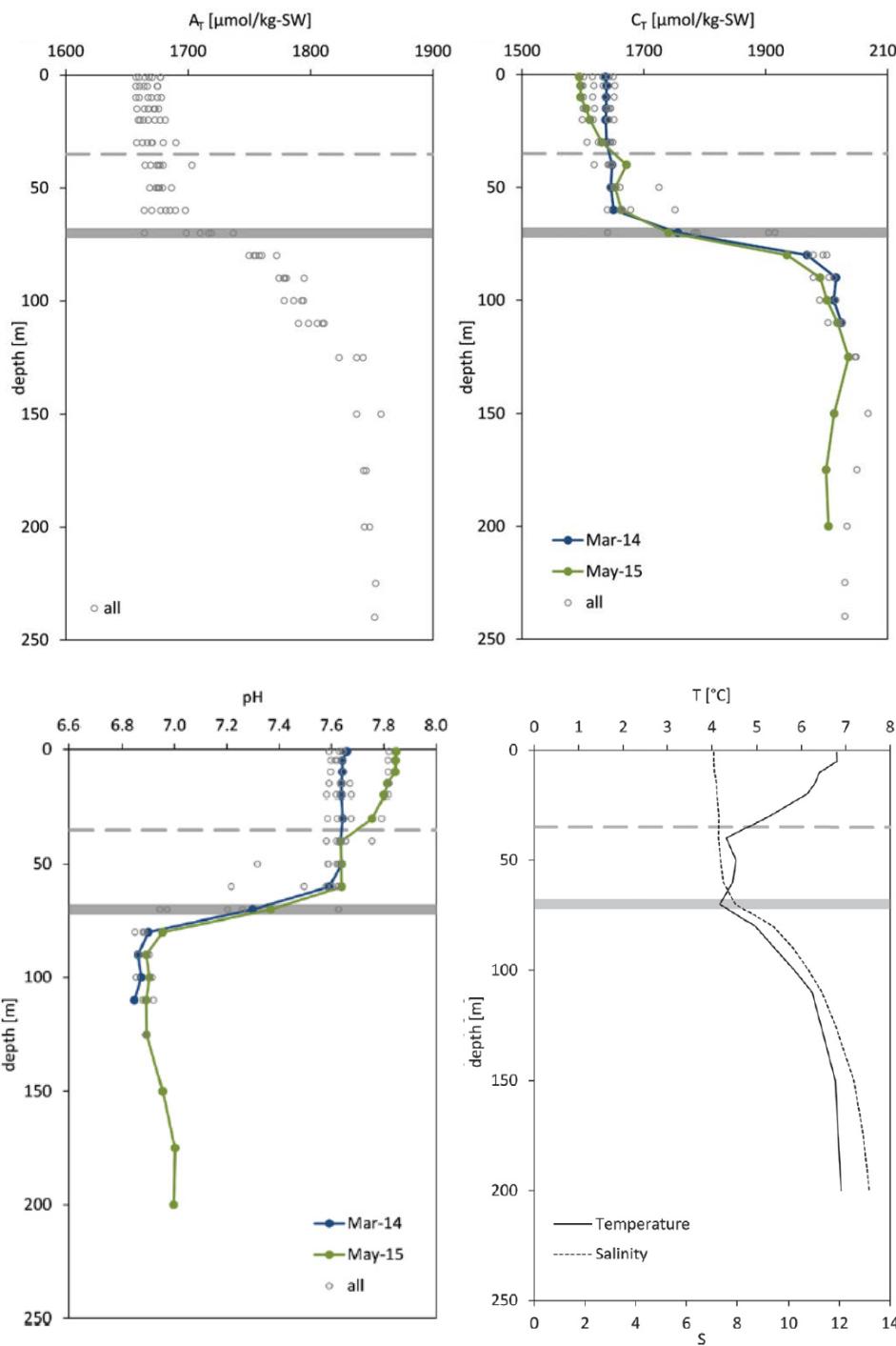
Source: SMHI



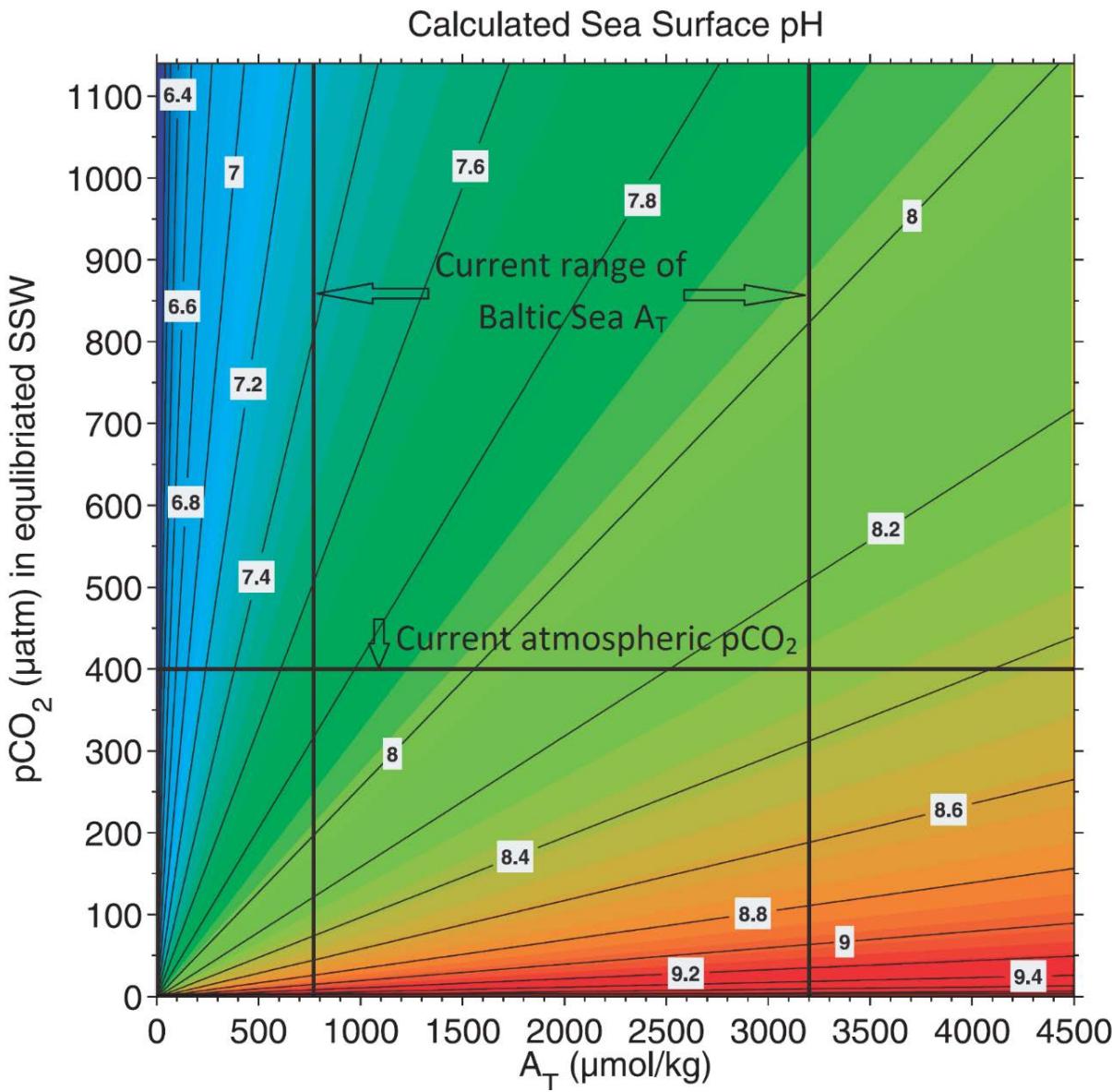
Source: balticseaweed.com



Source: Kuliński et al., in prep.

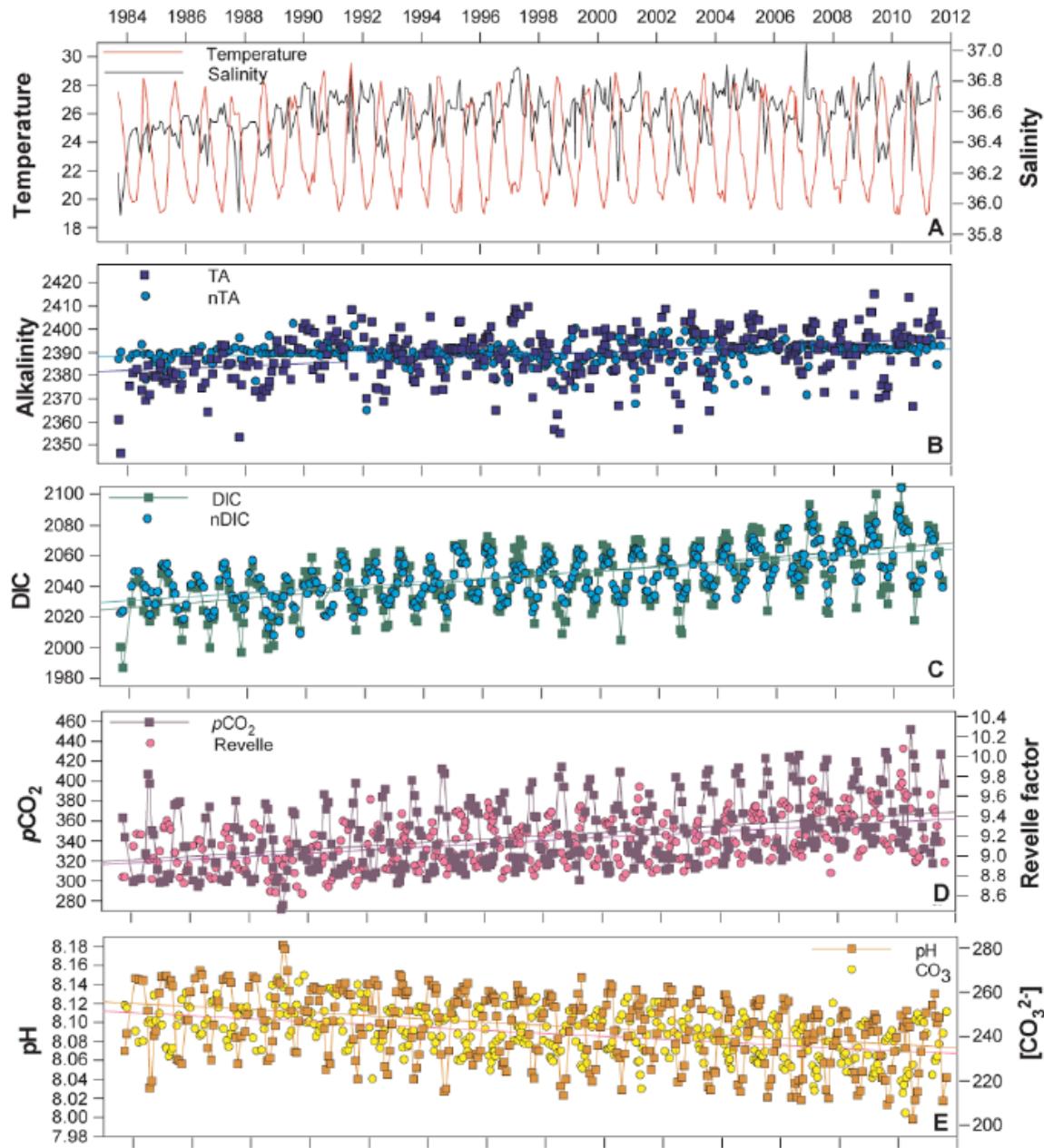


Source: Hammer et al., 2017

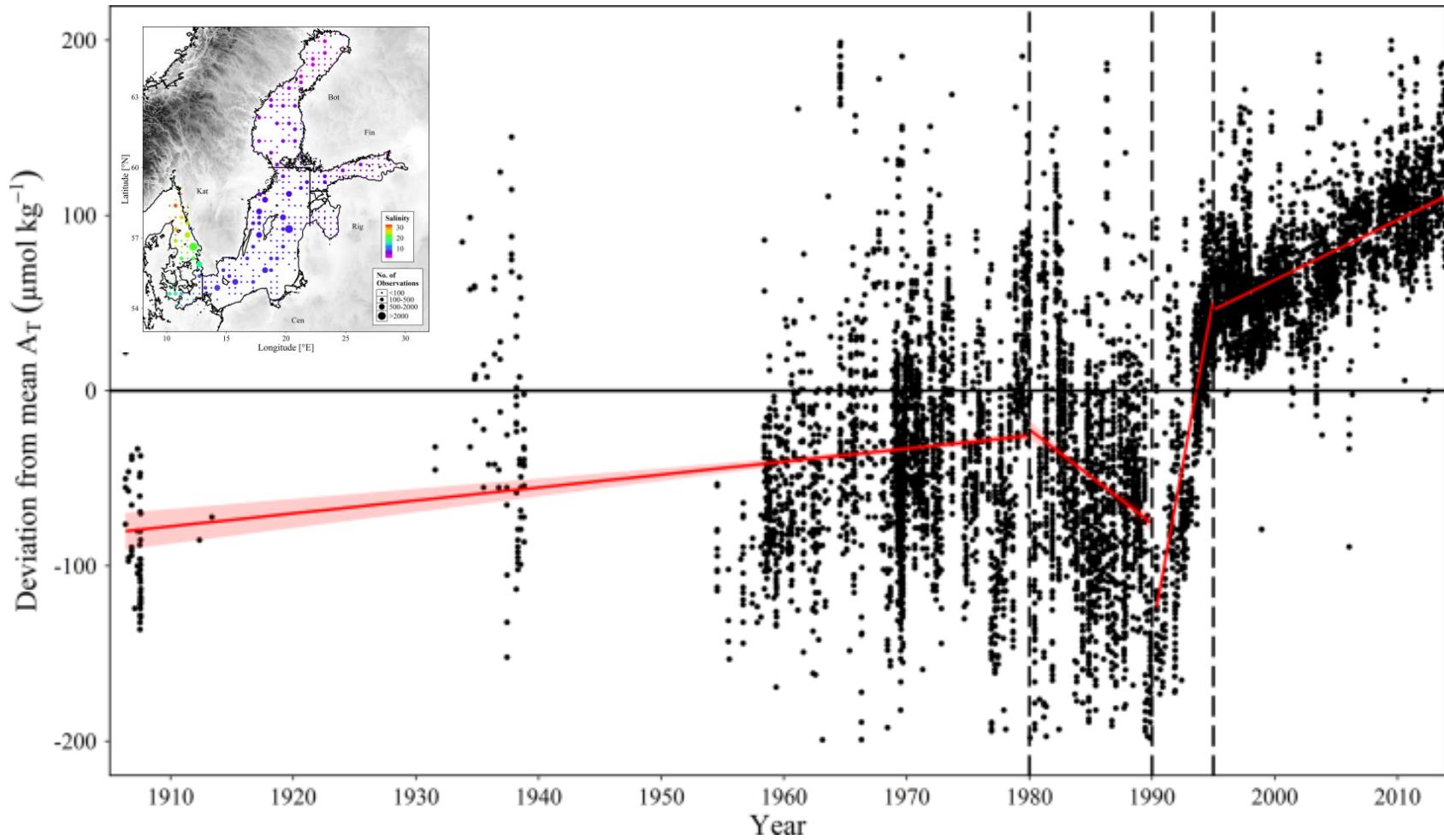


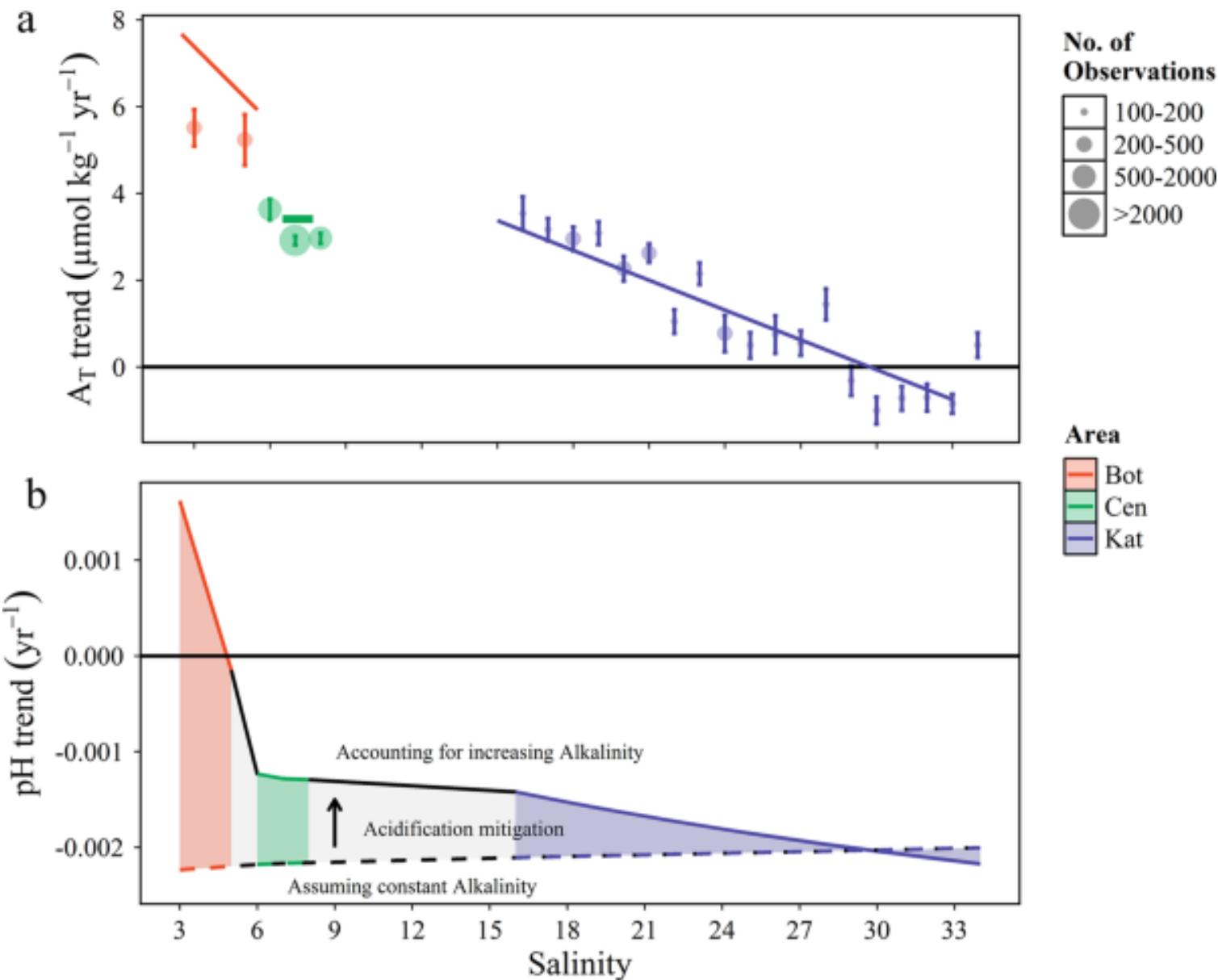
Source: Kuliński et al., 2017 modified after Omstedt et al., 2010

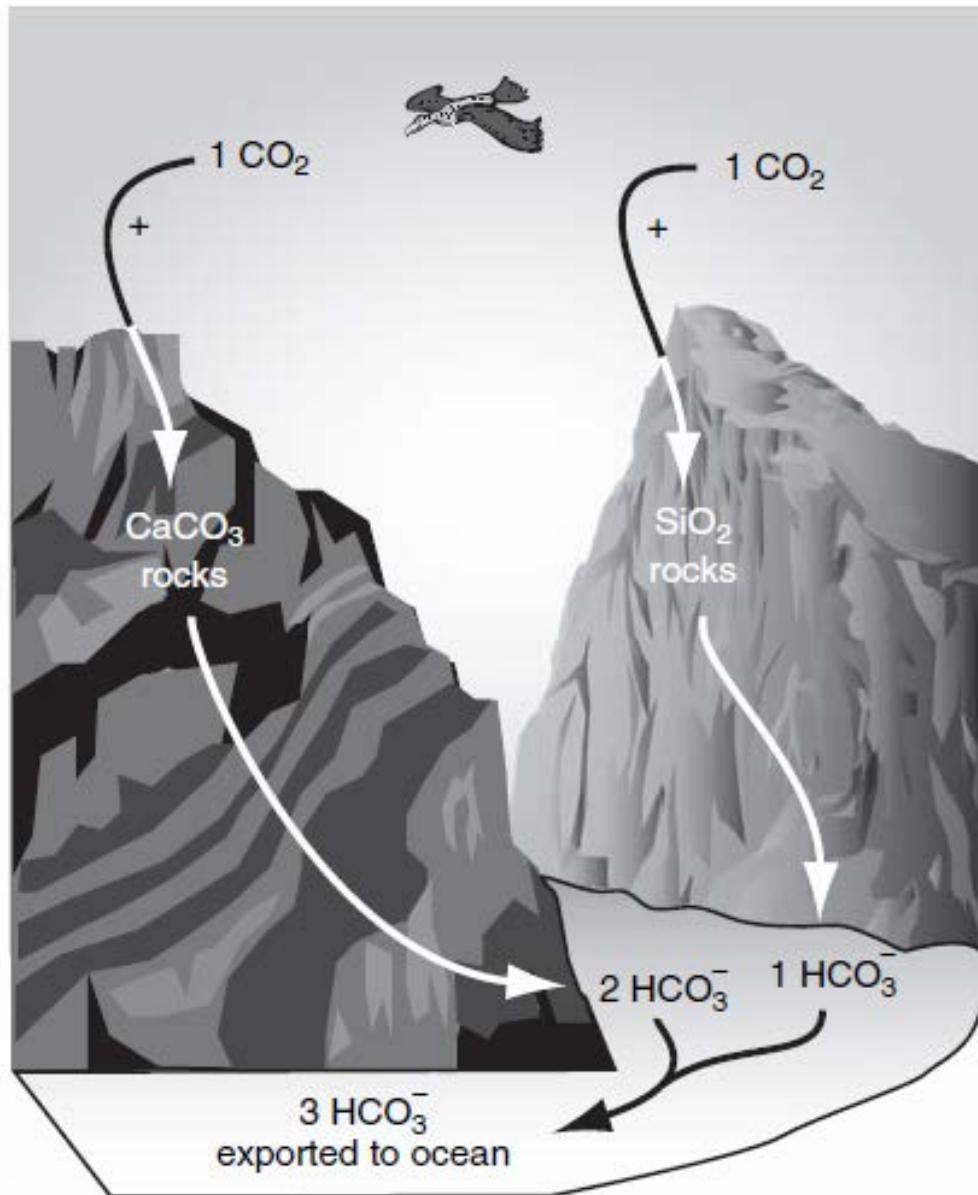
BATS - Bermuda Atlantic Time-series Study



Source: Bates et al., 2012

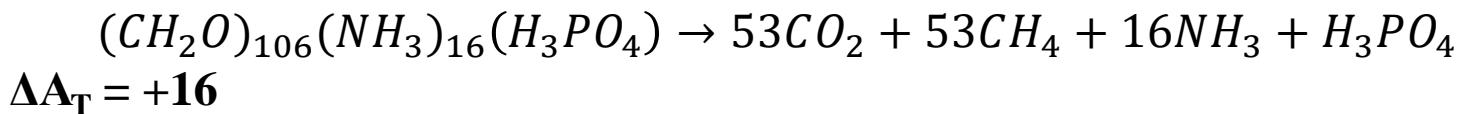
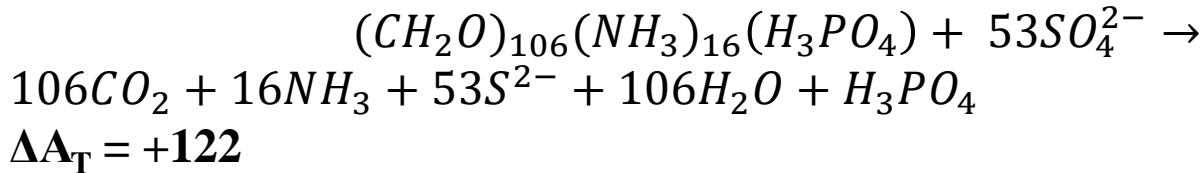
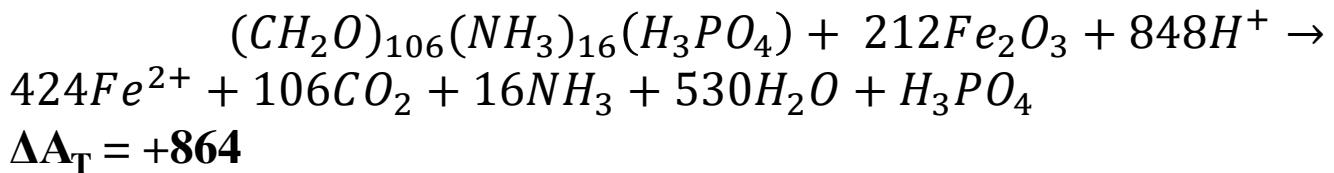
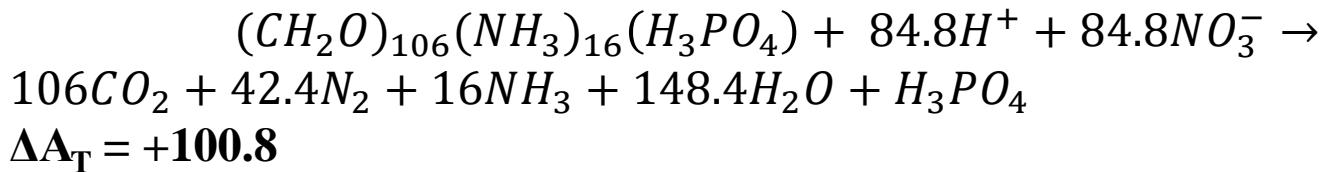
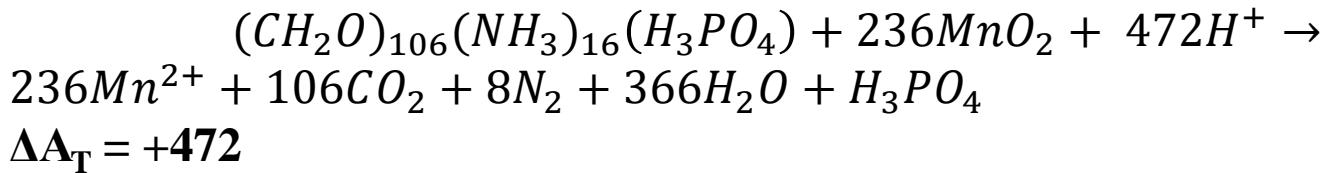
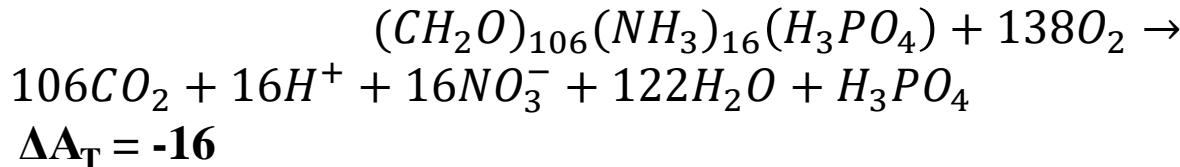


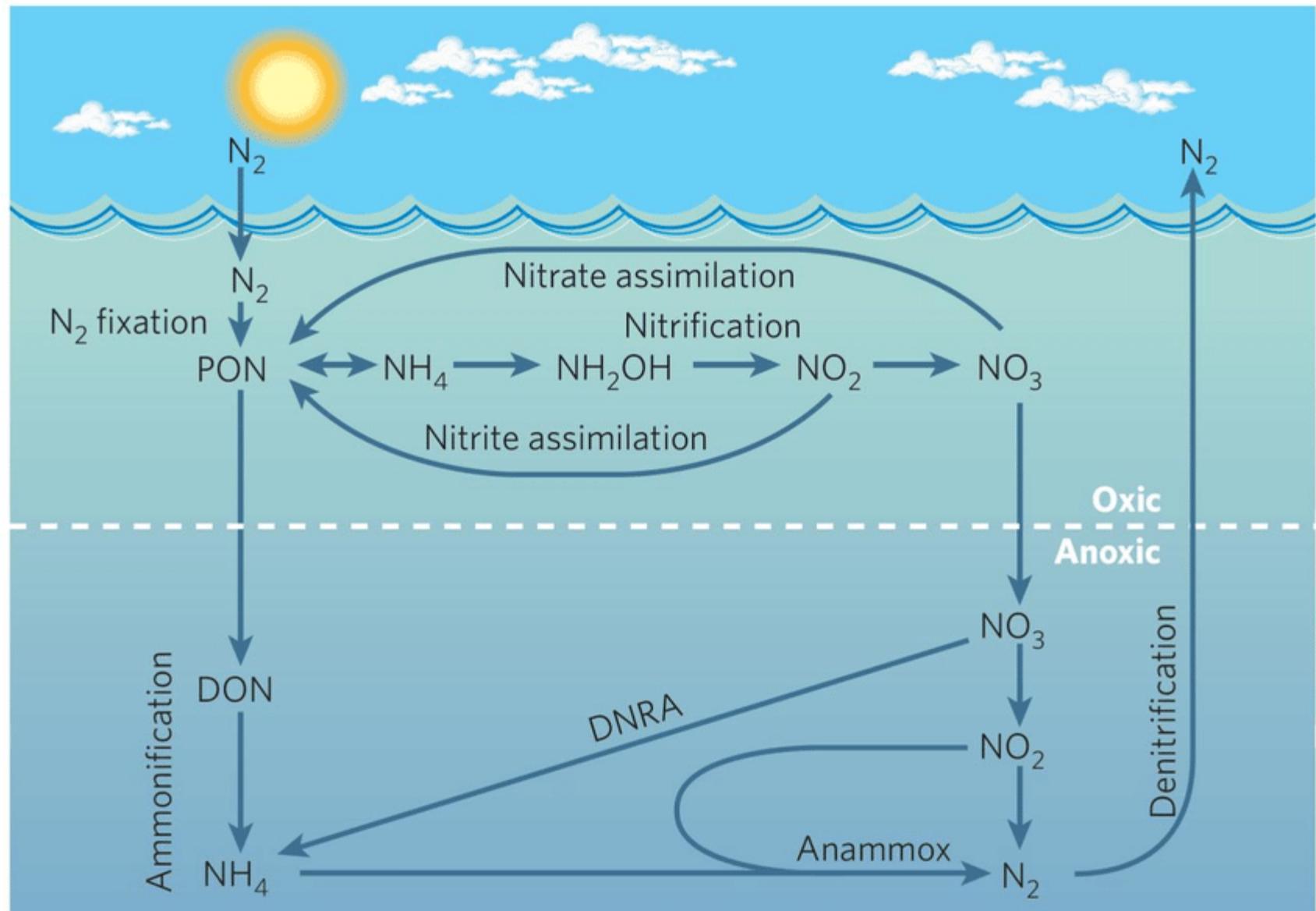


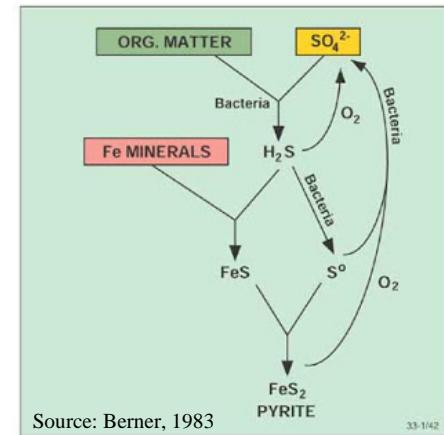
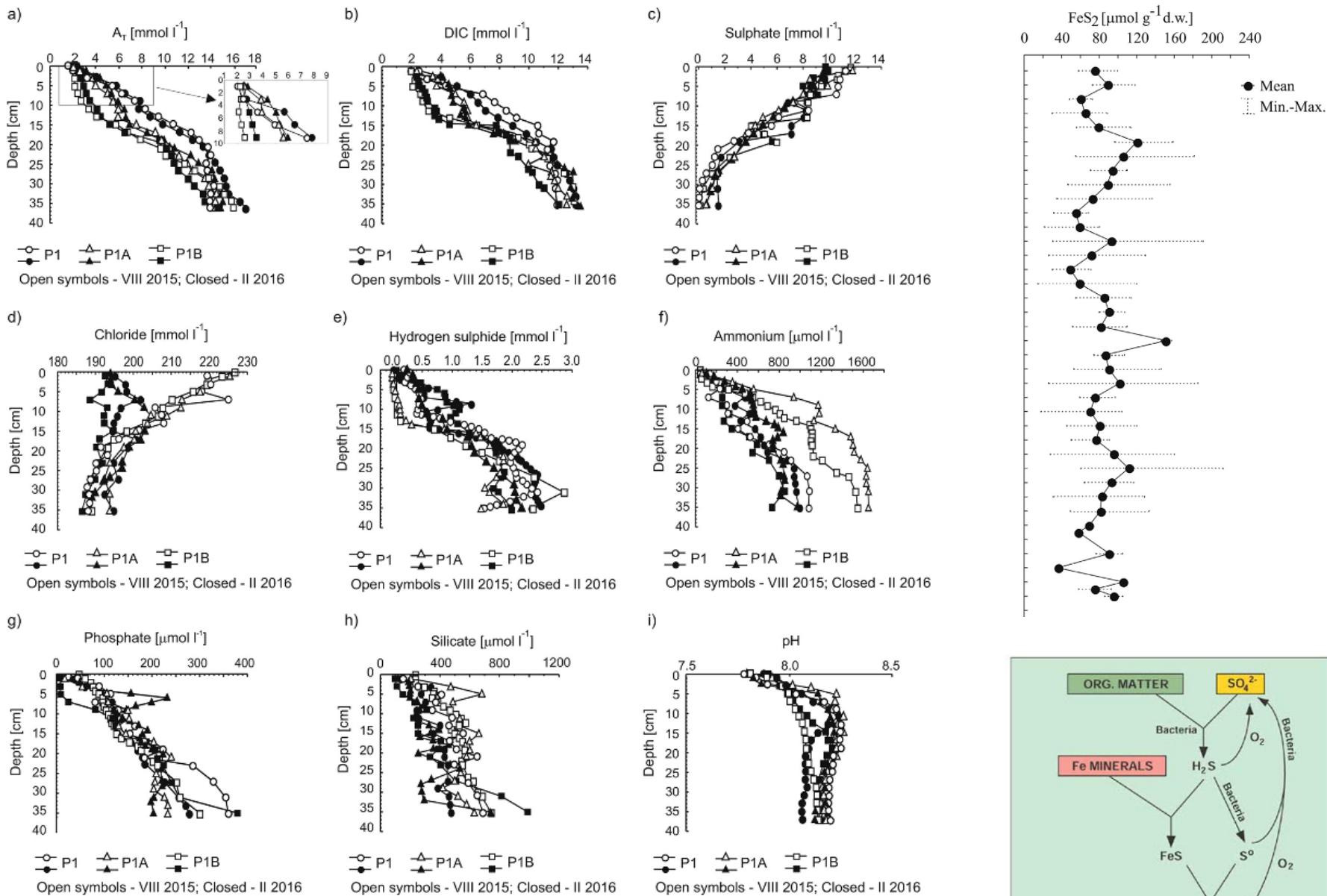


Source: Emerson & Hedges, 2008

Sediments – source of alkalinity

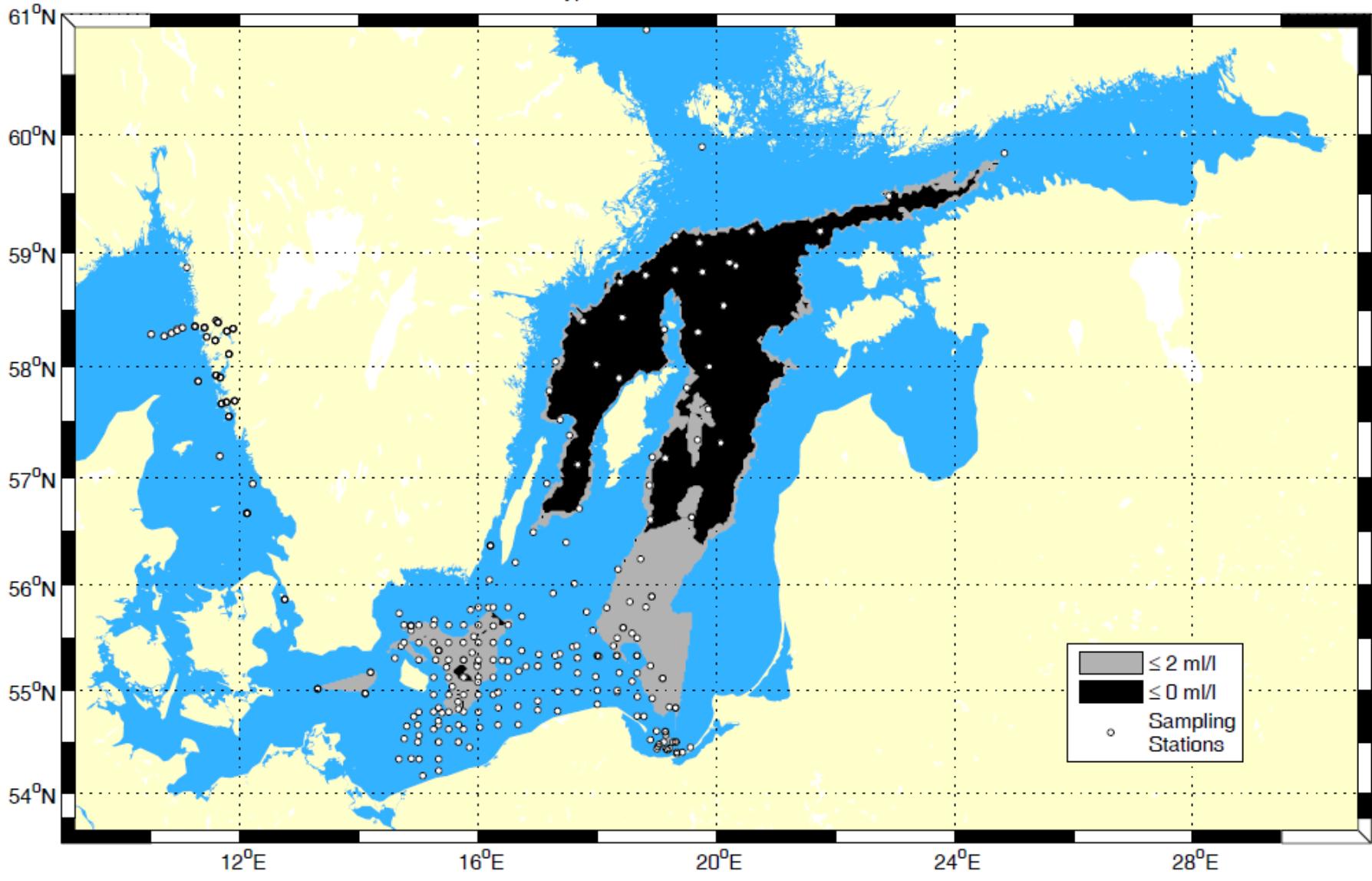






Source: Łukawska-Matuszewska & Graca, 2018

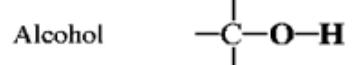
Extent of hypoxic & anoxic bottom water, Autumn 2017



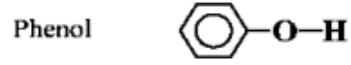
Created:
January
2018

Source: HELCOM

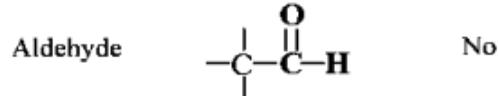
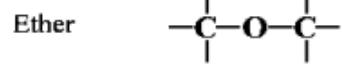
Group	Structure	Exchange H ?
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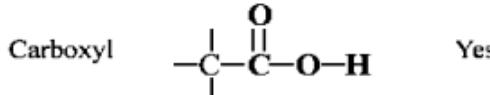
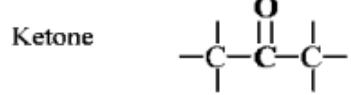
Yes



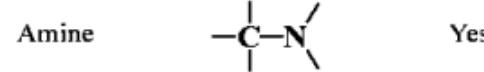
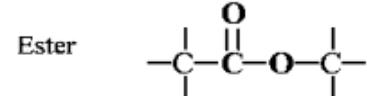
Yes



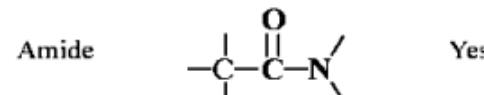
No



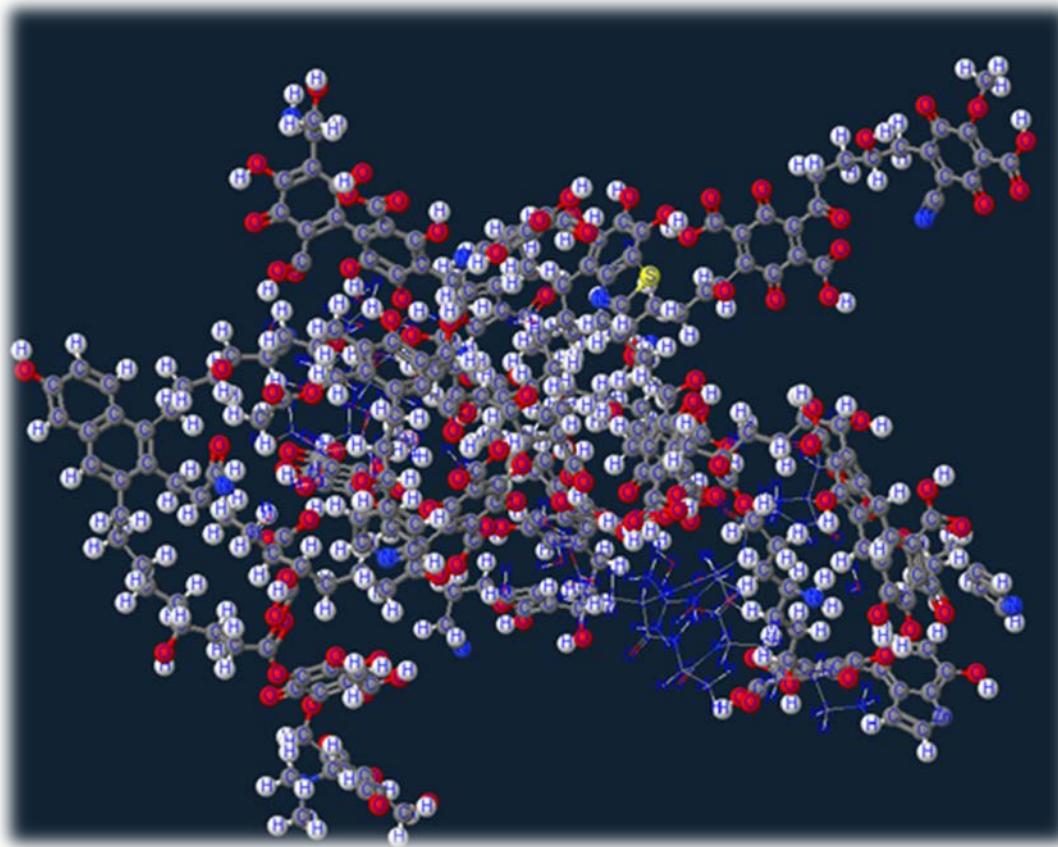
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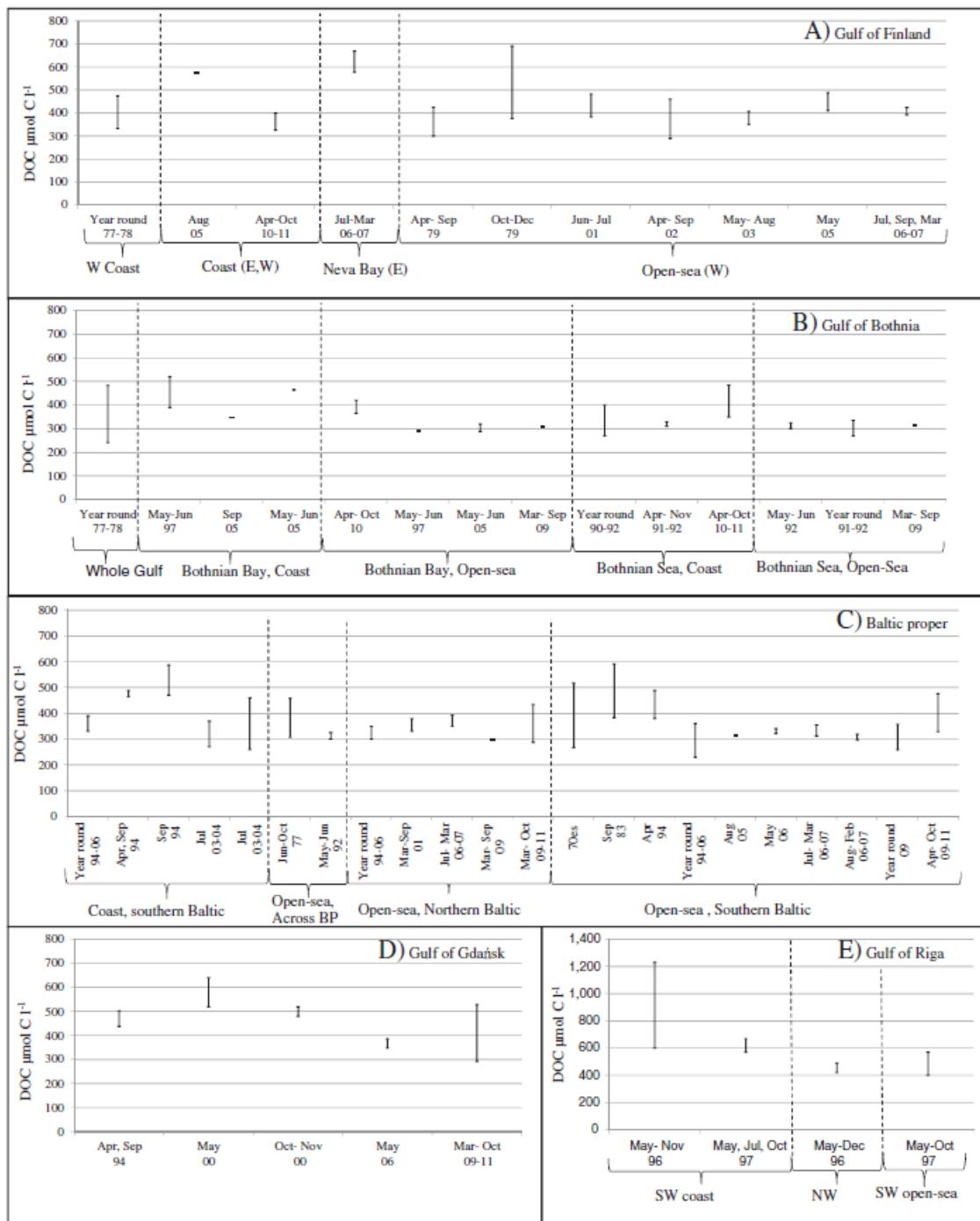
Yes



Yes

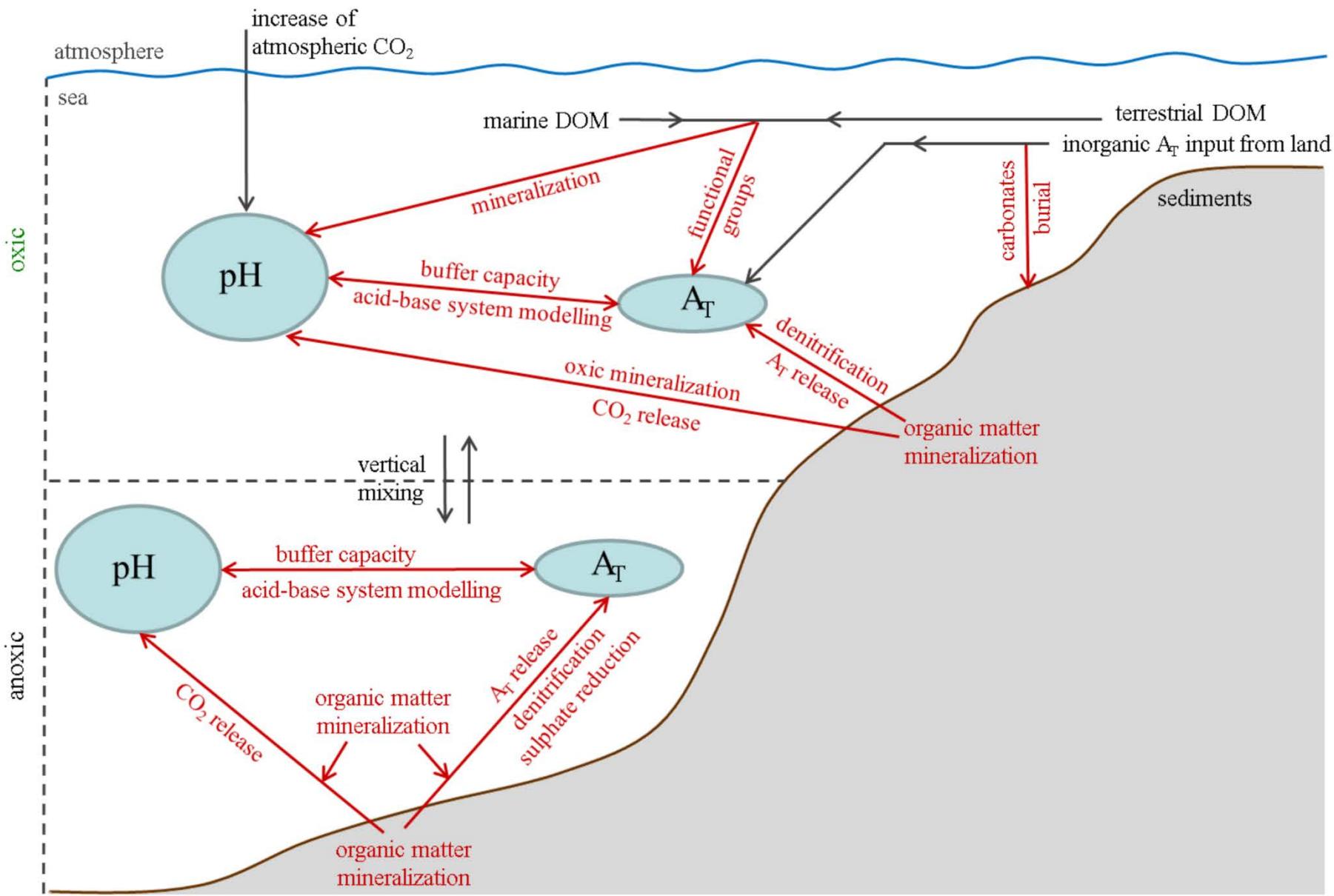


hypothetical structure of humic-like substances



Source: Hoikkala et al., 2015

Peculiarities of the acid-base system in the Baltic Sea



Question:

Name the nation we
all hate?

Answer:

Exami-Nation

